



The Community Land Model

Representing terrestrial processes in the Earth System

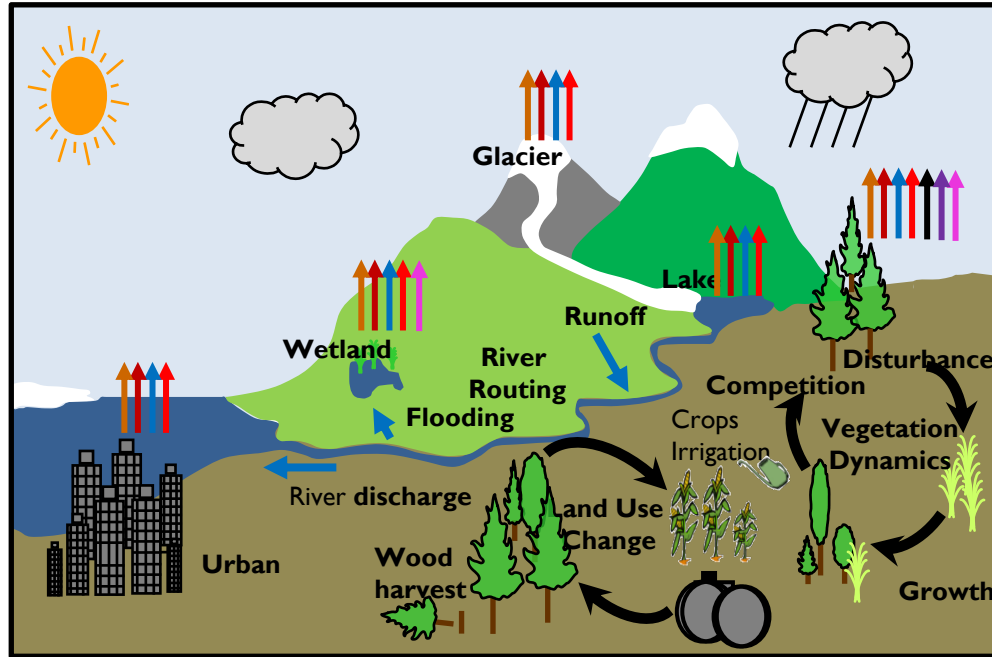
David Lawrence

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Climate and Global Dynamics Lab
Terrestrial Sciences Section
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NCAR is sponsored by the National Science Foundation



Land Modeling

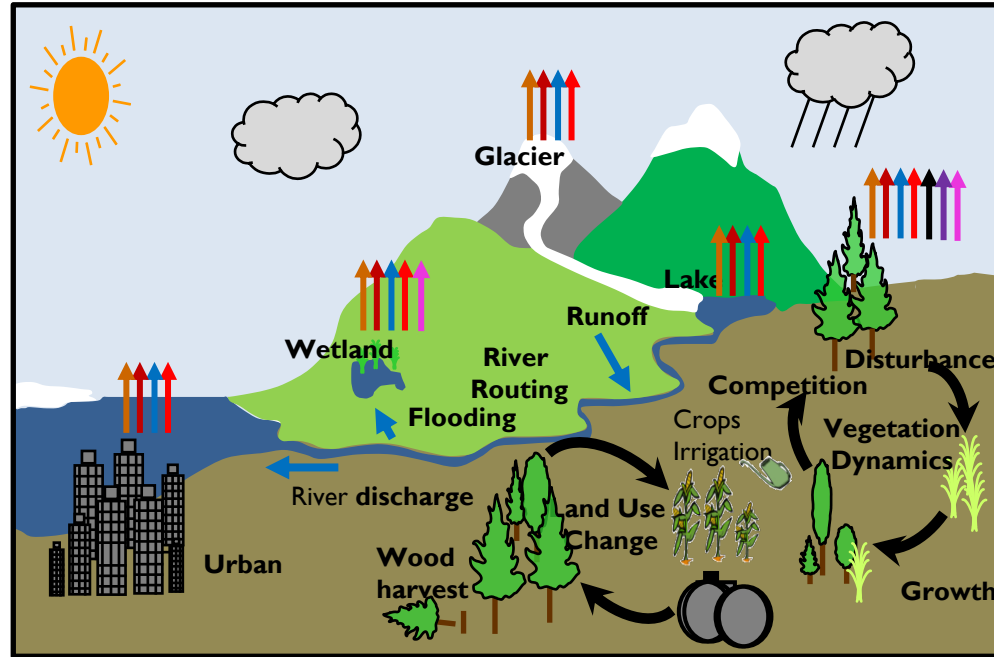


“Why?”

“Are you sure this is necessary?”



Land Modeling



Yes!

Land is the critical interface through which humanity affects and is affected by, adapts to, and mitigates global environmental change

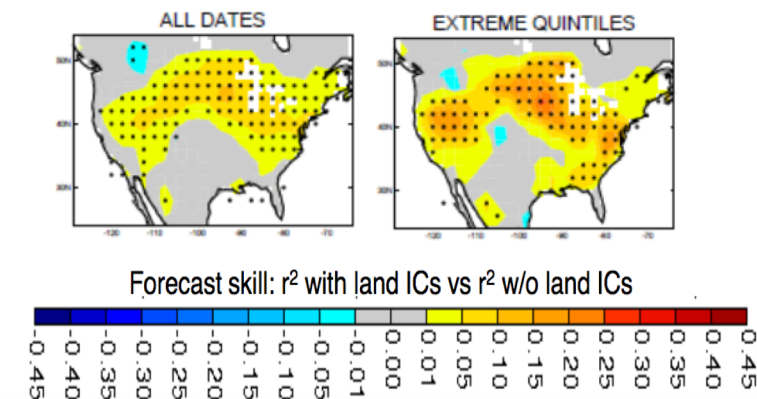
Land modelling, why?

Land-atmosphere interactions



- **When, where, and by how much do land fluxes influence atmosphere, surface temperature, clouds, precipitation, etc.?**
- **Land-driven predictability**
 - **Significant skill, especially when conditioned on amplitude of initial soil moisture anomaly**
 - **Increased land-atmosphere coupling in future warmer climate, increased land-driven skill?**
- **Land influence on extremes**

30-45 day forecast conditioned on SM



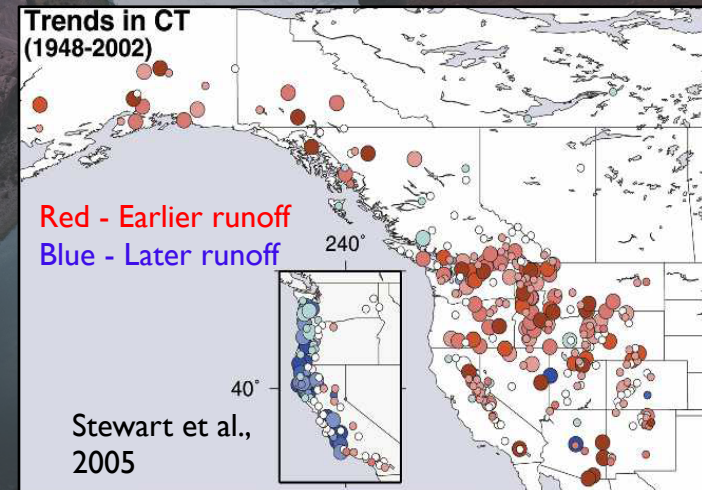
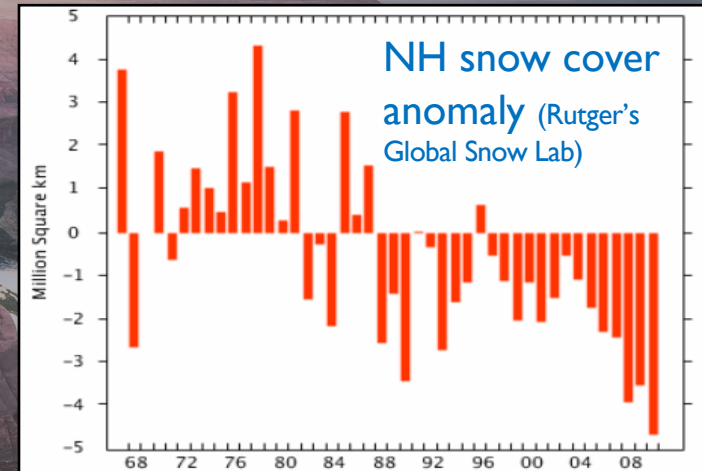
Koster et al., 2010



Land modeling, why?

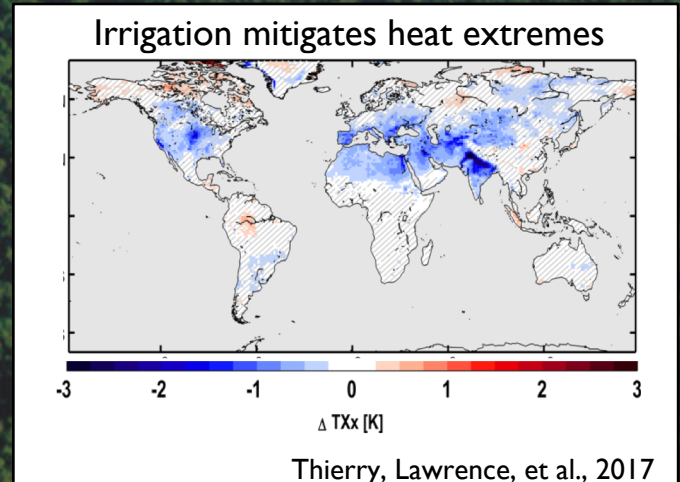
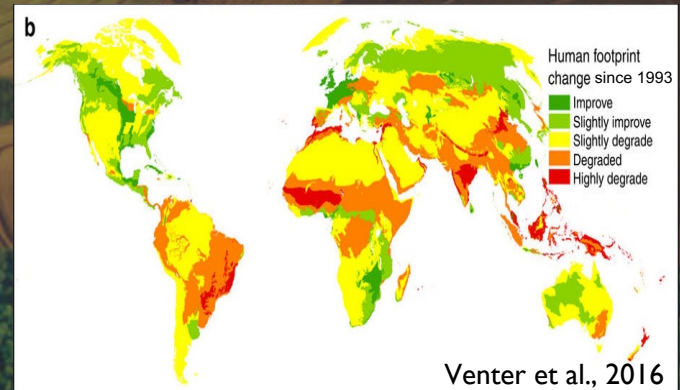
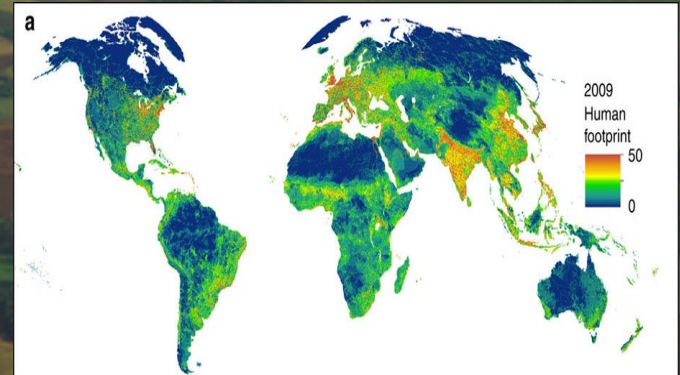
Water

- Land feedbacks on droughts and floods
- Snow-albedo and snow-soil T feedbacks
- Water and food security
 - >1/6th world population dependent on water from seasonal snowpacks
- Water – plant interactions
 - Plant water use efficiency likely to increase with CO₂
- Streamflow prediction



Land modeling why? Land-use and land-cover change

- ~25% non-ice land area undergone anthropogenic land-cover change
- ~80% non-ice land area under some form of land management
- Regionally, LULCC as impactful on surface climate as greenhouse gases
- ~1/3 of direct historic carbon emissions ($180 \pm 80 \text{ PgC}$ from land use, $\sim 400 \text{ PgC}$ from fossil fuel and cement),
- Deforestation: loss of Additional Sink Capacity yields indirect C impact
- Effectiveness of afforestation and biofuels for CO_2 mitigation
- Urban-rural differences in climate change impacts, e.g., heat stress



Land modeling, why?

Carbon and ecology

- Carbon and nitrogen cycle interactions and their impact on long term trajectory of terrestrial carbon sink
- High uncertainty in projected land C sink
 - Emissions driven RCP8.5: 795 to 1140 ppm (source of $\pm 1.2^{\circ}\text{C}$ uncertainty on top of 3.7°C projected change)
- Vulnerability of ecosystems to climate change as well as natural and human disturbances
- Ecosystem services
- Ecosystem management to mitigate climate change



The interdisciplinary evolution of land models



The interdisciplinary evolution of land models

Land as a lower boundary
to the atmosphere



Land as an integral component
of the Earth System

Surface Energy Fluxes

70's

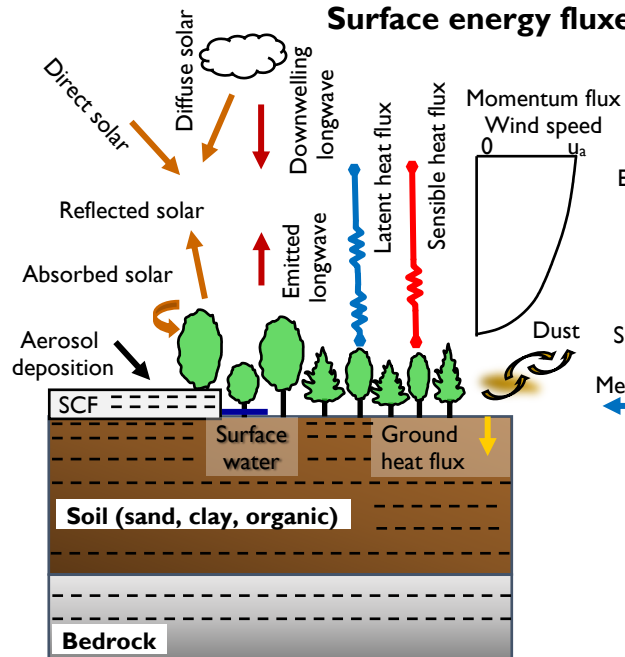
80's

90's

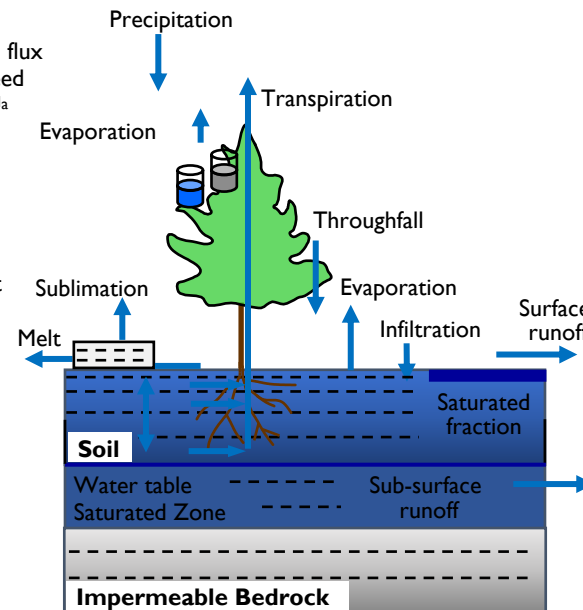
00's

10's

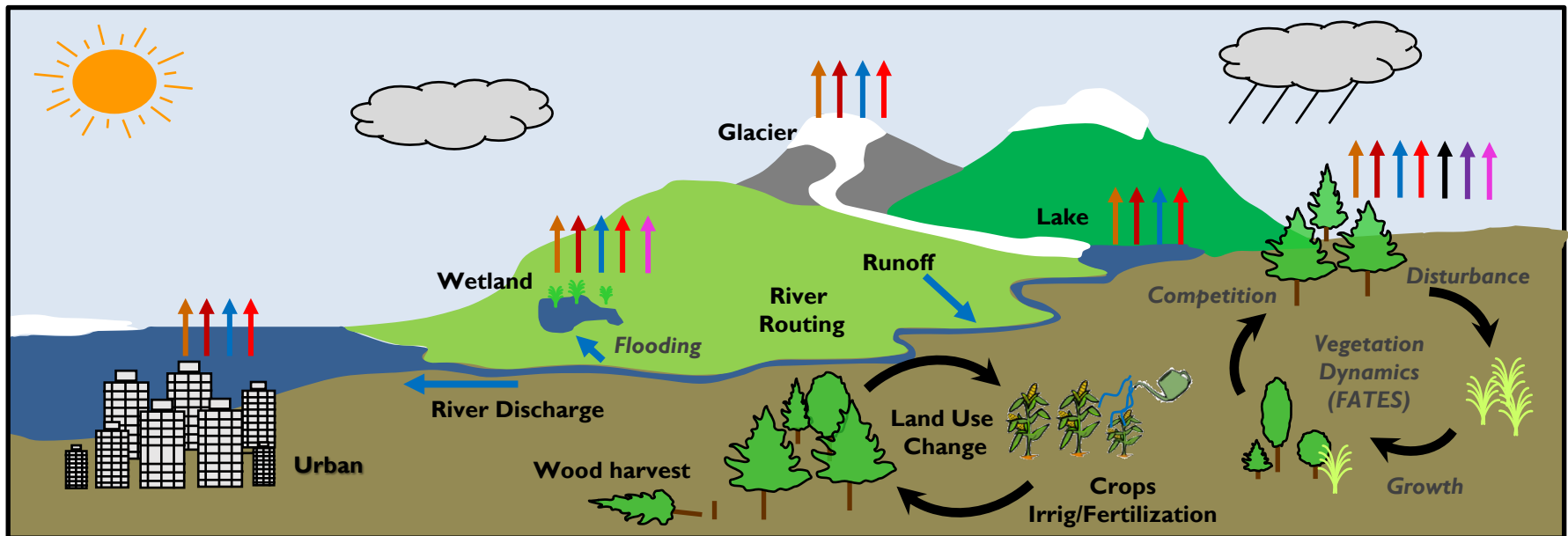
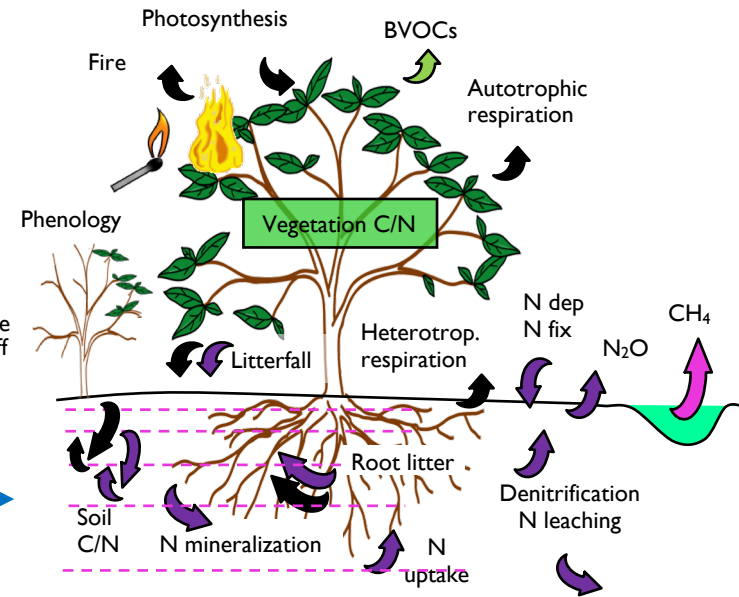
Surface energy fluxes

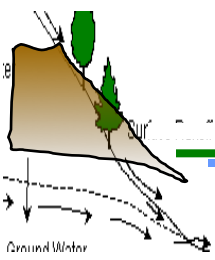


Hydrology



Biogeochemical cycles

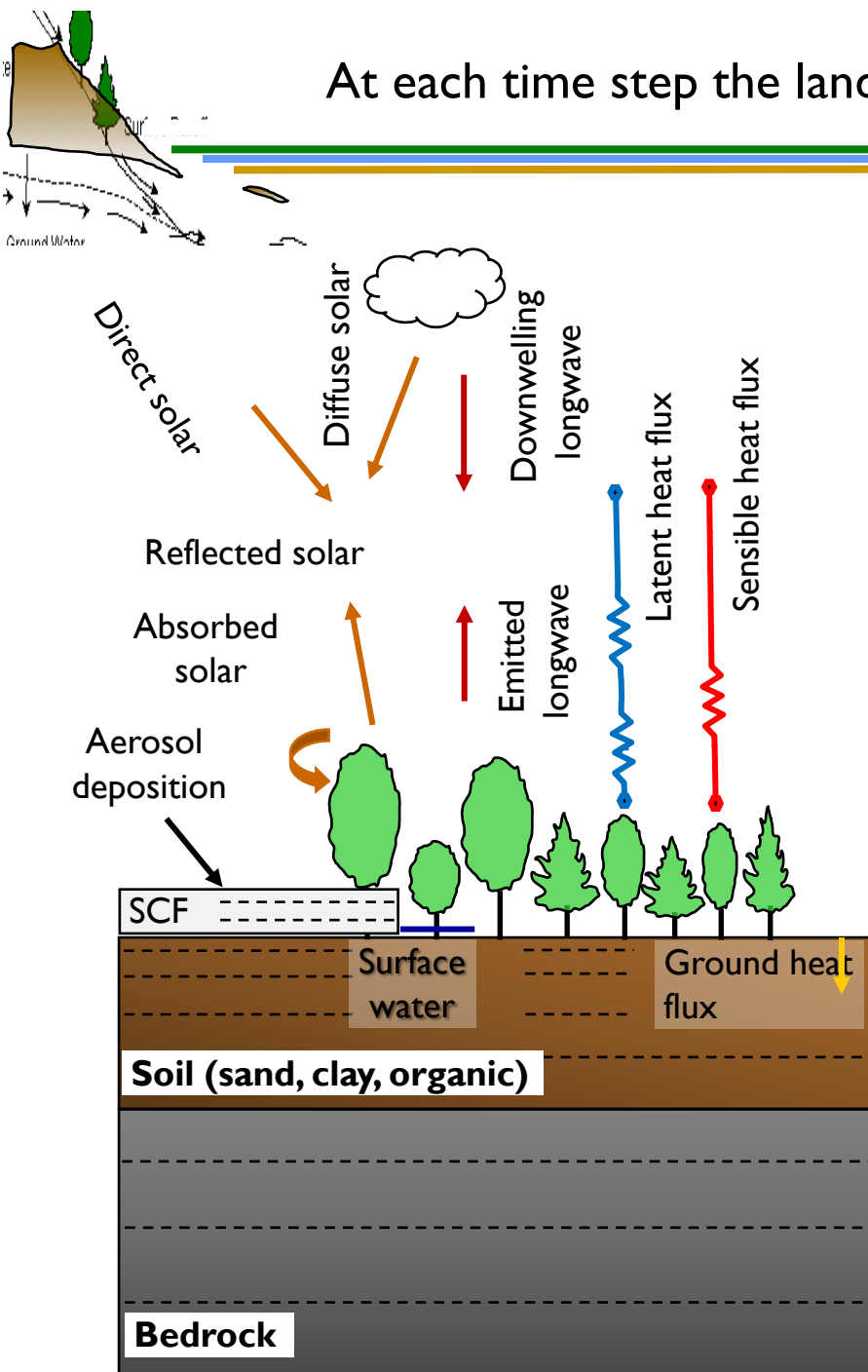




The role of a land model within an Earth System Model

- exchanges of momentum, energy, water vapor, CO_2 , dust, and other trace gases/materials between land surface and the overlying atmosphere (and routing of runoff to the ocean)
- states of land surface (e.g., soil moisture, soil temperature, canopy temperature, snow water equivalent, C and N stocks in vegetation and soil)
- characteristics of land surface (e.g., soil texture, surface roughness, albedo, emissivity, vegetation type, cover extent, leaf area index, and seasonality)

At each time step the land model solves Surface Energy Balance



$$S^{\uparrow} - S^{\downarrow} + L^{\uparrow} - L^{\downarrow} = \lambda E + H + G$$

S^{\uparrow} , S^{\downarrow} are down(up)welling solar radiation,

L^{\uparrow} , L^{\downarrow} are up(down)welling longwave rad,

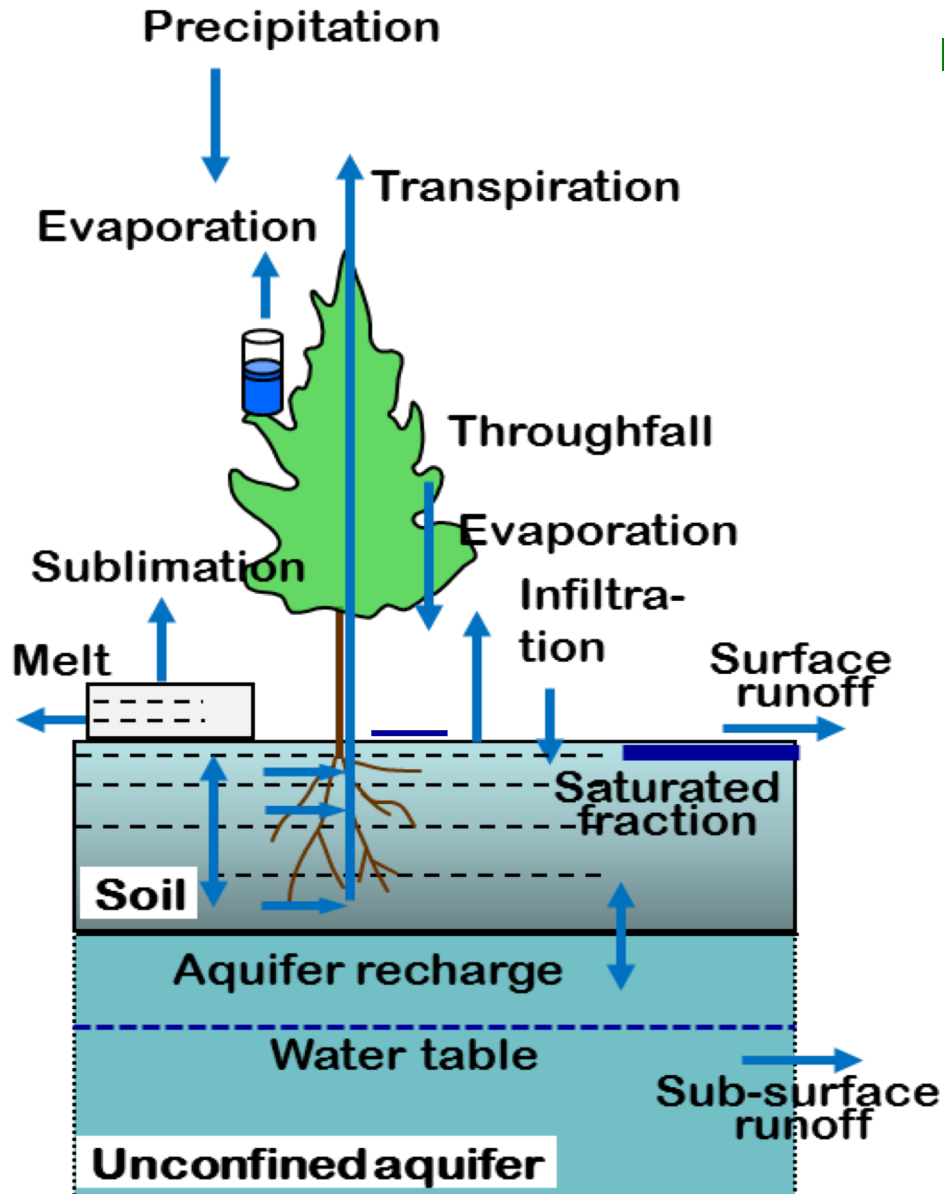
λ is latent heat of vaporization,

E is evaporation,

H is sensible heat flux

G is ground heat flux

... and the Surface Water Balance



$$P = E_S + E_T + E_C + R +$$

$$(\Delta W_{\text{soi}} + \Delta W_{\text{snw}} + \Delta W_{\text{sfcw}} + \Delta W_{\text{can}}) / \Delta t$$

P is rainfall/snowfall,

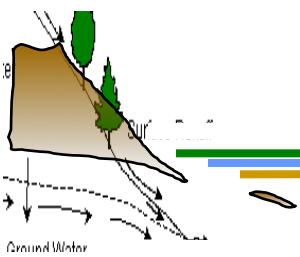
E_S is soil evaporation,

E_T is transpiration,

E_C is canopy evaporation,

R is runoff (surf + sub-surface),

$\Delta W_{\text{soi}} / \Delta t$, $\Delta W_{\text{snw}} / \Delta t$, $\Delta W_{\text{sfcw}} / \Delta t$, $\Delta W_{\text{can}} / \Delta t$,
are the changes in soil moisture, surface
water, snow, and canopy water over a
timestep

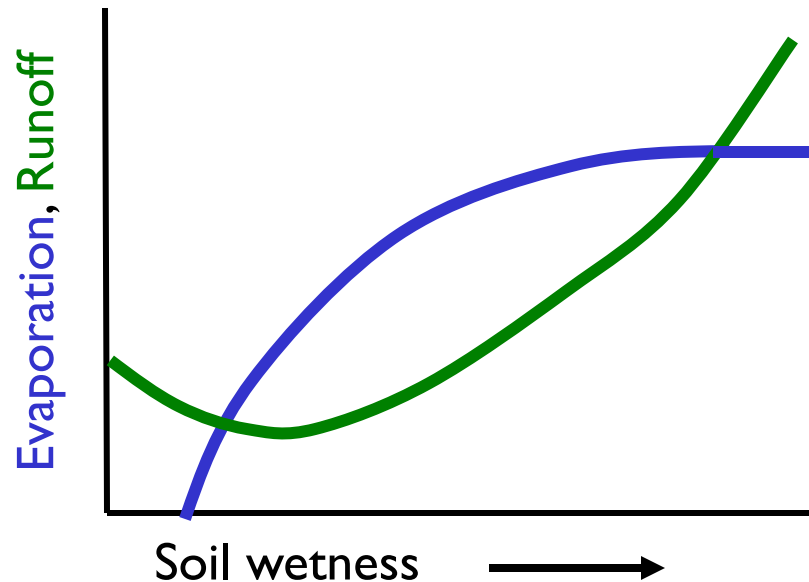


Terrestrial water and energy cycles intricately linked

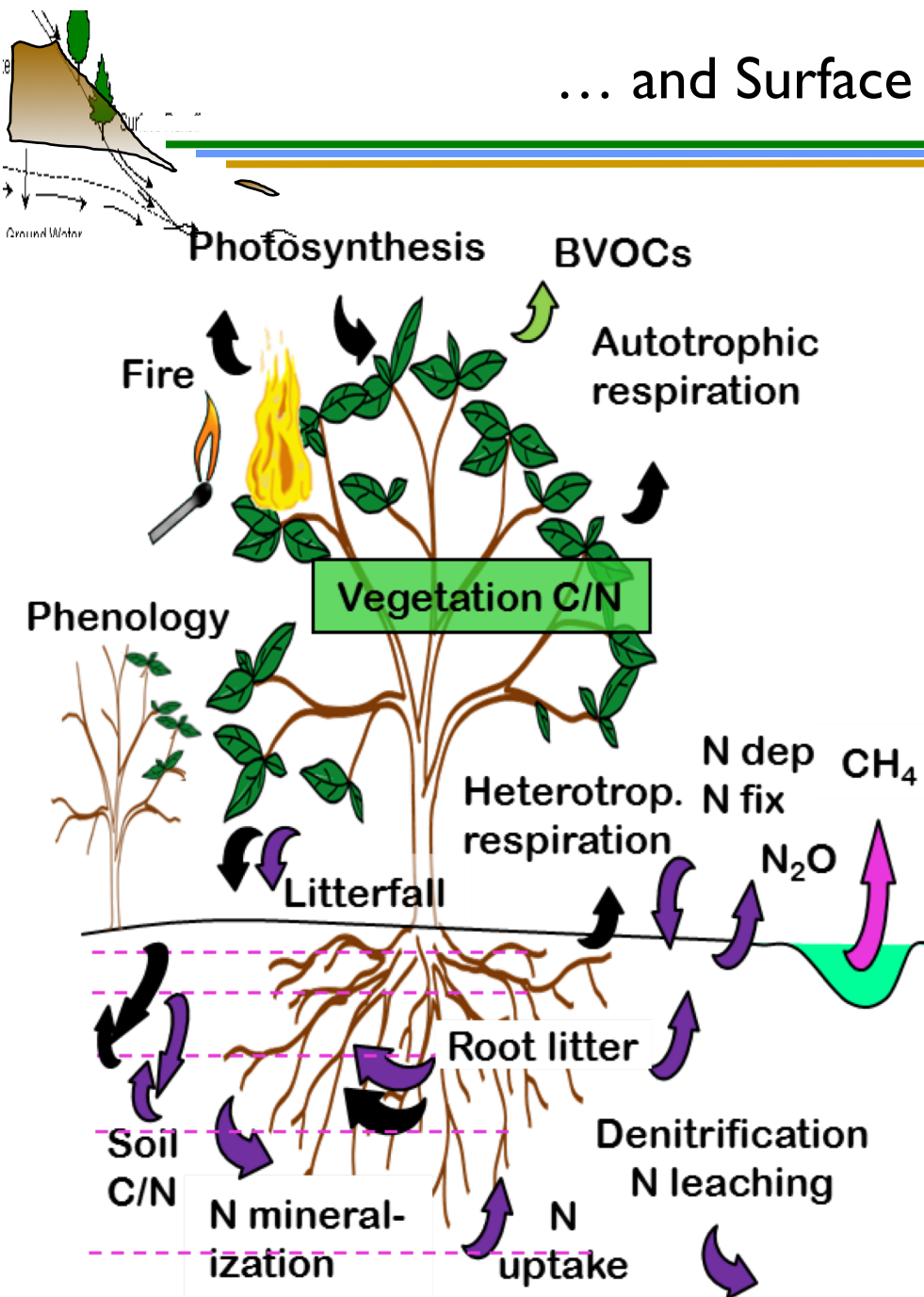
“The ability of a land-surface scheme to model evaporation correctly depends crucially on its ability to model runoff correctly. The two fluxes are intricately related through soil moisture.”

(Koster and Milly, 1997).

Runoff and evaporation both vary non-linearly with soil moisture



... and Surface Carbon Exchange



$$\text{NEE} = \text{GPP} - \text{HR} - \text{AR} - \text{Fire} - \text{LUC}$$

NEE is net ecosystem exchange

GPP is gross primary productivity

HR is heterotrophic respiration

AR is autotrophic respiration

Fire is carbon flux due to fire

LUC is C flux due to land use change



Land complexity: Submodels of CLM

– Biogeophysics

- Photosynthesis and stomatal resistance
- Hydrology
- Snow
- Soil thermodynamics
- Surface albedo and radiative fluxes

– Biogeochemistry

- Carbon / nitrogen pools, allocation, respiration
- Vegetation phenology
- Decomposition
- Plant mortality
- External nitrogen cycle
- Methane production and emission

– Vegetation dynamics

– Urban

– Crop and irrigation

– Lakes

– Glaciers and ice sheets

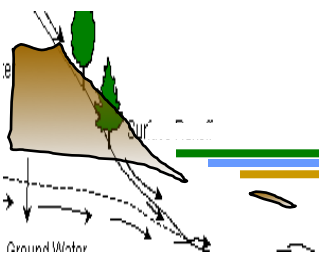
– Fire and fire emissions

– Dust emissions

– River flow

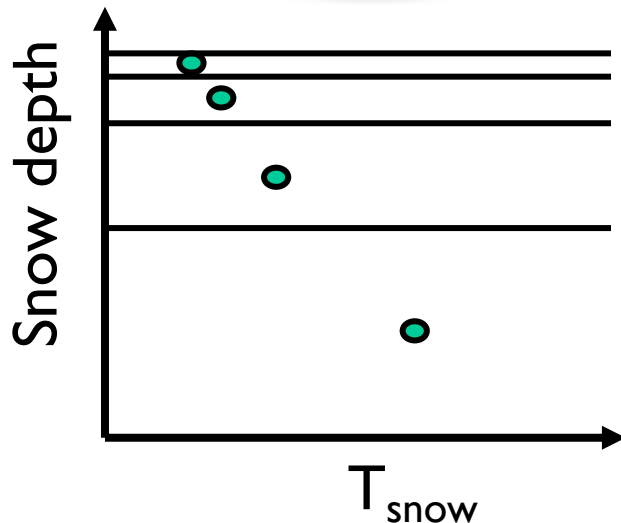
– Biogenic Volatile Organic Compound emissions

Land model complexity: Snow model example



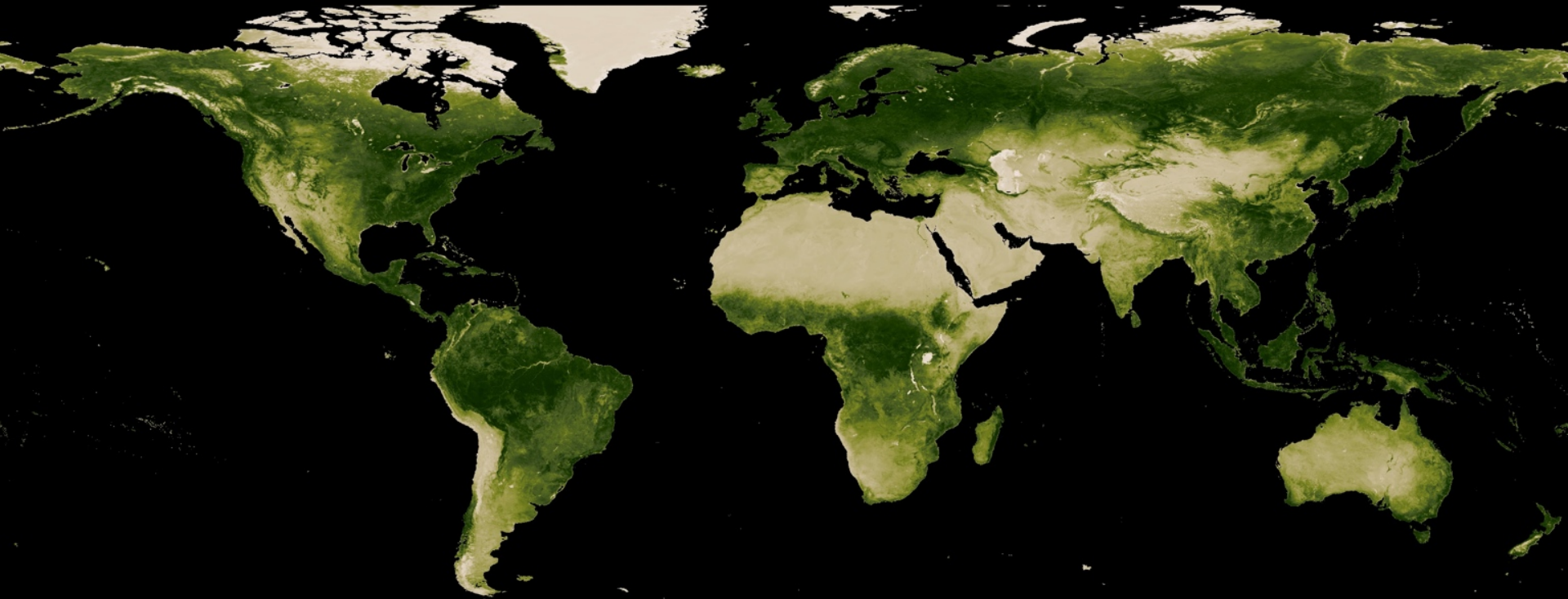
State Variables

$$N, w_{liq,i}, w_{ice,i}, \Delta z_i, T_i$$

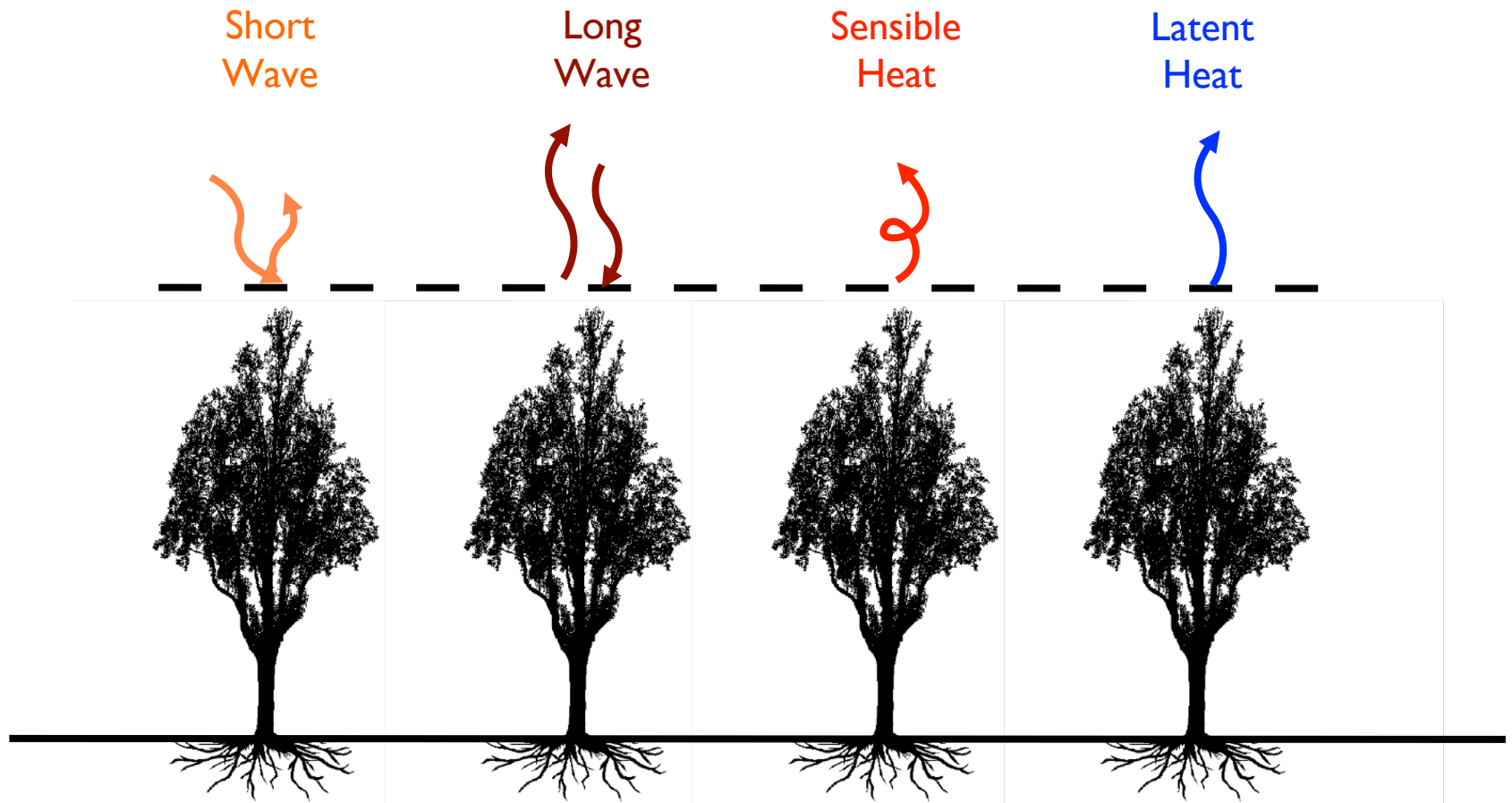


- Up to 10-layers of varying thickness
- Represented processes
 - Accumulation and fresh snow density $f(T, \text{wind})$
 - Snow melt and refreezing
 - Snow aging
 - Water and energy transfer across snow layers
 - Snow compaction
 - destructive metamorphism due to temperature and wind
 - overburden
 - melt-freeze cycles
 - Sublimation
 - Aerosol (black carbon, dust) deposition
 - Canopy snow storage and unloading
 - Canopy snow radiation
 - Snow burial of vegetation
 - Snow cover fraction
- Missing processes
 - Blowing snow
 - Subgrid variations in snow depths
 - Depth hoar

Plants ↔ Climate



How do Plants and Climate Interact?



Terrestrial Surface Energy Budget

How do Plants and Climate Interact?

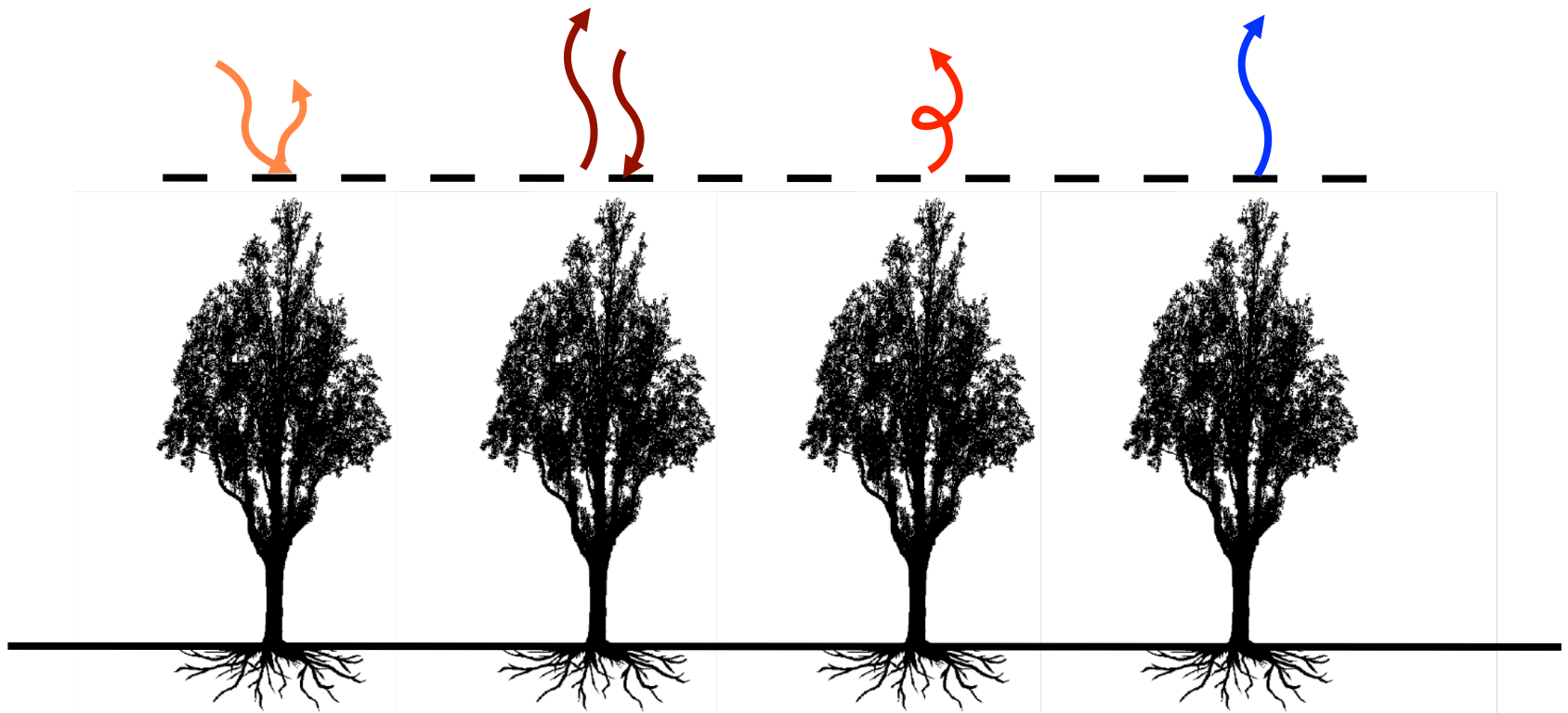
Albedo

Sunlight

Long
Wave

Sensible
Heat

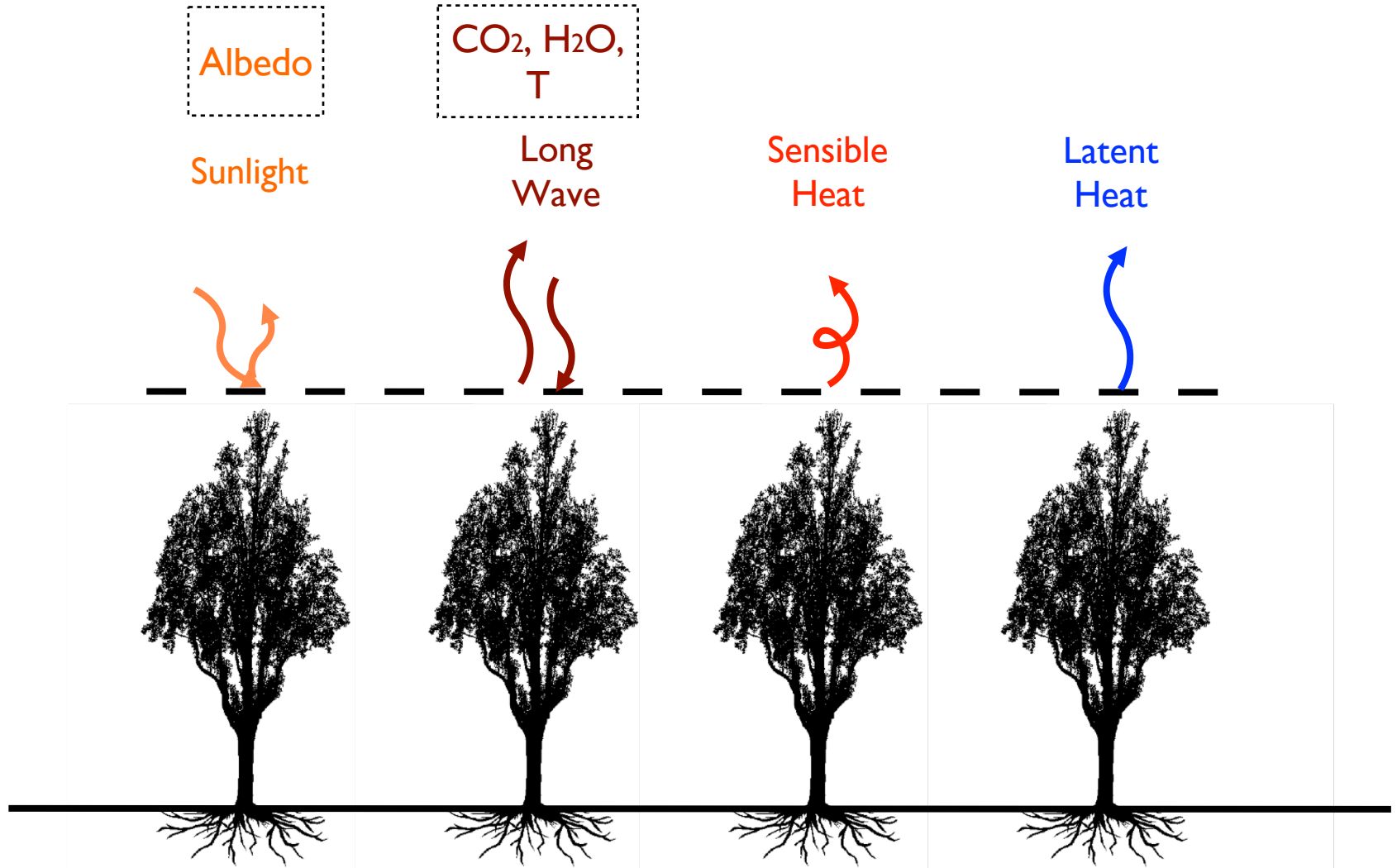
Latent
Heat



A scenic landscape photograph showing a dense forest of evergreen trees in the foreground and middle ground. In the background, rolling mountain ranges are visible under a clear blue sky with a few wispy clouds. The foreground is framed by the branches and green leaves of trees, some of which are in sharp focus while others are blurred. The overall scene is bright and sunny, with a mix of green and brown tones from the vegetation and the blue of the sky and distant mountains.

Albedo varies by plant type

How do Plants and Climate Interact?



How do Plants and Climate Interact?

Albedo

Sunlight



CO₂, H₂O,
T

Long
Wave

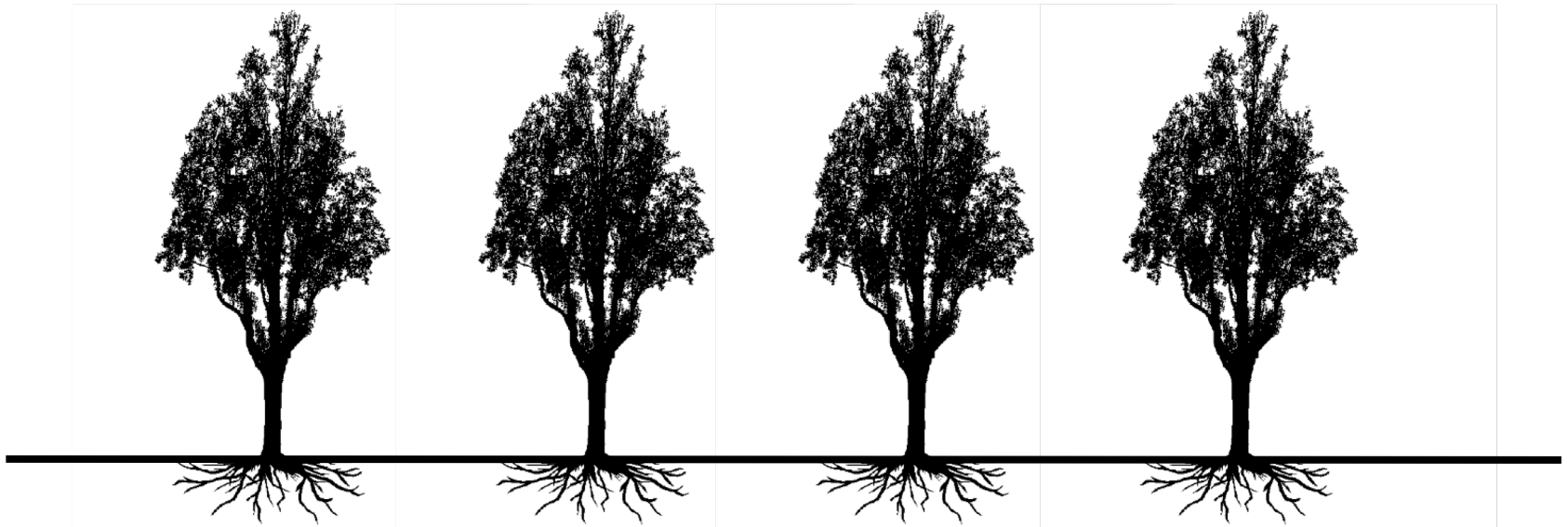


Roughness

Sensible
Heat



Latent
Heat



How do Plants and Climate Interact?

Albedo

Sunlight



CO₂, H₂O,
T

Long
Wave

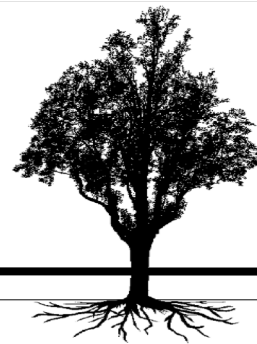
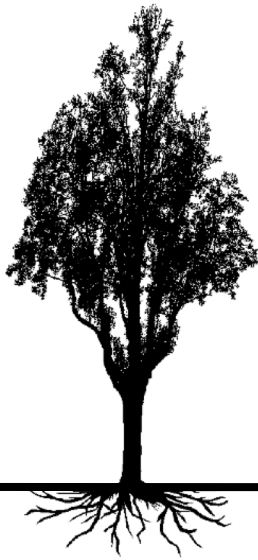


Roughness

Sensible
Heat



Latent
Heat



How do Plants and Climate Interact?

Albedo

Sunlight



CO₂, H₂O,
T

Long
Wave



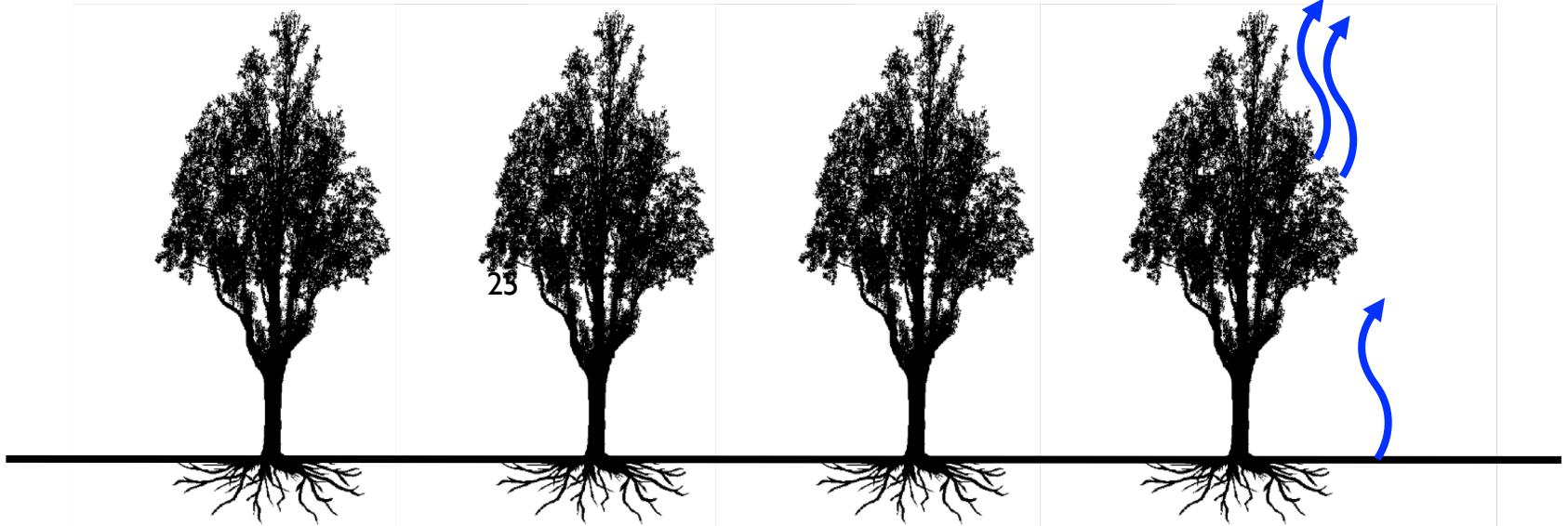
Roughness

Sensible
Heat

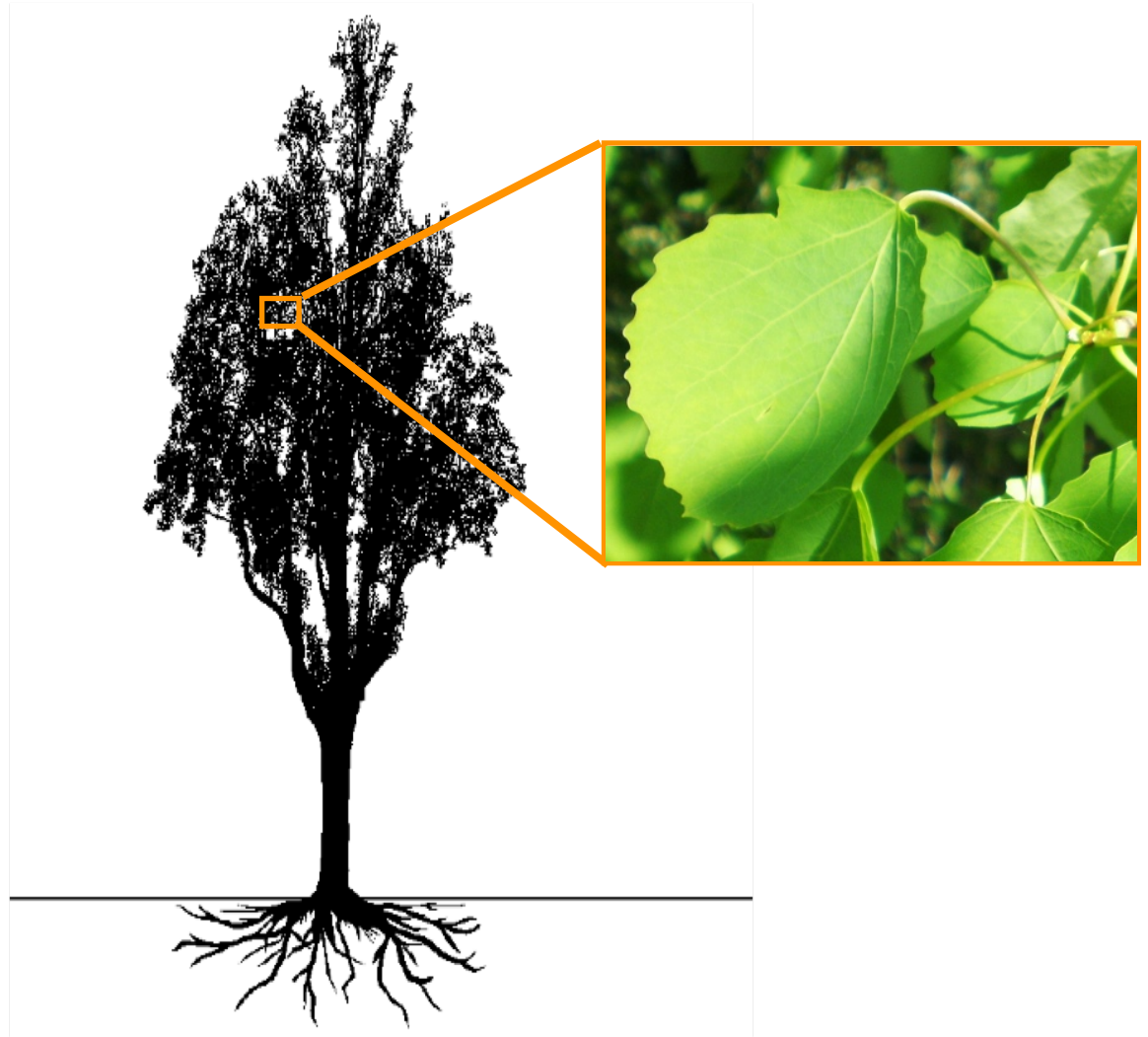


Evaporation,
Transpiration

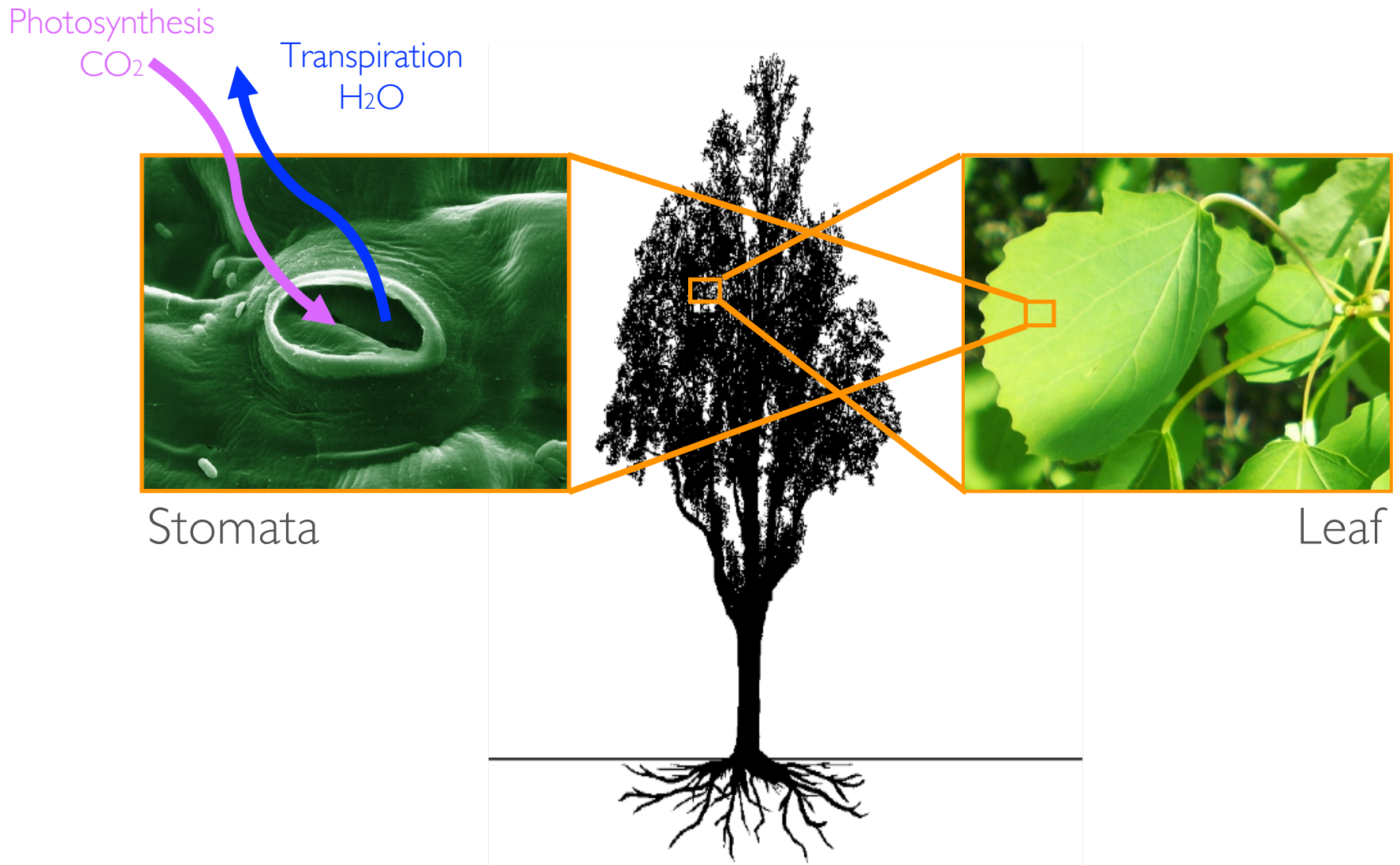
Latent
Heat



Transpiration flux of water



Carbon in, water out



Plant physiological controls on CO₂ exchange and transpiration

Function of solar radiation, humidity deficit, soil moisture, [CO₂], temperature, leaf N content

Photos: Wikimedia Commons

Δ Plants \Rightarrow Δ Surface Energy Budget

Albedo

Sunlight



CO₂, H₂O,
T

Long
Wave



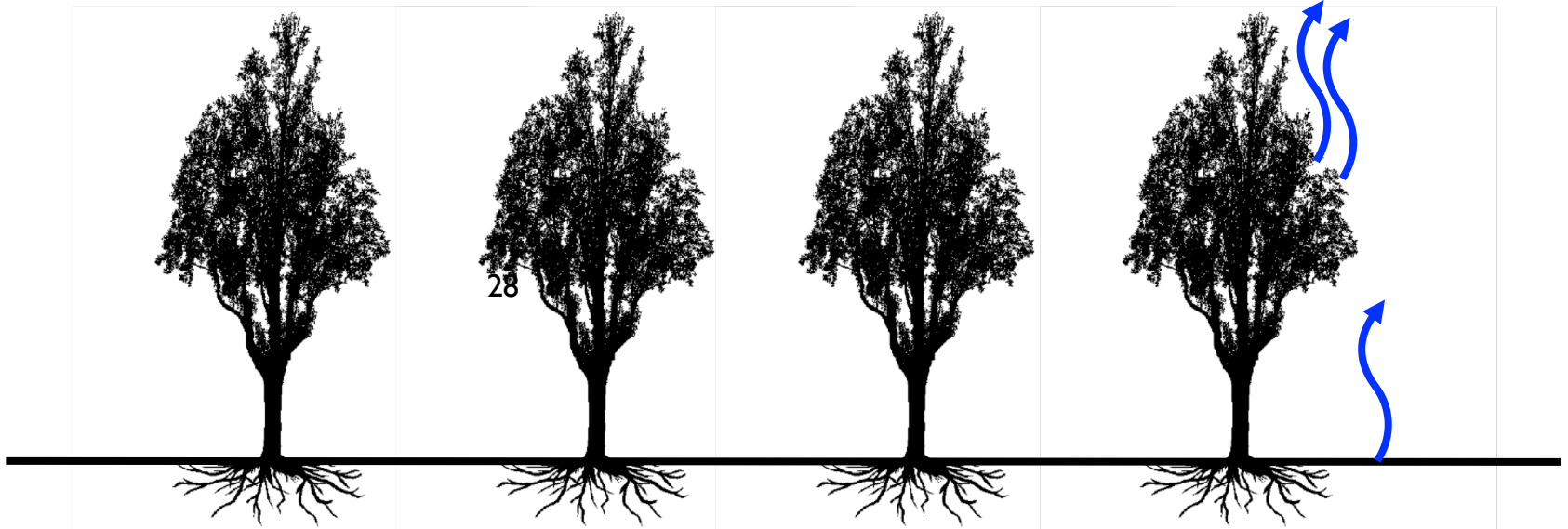
Roughness

Sensible
Heat

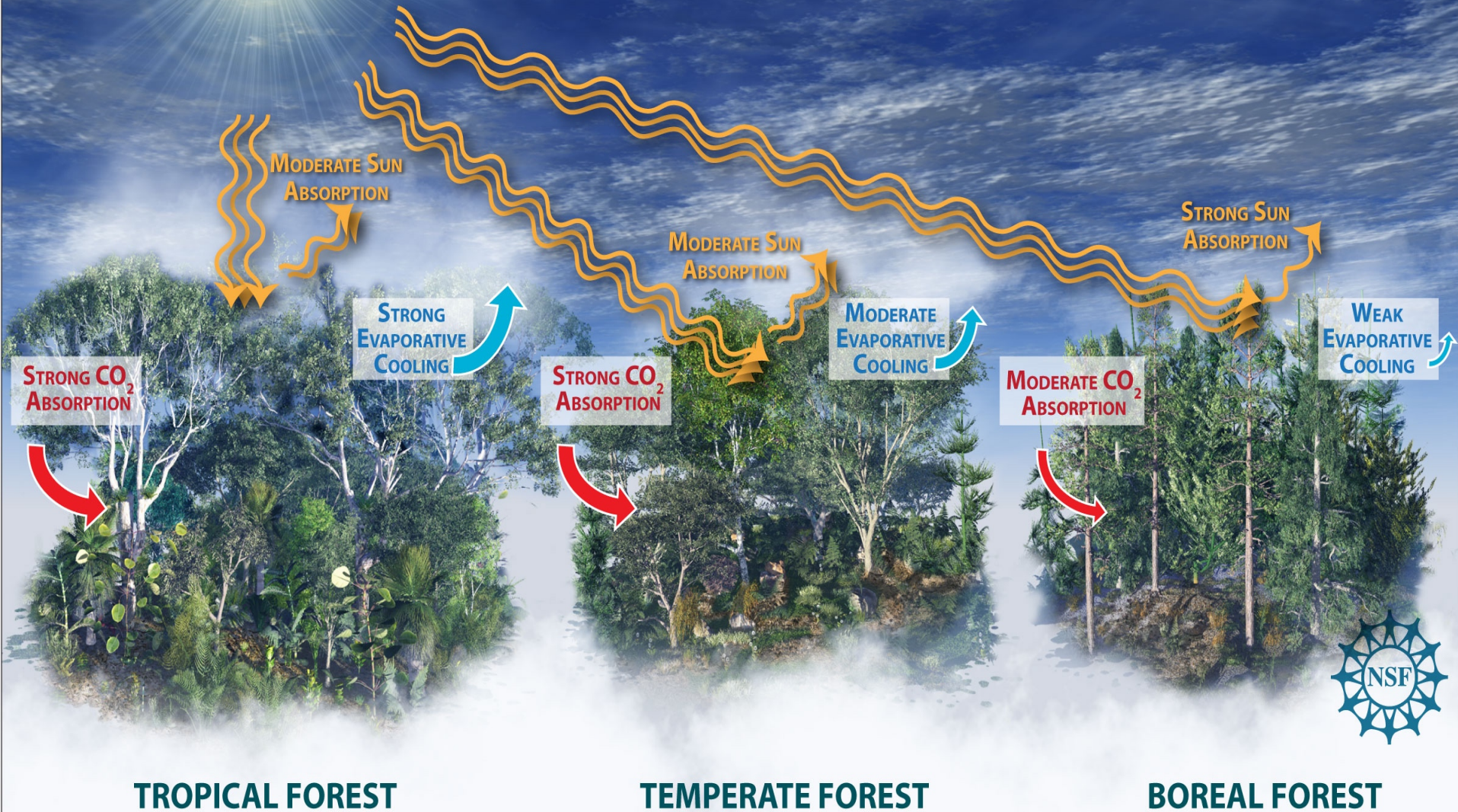


Evaporation,
Transpiration

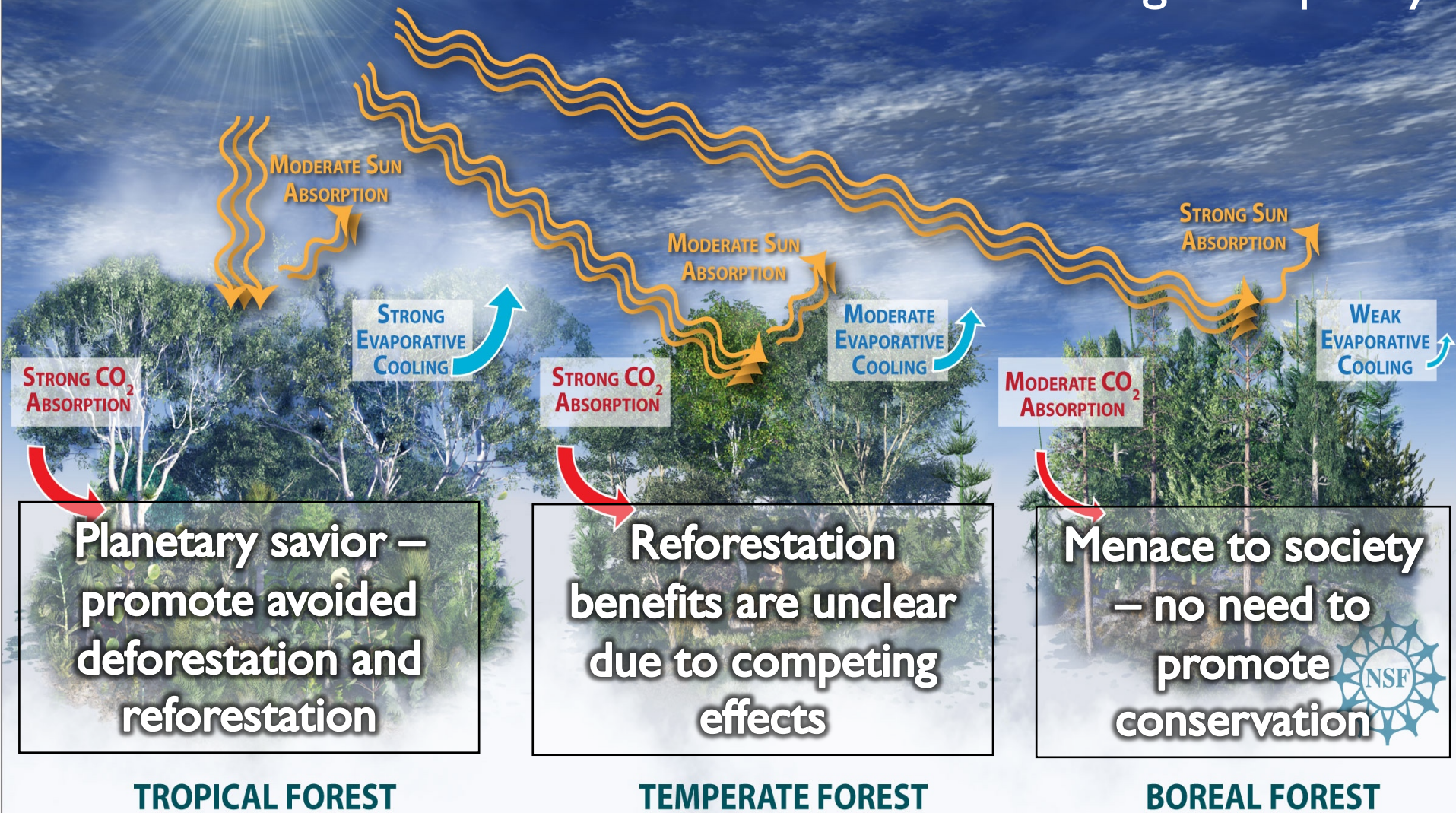
Latent
Heat



Not all forest ecosystems have the same impact on climate



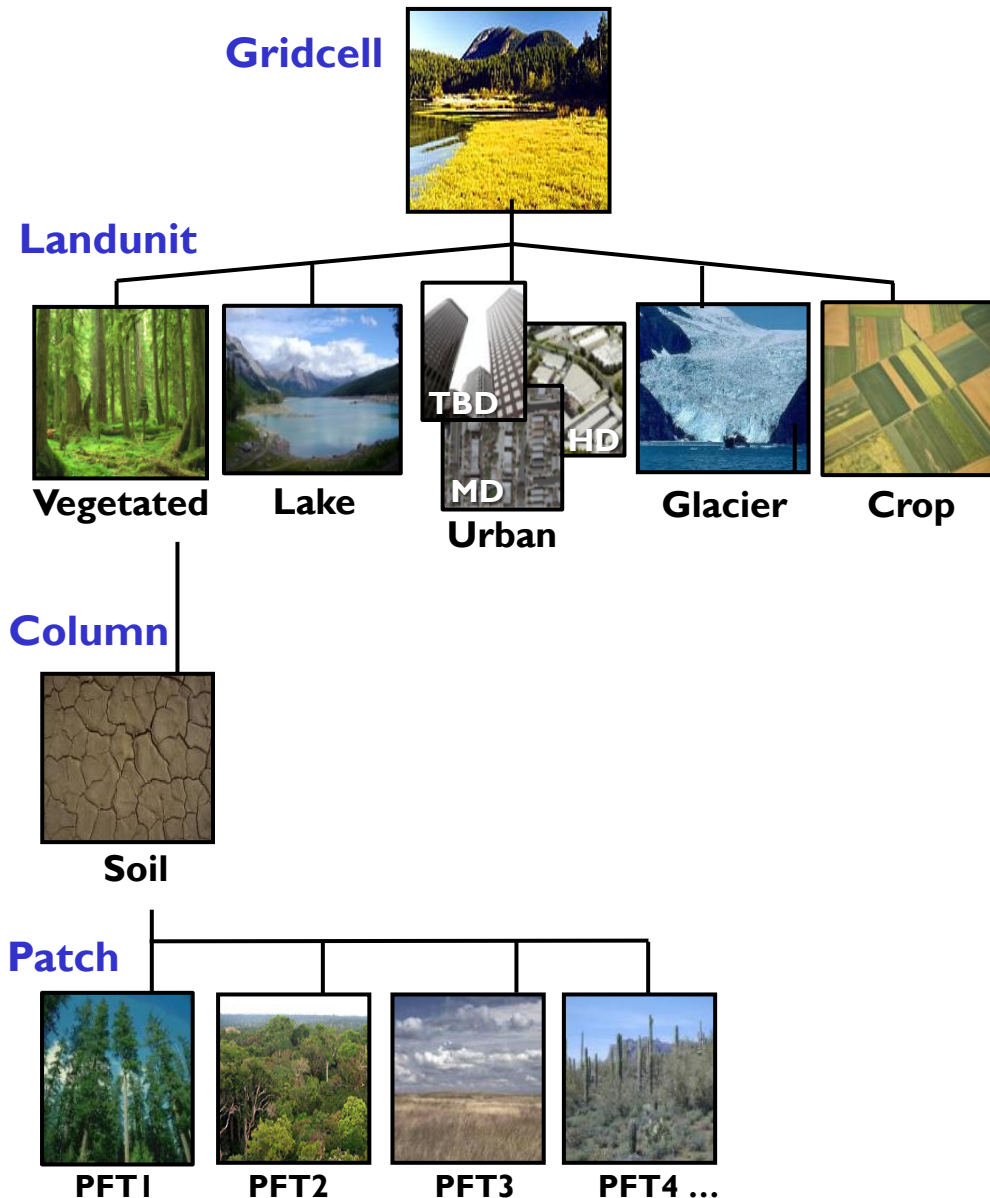
Differences in ecosystem functioning have implications for land climate mitigation policy



Land Modeling Challenges: Land surface heterogeneity



Land surface heterogeneity: Subgrid tiling



Plant Functional Types:

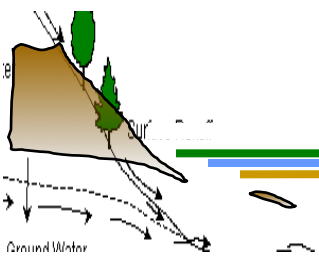
0. Bare

Tree:

1. Needleleaf Evergreen, Temperate
2. Needleleaf Evergreen, Boreal
3. Needleleaf Deciduous, Boreal
4. Broadleaf Evergreen, Tropical
5. Broadleaf Evergreen, Temperate
6. Broadleaf Deciduous, Tropical
7. Broadleaf Deciduous, Temperate
8. Broadleaf Deciduous, Boreal

Herbaceous / Understorey:

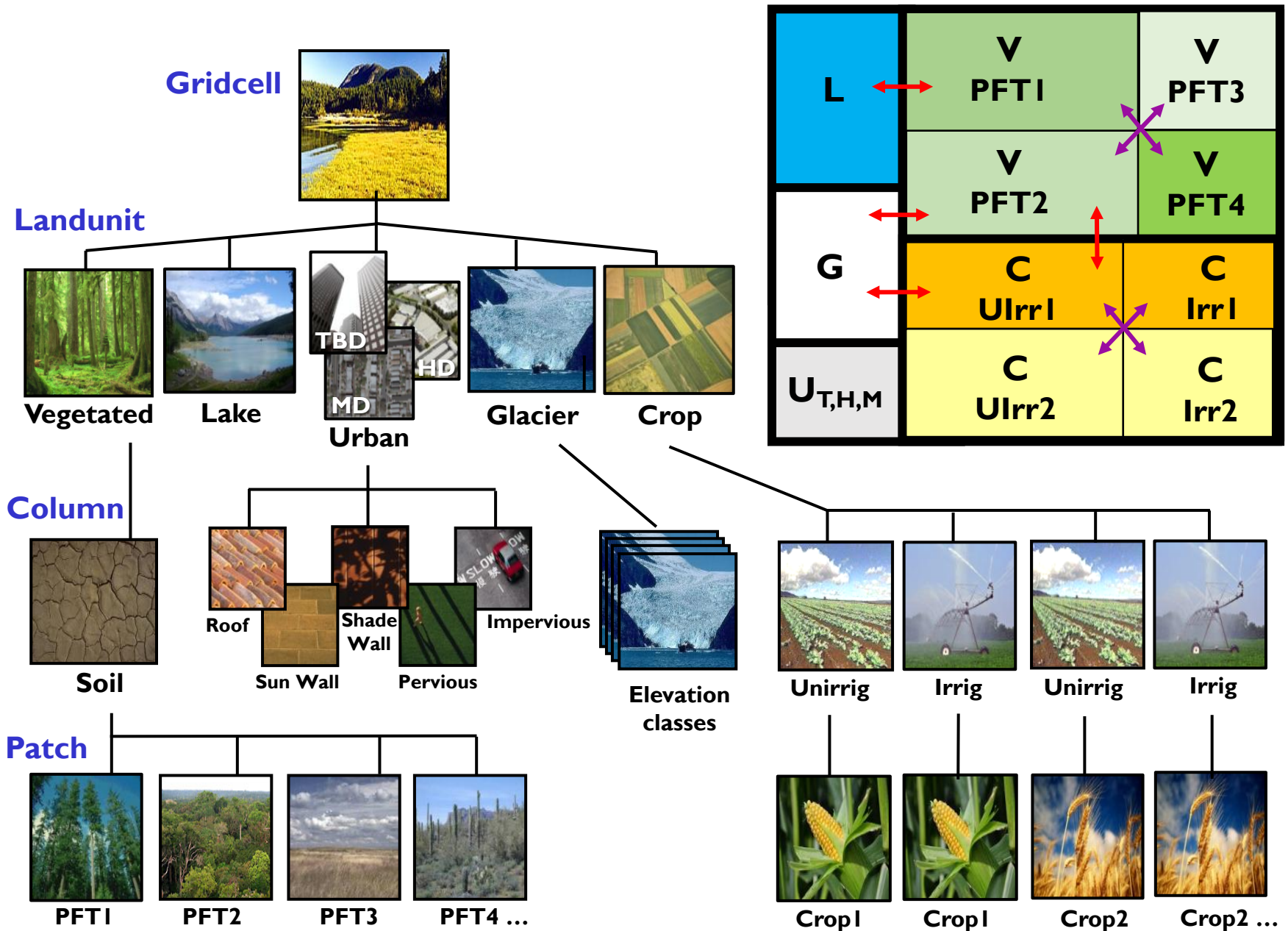
9. Broadleaf Evergreen Shrub, Temperate
10. Broadleaf Deciduous Shrub, Temperate
11. Broadleaf Deciduous Shrub, Boreal
12. C3 Arctic Grass
13. C3 non-Arctic Grass
14. C4 Grass
15. Crop

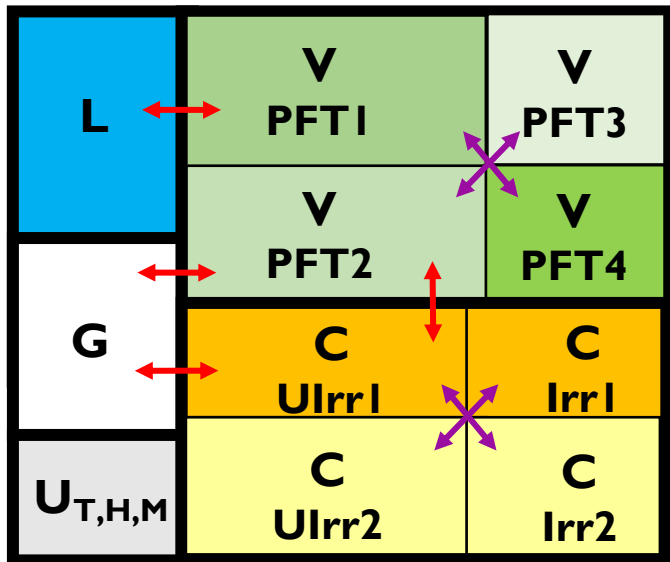
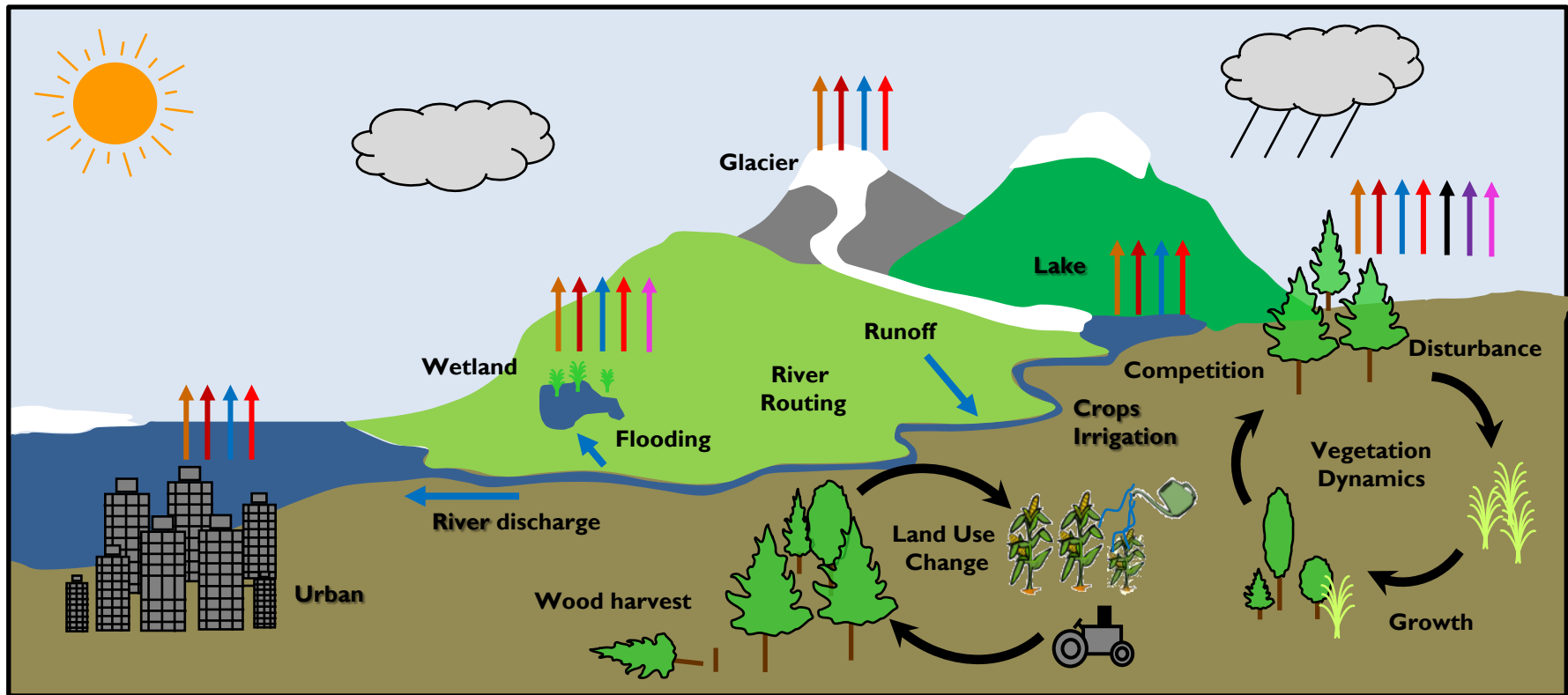


Plant Functional Type Parameters

- Optical properties (visible and near-infrared):
 - Leaf angle
 - Leaf reflectance
 - Stem reflectance
 - Leaf transmittance
 - Stem transmittance
- Fire:
 - Combustion completeness
 - Fire mortality
- Land models are parameter heavy!!!
- Morphological properties:
 - Leaf area index (annual cycle)
 - Stem area index (annual cycle)
 - Leaf dimension, leaf orientation
 - Roughness length/displacement height
 - Canopy top and bottom height
 - Root depth and distribution
- Photosynthetic parameters:
 - Specific leaf area
 - m (slope of conductance-photosynthesis relationship)
 - V_{cmax} (maximum rate of carboxylation)
 - Leaf carbon to nitrogen ratio
 - Fraction of leaf nitrogen in Rubisco
 - Root conductivity, plant conductivity

Land surface heterogeneity: Subgrid tiling

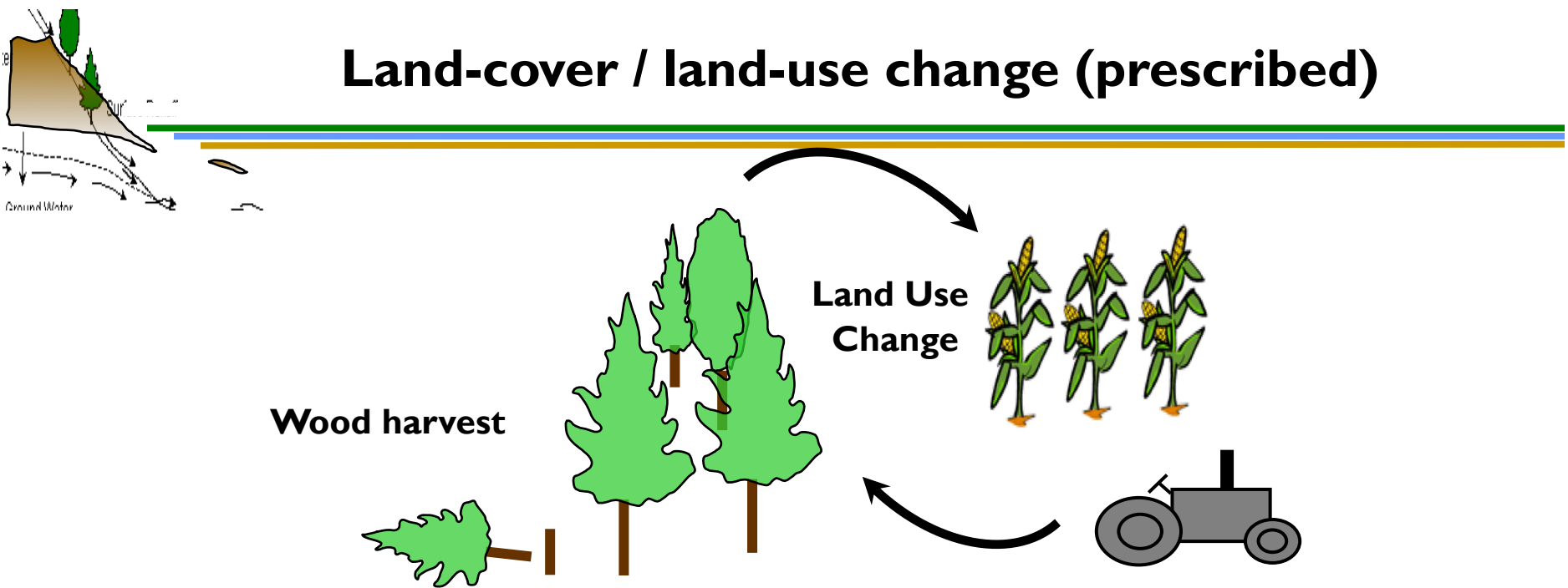




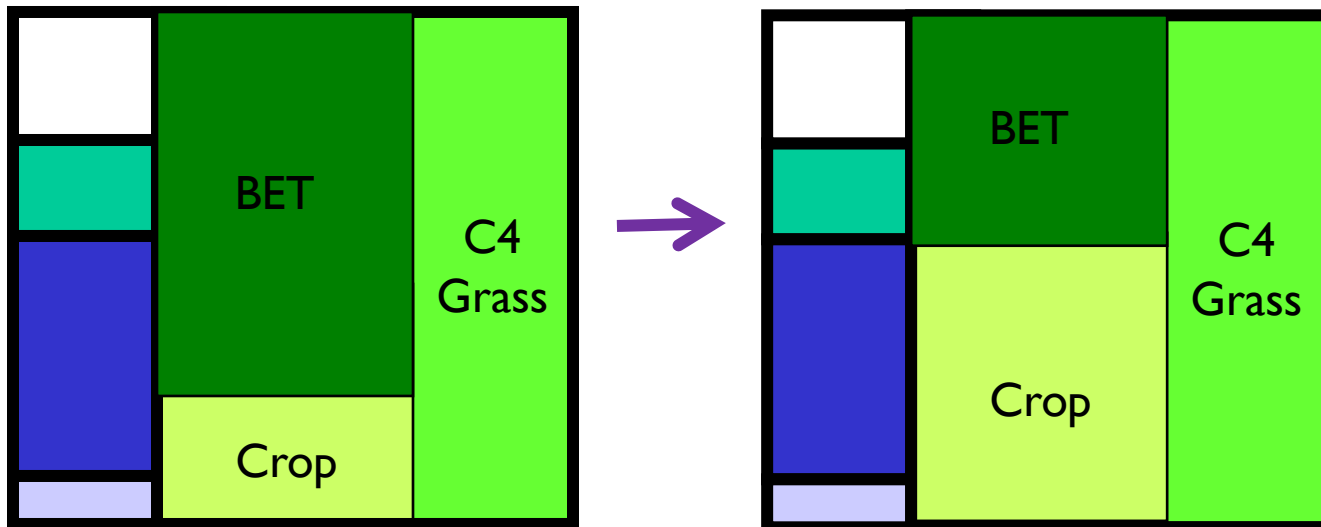
Landscape-scale dynamics

Long-term dynamical processes that affect fluxes in a changing environment (disturbance, land use, succession)

Land-cover / land-use change (prescribed)



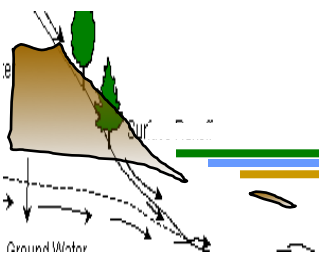
Deforestation example



Land Modeling Challenges: Land surface heterogeneity



Parameterize impact of subgrid-scale soil moisture heterogeneity

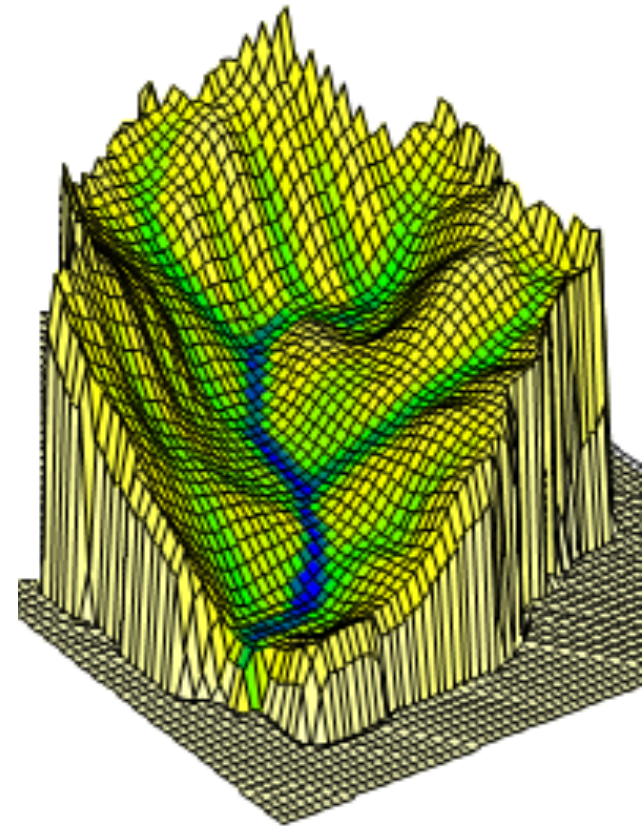


A major control on soil moisture heterogeneity and thus runoff is topography.

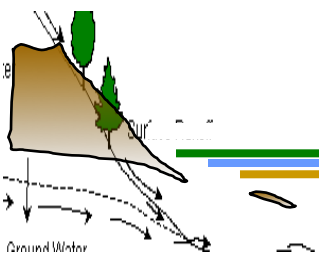
Lowland soils tend to be zones of high soil moisture content, while upland soils tend to be progressively drier.

Three main sources of runoff:

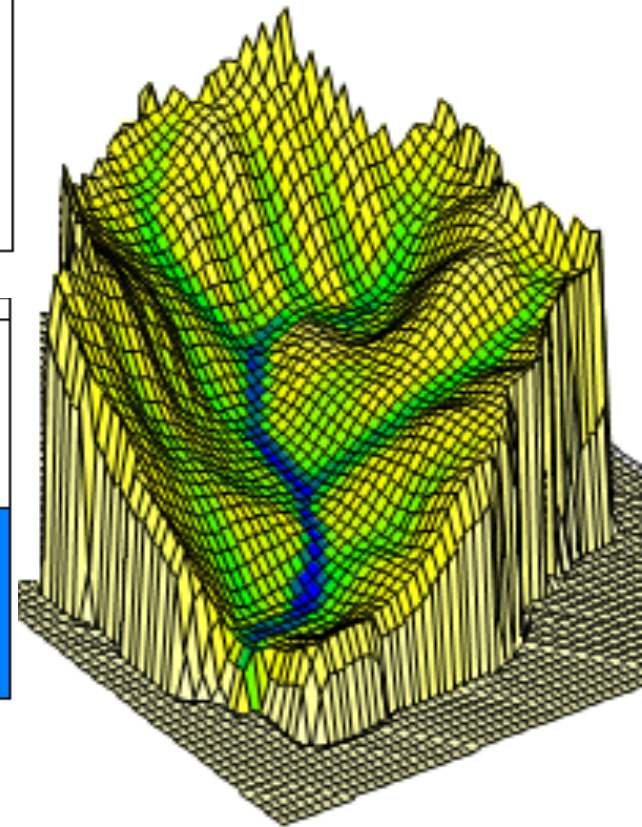
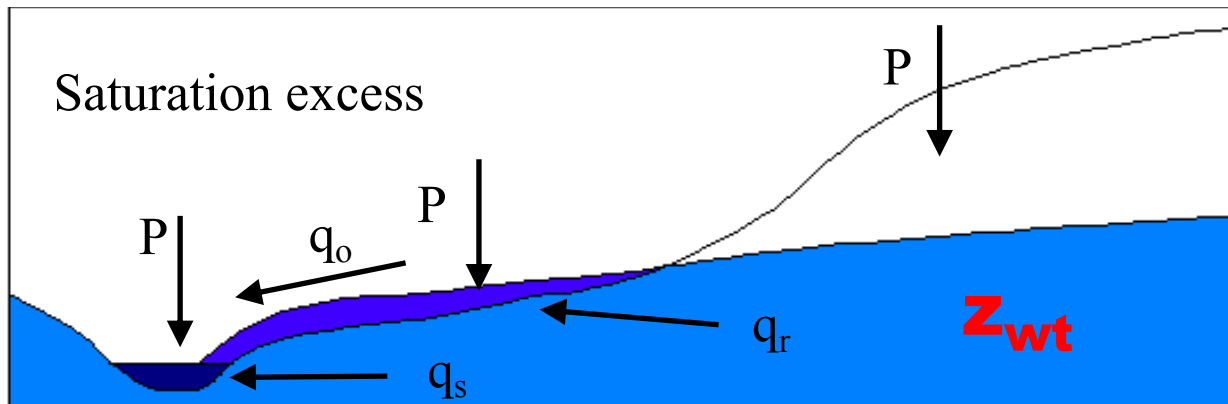
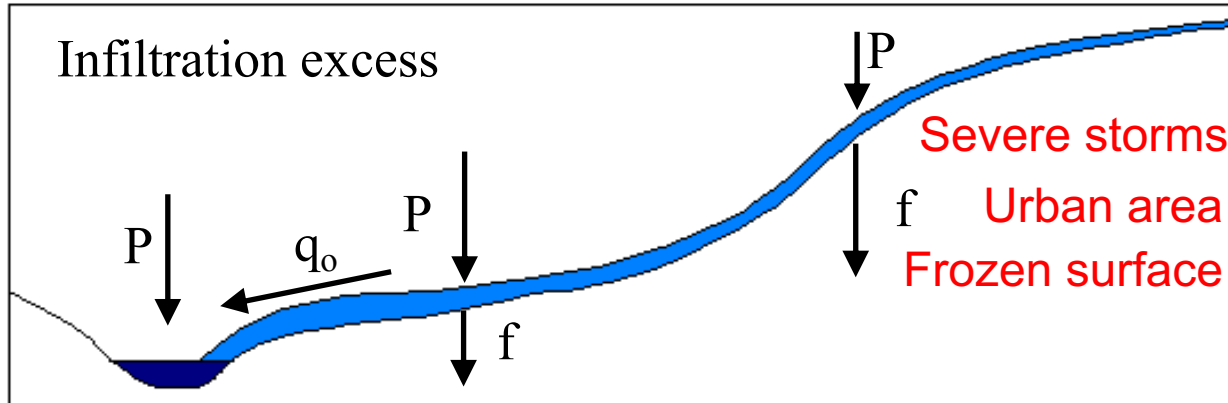
- Infiltration excess
- Saturation excess
- Baseflow (drainage)



Subgrid-scale soil moisture heterogeneity



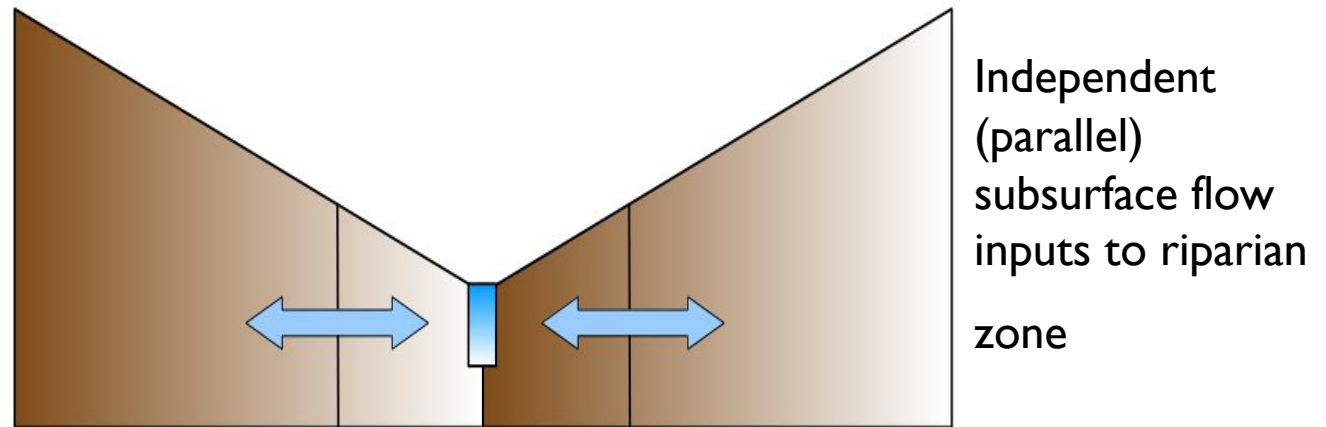
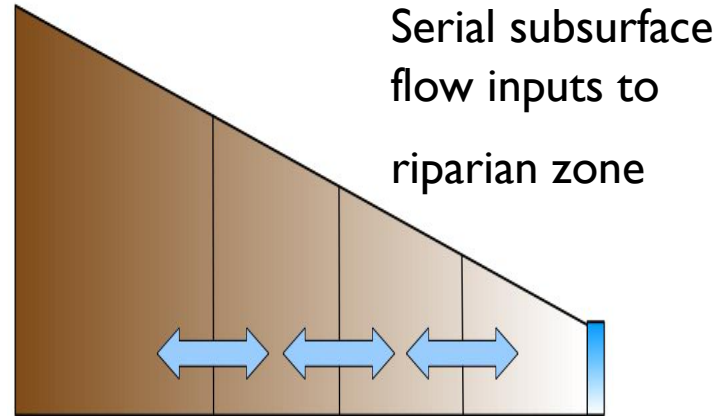
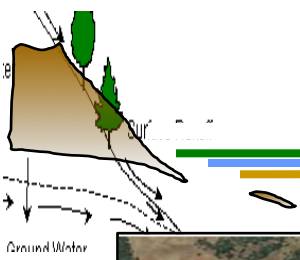
SIMTOP:TOPMODEL-based runoff



**Natural vegetation patterns imply controls from
soil moisture convergence, slope, and aspect**



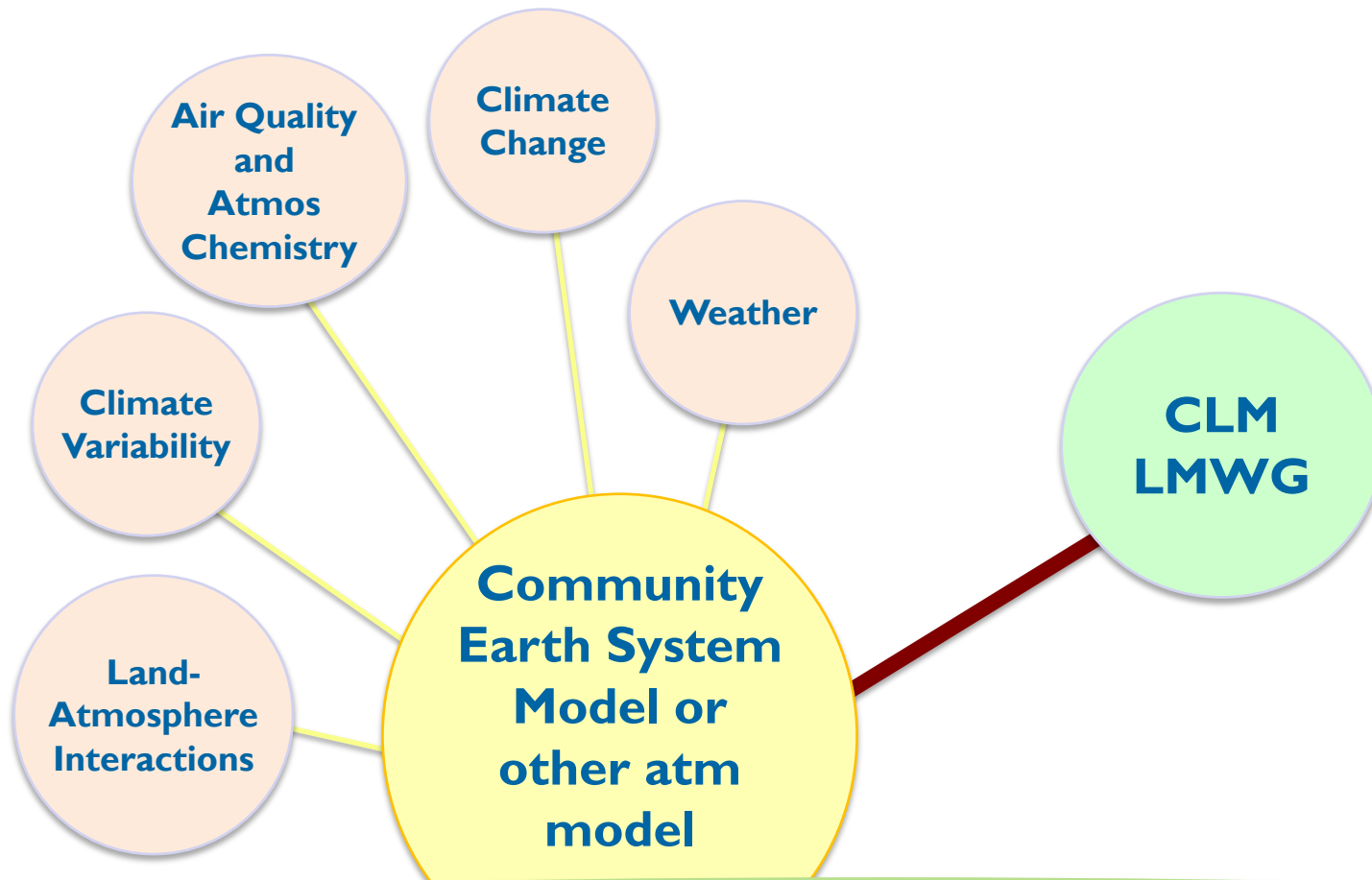
Representative hillslopes (CLM5 option)



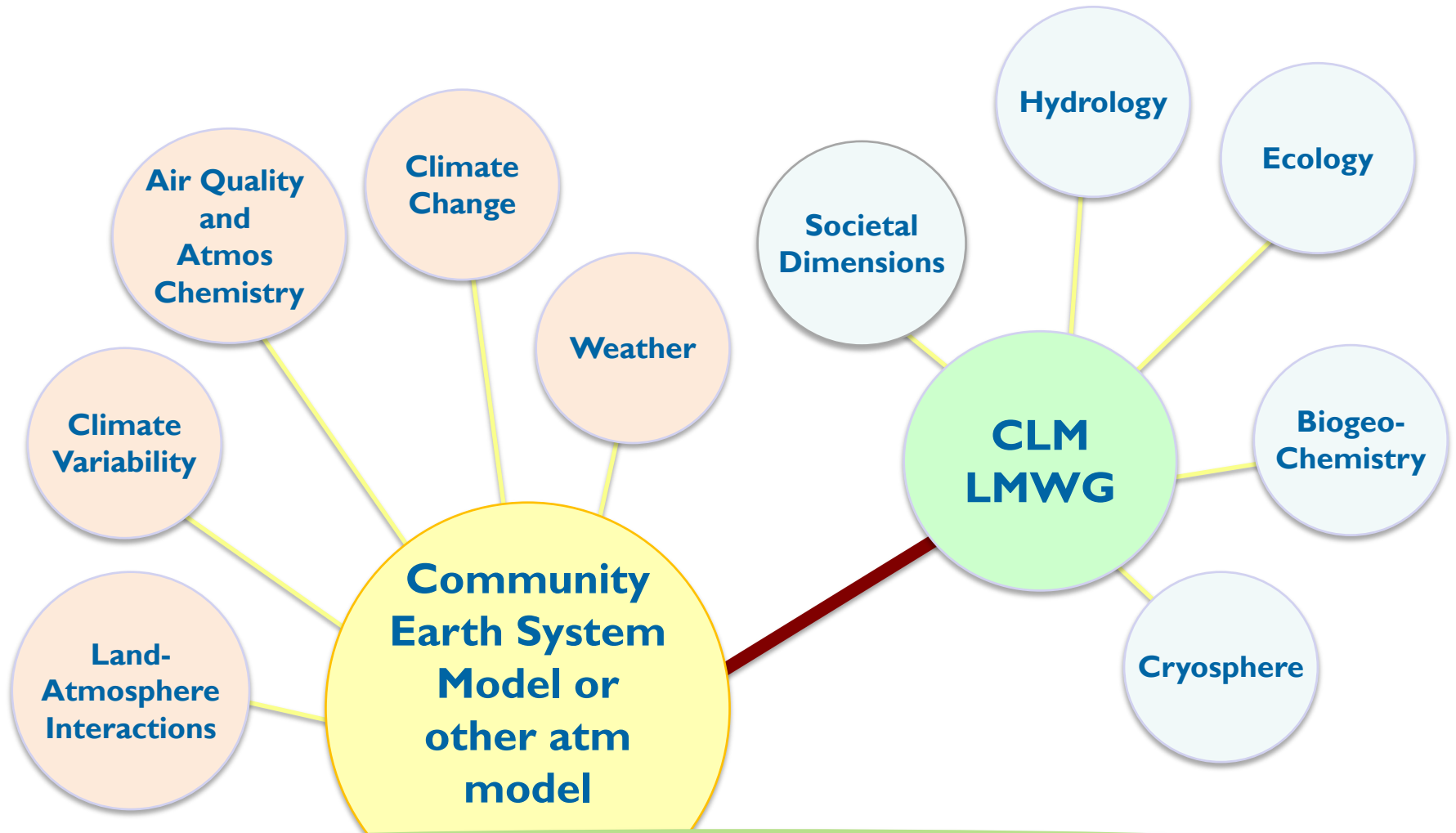
Model development and assessment



CLM as a community modeling tool



CLM as a community modeling tool



CLM4 (June 2010)



CLM4.5 (June 2013)



- Carbon and nitrogen prognostic vegetation
- Transient land cover and wood harvest
- 'Permafrost-enabled' deep ground
- Aerosol deposition
- Simple groundwater
- Urban model
- Vertically-resolved soil C/N
- Co-limitation and acclimation of photosynthesis
- Variable river flow rates
- Natural CH₄ emissions
- Human triggering and suppression of fire
- Cold region hydrology
- Revised lake model
- Multiple urban density classes

CLM4 (June 2010)



CLM4.5 (June 2013)



CLM5 (Feb 2018)



- Carbon and nitrogen prognostic vegetation
- Transient land cover wood harvest
- 'Permafrost-enabled' deep ground
- Aerosol deposition
- Simple groundwater
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- Vertically-
- Co-limitation of photosynthesis
- Variable riverine
- Natural CO₂
- Human trace gas suppression
- Cold region
- Revised land use
- Multiple urban classes

- Flexible leaf stoichiometry, leaf N optimize for photosynthesis
- Carbon costs for plant N uptake
- Plant hydraulics w/ hydraulic redistribution, *Ecosystem demography (FATES), ozone damage*
- Spatially explicit soil depth (0.4 – 8.5m), dry surface layer, revised GW, canopy interception, *representative hillslopes*
- MOSART river model (hillslope → tributary → main channel)
- Canopy snow, snow dens (T, wind), simple firn model
- Global crop model (8 crop types), transient irrigation and fertilization, *shifting cultivation*
- Dynamic landunits (nat veg ↔ crop, *glacier ↔ nat veg,*)
- Urban heating and AC, heat stress indices
- Carbon isotopes
- *Coupled fire trace gas emissions*

Land management in CLM5

Included in default CLM5

- Global crop model with 8 basic crop types; planting, grain fill, harvest
- Crop irrigation
- Crop industrial fertilization
- Wood harvest
- Urban environments
- Anthropogenic fire ignition and suppression

Corn*



Soy*

Winter wheat



Cotton

Sugarcane



Rice

* Temperate and tropical varieties

Fertilization



Irrigation



CLM4 (June 2010)

CLM4.5 (June 2013)

CLM5 (Feb 2018)



A central challenge: Model assessment

Are land models getting better or just more complex?

Do land models need to be more complex to be better?

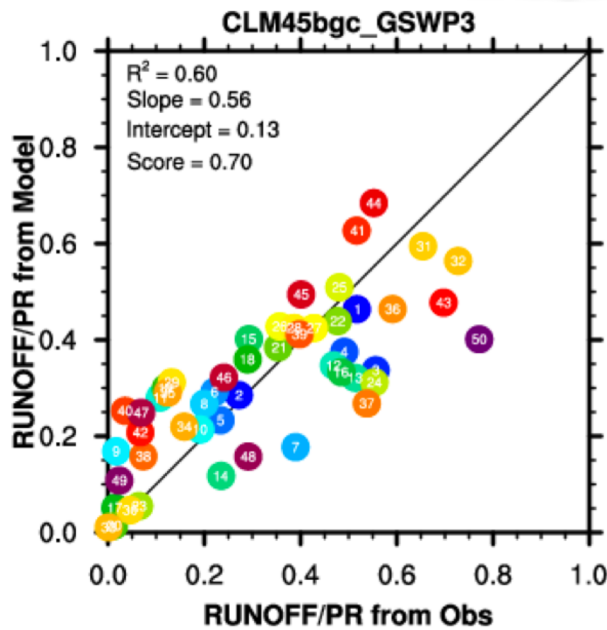
How do we interpret results from disparate set of models with varying degrees of comprehensiveness and complexity?

CMIP5 models, TRENDY models

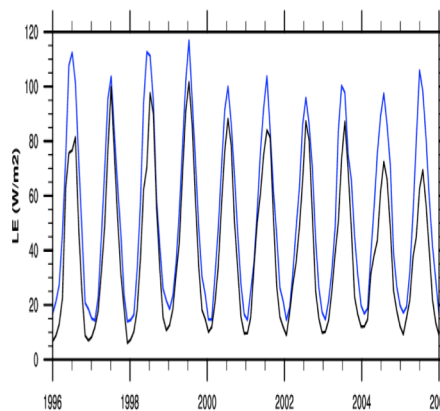


International Land Model Benchmarking (ILAMB) Package

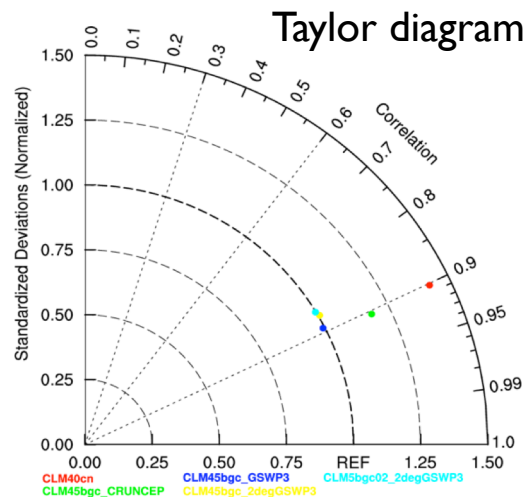
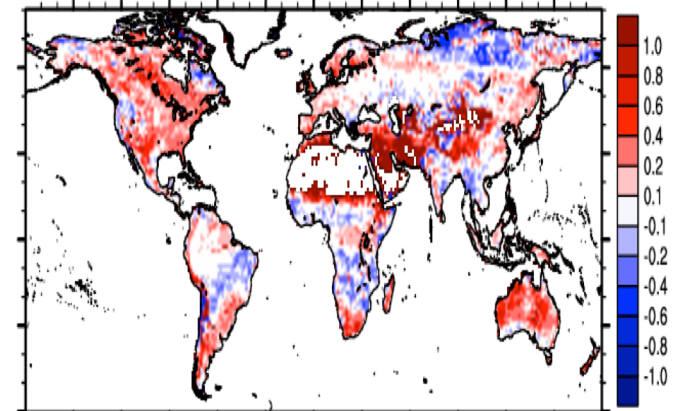
Land diagnostics package (25+ variables, 60+ datasets) with metrics for RMSE, bias, spatial pattern corr, interannual variability, functional relationships



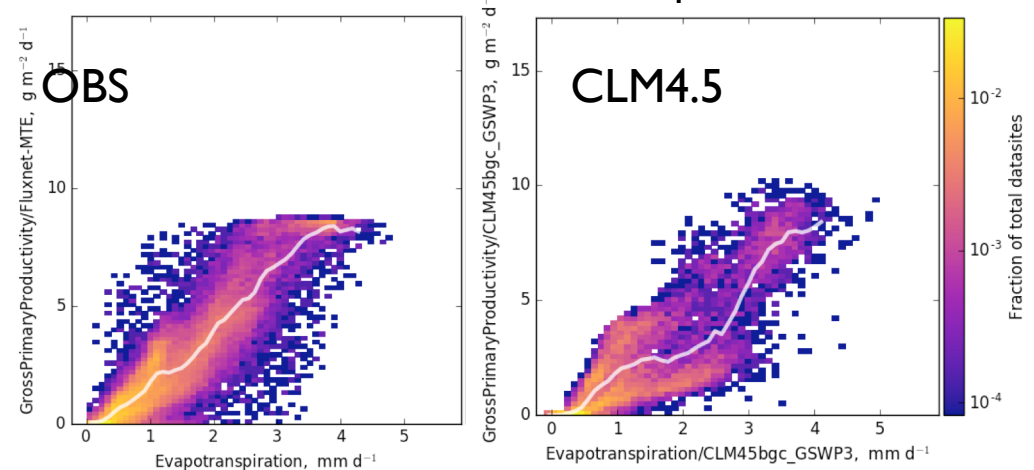
Tower Site



Global bias, relative bias, RMSE



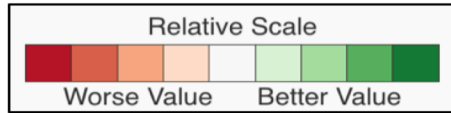
Functional relationships



CLM land-only forced with GSWP3

for full CLM results:

www.cesm.ucar.edu/experiments/cesm2.0/land/diagnostics/clm_diag_ILAMB.html



©

CLM4
CLM4.5
CLM5

Biomass



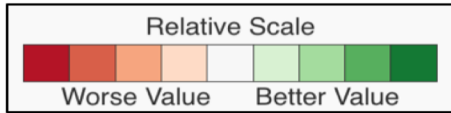
International Land Model Benchmarking (ILAMB) project

- Integrates analysis of ~30 variables against 60+ global, regional, and site-level observational datasets
- Graphics and scoring system for
 - RMSE
 - bias
 - seasonal cycle phase
 - spatial patterns
 - interannual variability
 - variable-to-variable relationships

CLM land-only forced with GSWP3

for full CLM results:

www.cesm.ucar.edu/experiments/cesm2.0/land/diagnostics/clm_diag_ILAMB.html



	CLM4	CLM4.5	CLM5
Ecosystem and Carbon Cycle			
Biomass	Orange	Light Green	Green
Burned Area	Orange	Light Green	Green
Carbon Dioxide	Orange	Light Green	Light Green
Gross Primary Productivity	Orange	Light Green	Light Green
Leaf Area Index	Orange	Light Green	Green
Global Net Ecosystem Carbon Balance	Light Green	Light Green	Green
Net Ecosystem Exchange	Orange	Light Green	Light Green
Ecosystem Respiration	Orange	Light Green	Light Green
Soil Carbon	Orange	Light Green	Light Green
Hydrology Cycle			
Evapotranspiration	Orange	Light Green	Light Green
Evaporative Fraction	Orange	Light Green	Light Green
Latent Heat	Orange	Light Green	Light Green
Runoff	Light Green	Light Green	Orange
Sensible Heat	Light Green	Light Green	Light Green
Terrestrial Water Storage Anomaly	Orange	Light Green	Light Green
Permafrost	Light Green	Orange	Light Green
Radiation and Energy Cycle			
Albedo	Light Green	Light Green	Light Green
Surface Upward SW Radiation	Light Green	Light Green	Light Green
Surface Net SW Radiation	Light Green	Light Green	Light Green
Surface Upward LW Radiation	Light Green	Light Green	Orange
Surface Net LW Radiation	Light Green	Light Green	Orange
Surface Net Radiation	Light Green	Light Green	Orange

- For majority of variables, progression in simulation quality from CLM4 to CLM5
- Why?
 - Improvements in mechanistic treatment of processes (e.g., hydrology, plant N processes, land use)
 - But, at same time, many more moving parts, additional unconstrained parameters



RUBISCO



Ground Water

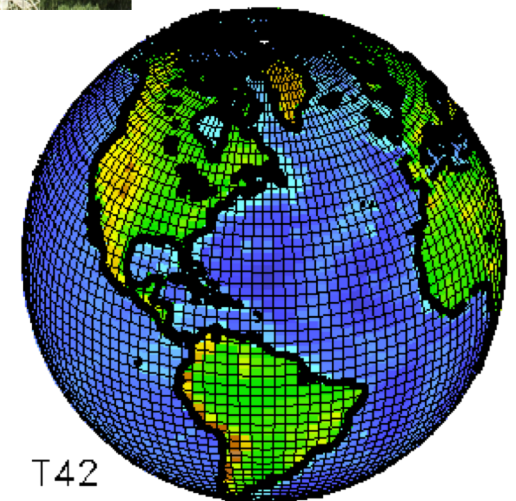
Model configurations

- SP (satellite phenology, prescribed vegetation)
- BGC (prognostic carbon, vegetation)
- BGC-crop (default in CESM2, same as BGC with crops)
- BGC no-anthro
- BGC FATES
- + many options for individual parameterizations (i.e., can revert to CLM4.5)



Spatial configurations

- Global (low and high resolution)
- Regional
- Single point (tower site)
- Irregular grids (cubed sphere, *basin*)





CLM as a research tool

Options to reduce complexity

- CH₄ emissions
- Carbon isotopes
- Land-use change
- VOC emissions
- Plant Hydraulics
- Soil structure (15-level vs 25-level)

Options to increase complexity

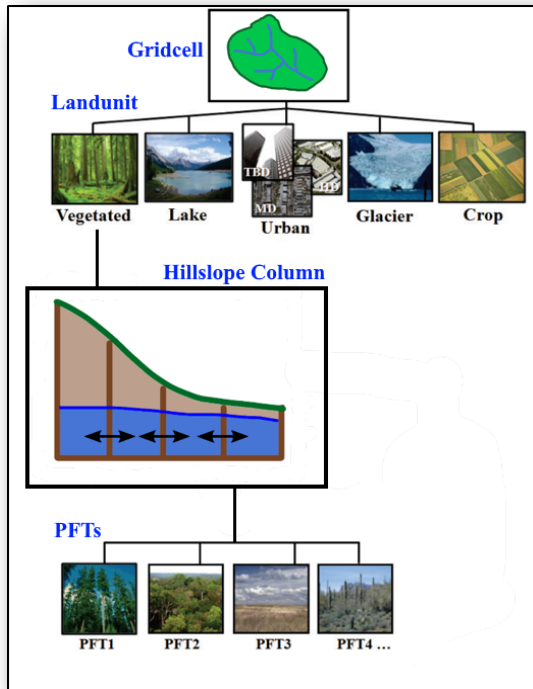
- Representative hillslopes
- FATES (Ecosystem dynamics)
- Fire trace gas emissions
- Additional land management
- Flooding
- Ozone damage to plants
- *Water tracers (available soon)*

Resources:

- CLM5 release webpage: www.cesm.ucar.edu/models/cesm2/land/
- CLM code repository: github.com/ESCOMP/ctsm
- CLM tutorial (Feb 2019): www.cgd.ucar.edu/events/2019/ctsm/
- Lawrence et al. (2019), in review *JAMES*

Some priorities and plans for next generation CLM

- Water and food security in context of climate variability, change, and extreme weather
- Ecosystem vulnerability and impacts on carbon cycle and ecosystem services
- Sources of predictability from land processes
- Impacts of land use and land-use change on climate, carbon, water, and extremes

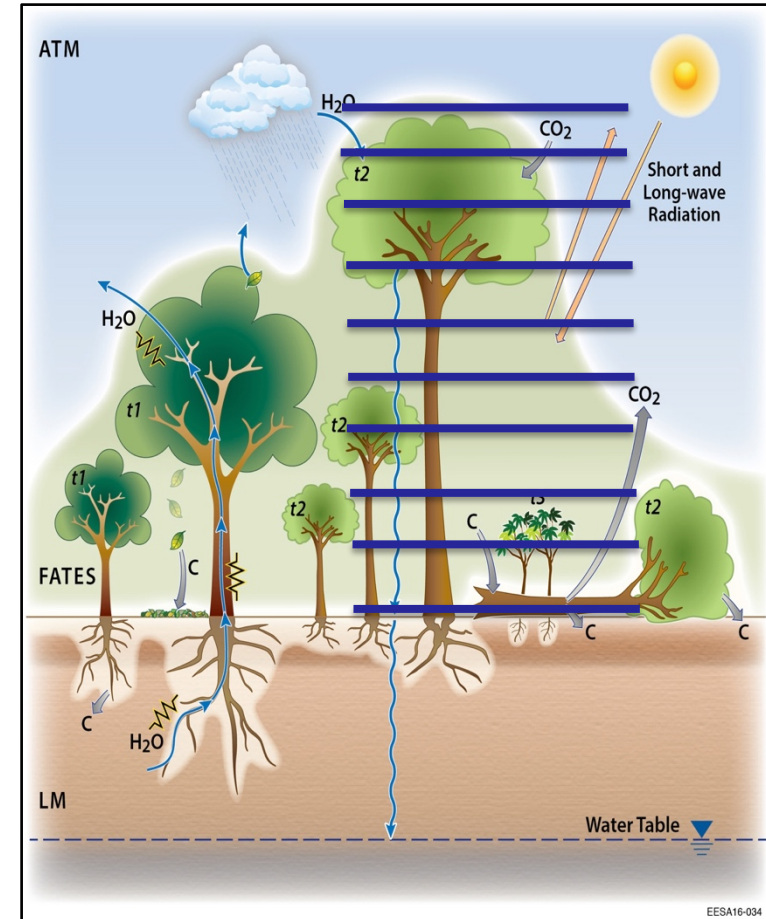


Lateral fluxes of water



Water and land management

Ecosystem Demography / Multi-layer canopy



Questions?

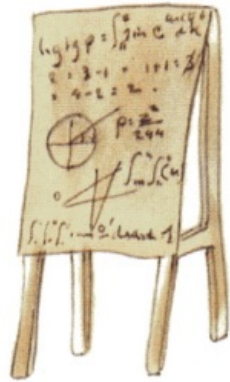


Extra slides

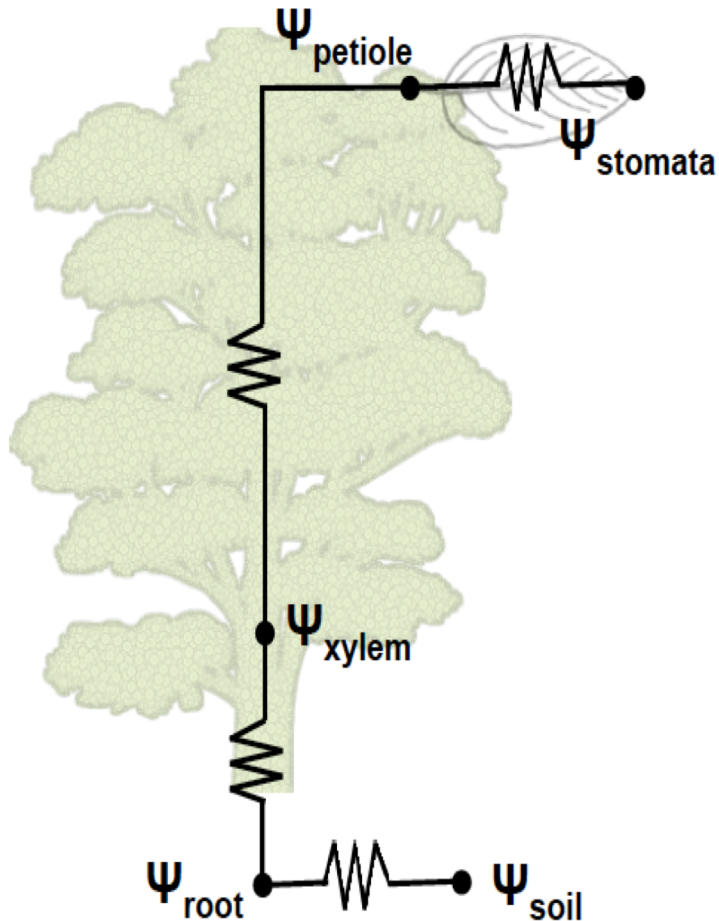


The diagram illustrates the hydrological cycle on a slope. It shows a cross-section of the ground with a brown surface and a tan interior. A green tree is on the left, and a green rectangular area labeled 'Silt' is on the right. Arrows indicate the flow of water: solid arrows for surface runoff, dashed arrows for infiltration into the ground, and a solid arrow for groundwater flow. The label 'Ground Water' is at the bottom left.

- Easy to run the model and get an answer
- Much harder to understand why you got that answer
- CLM is a very complex, multidisciplinary model



CLM5: Plant Hydrodynamics

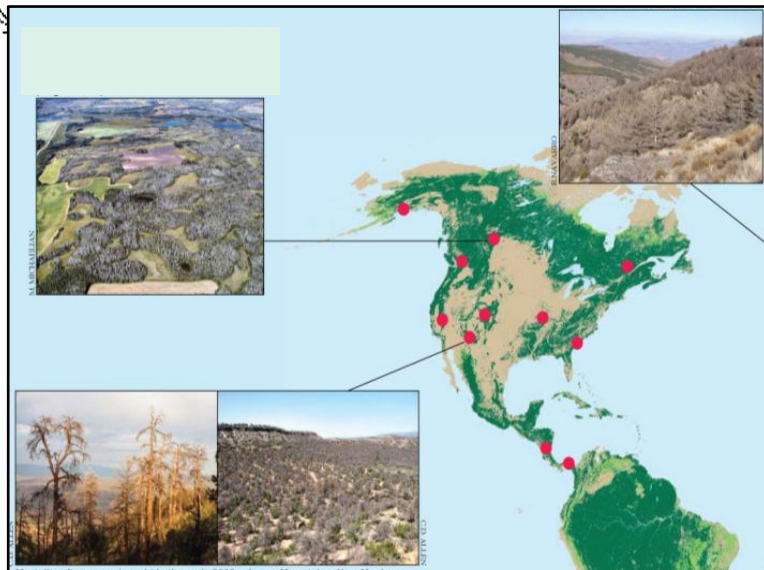
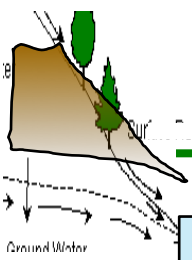


Why plant hydrodynamics

- BTRAN (soil moisture stress), and its parameters, θ_{crit} and θ_{wilt} , have no physical meaning and cannot be measured.
- Flux tower ET convolutes transpiration with canopy and soil evap making it difficult to use for process-level assessment. With plant hydrodynamics, sap flow measurements could be utilized.
- Satellites increasingly observe properties related to canopy or leaf water content.

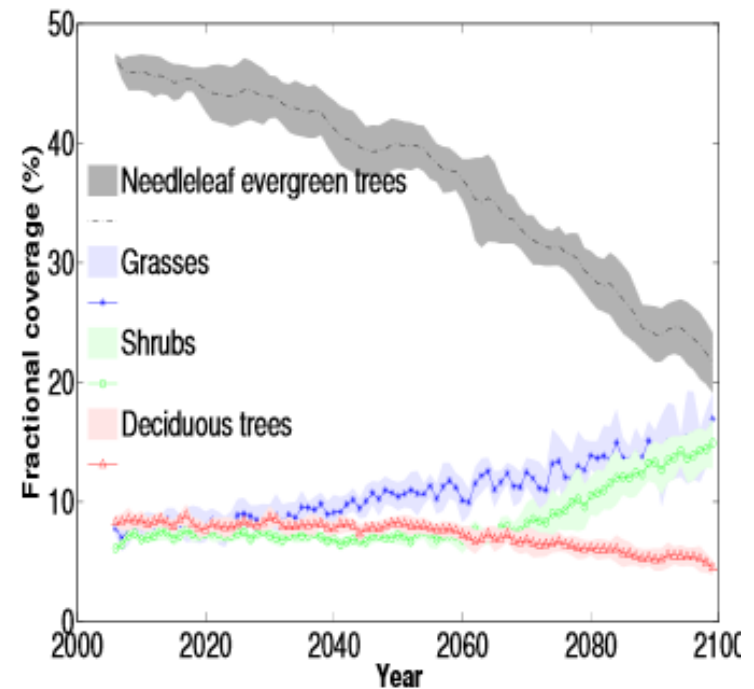
Ecosystem vulnerability to climate change

e.g., how vulnerable are western US forests to climate change?



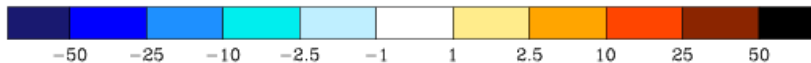
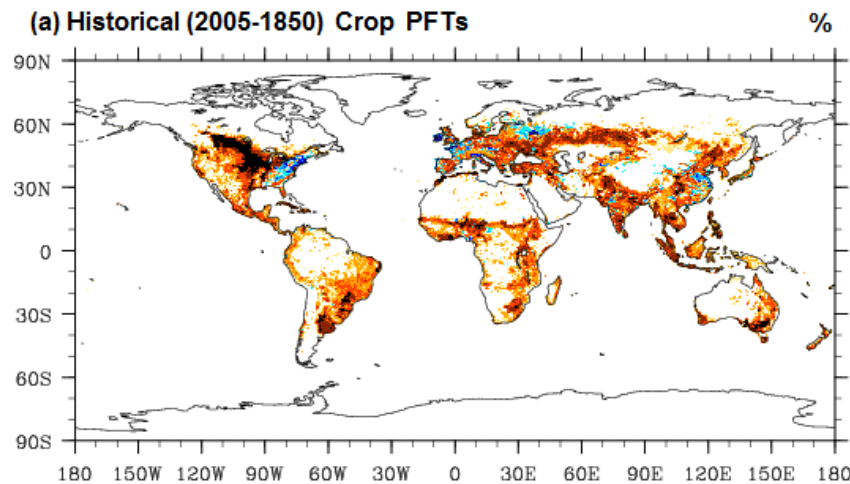
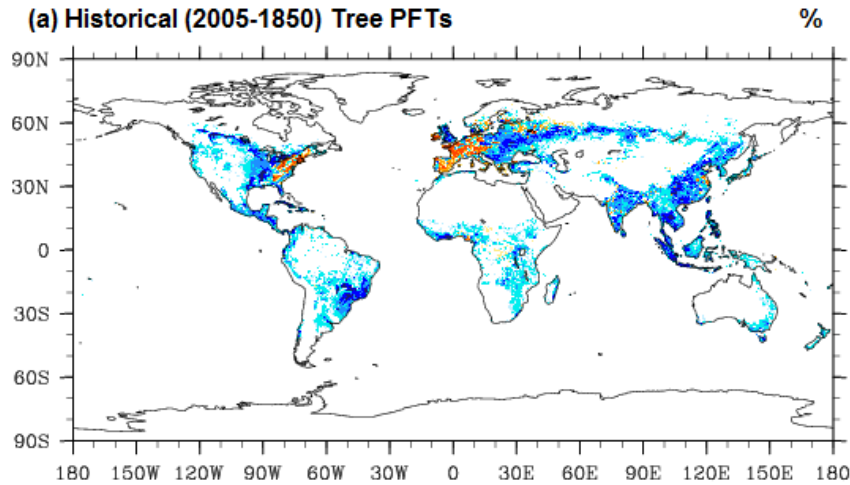
But ... these results are likely unreliable; tree response to soil moisture deficits represented in ad hoc way in land models. Forest loss is complex problem that requires combined consideration of climate, hydrology, ecology, and plant physiology and diversity

CLM4(DGVM), suggests widespread die-off of forests by 2100, but simple representation of hydrology, plant water use, mortality, ecosystem dynamics

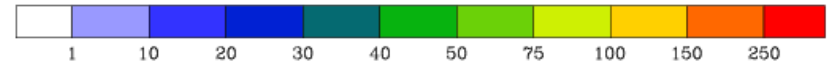
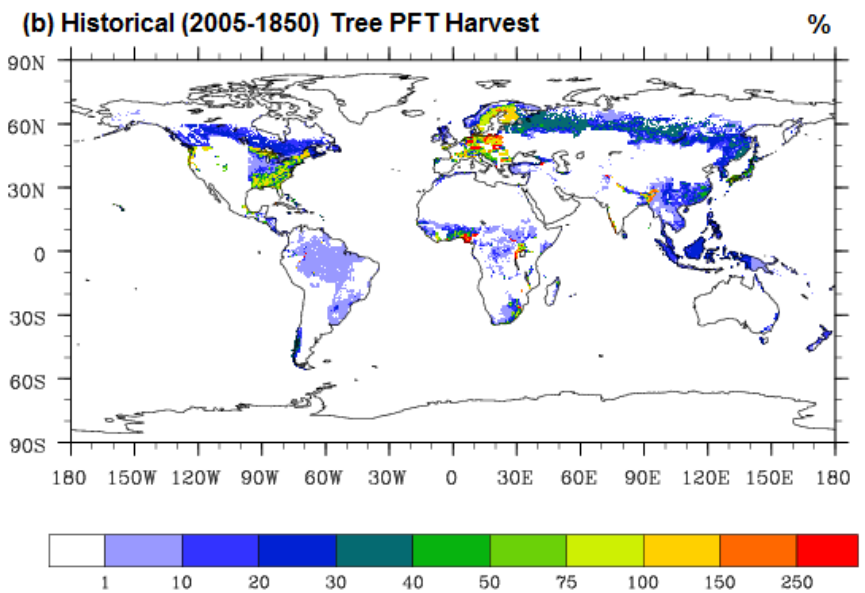


Historical land use & land cover change, 1850-2005

Change in tree and crop cover (% of grid cell)



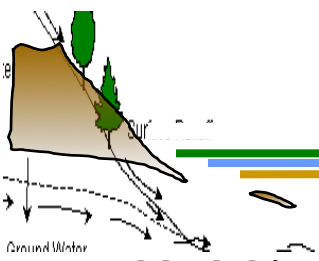
Cumulative percent of grid cell harvested



Historical LULCC

- ☐ Loss of tree cover and increase in cropland
- ☐ Farm abandonment and reforestation in eastern U.S. and Europe
- ☐ Extensive wood harvest

Many paths to improve models and reduce model uncertainty



Model intercomparisons (MIPs)

- CMIP6: carbon cycle, land use, land-atmosphere coupling, ...
- Range of plausible outcomes, but more models \neq better results

Model benchmarking

- Comprehensive model evaluation against observations

Real-world experiments and models

- FACE, N addition

Model-data fusion

- Data assimilation, parameter estimation

“Discover” critical missing process

- Add another process that is ecologically or hydrologically important but poorly known at the global scale. Tune a key parameter to get a good simulation.

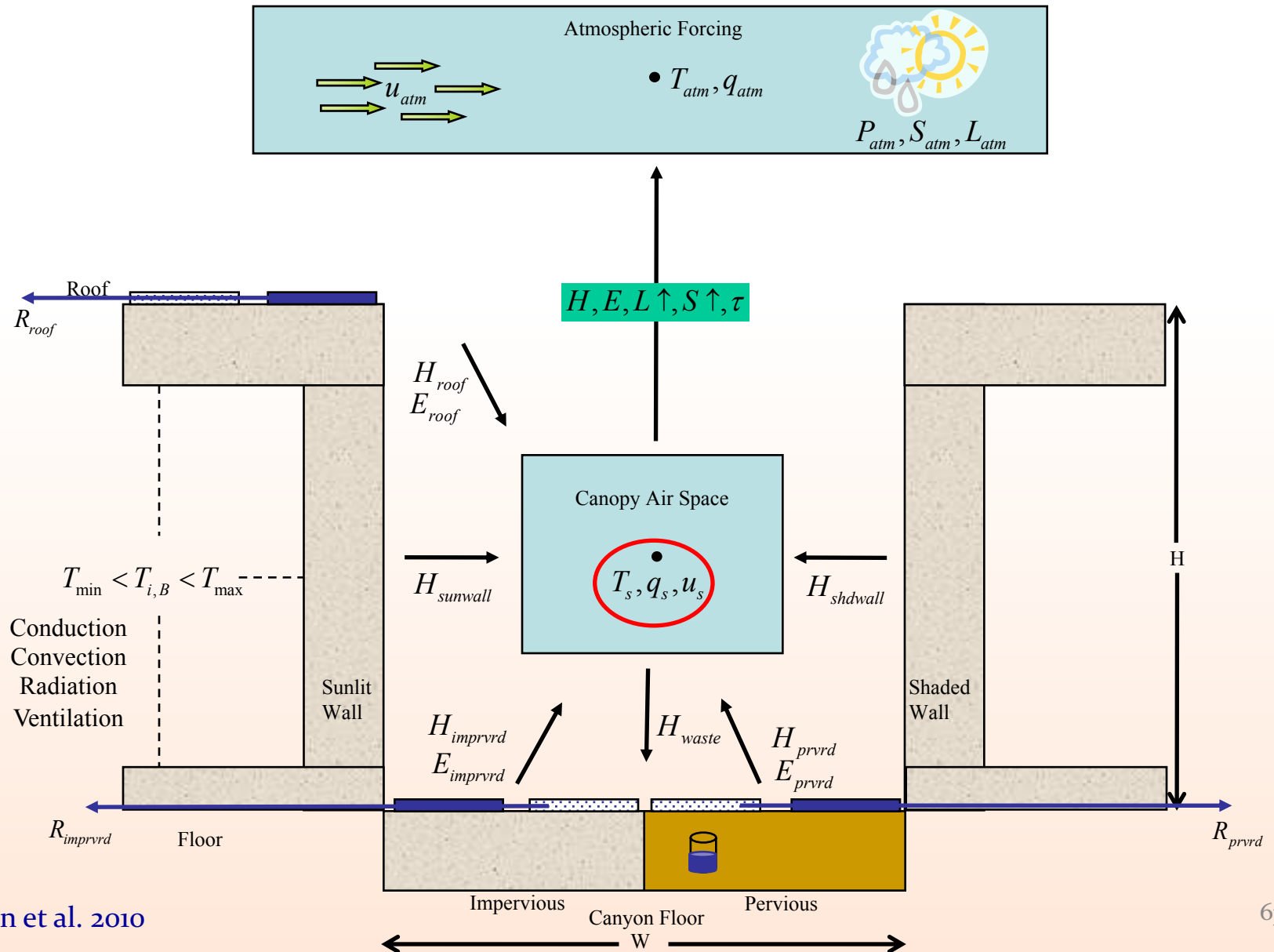
Model intracomparison

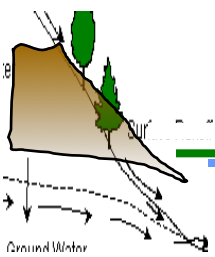
- Focus on model structural uncertainty to identify processes contributing to uncertainty

Model hierarchy

- CLM
- Process models (multilayer canopy, MIMICS)
- Simple land models (Marysa Lague)

Urban Model

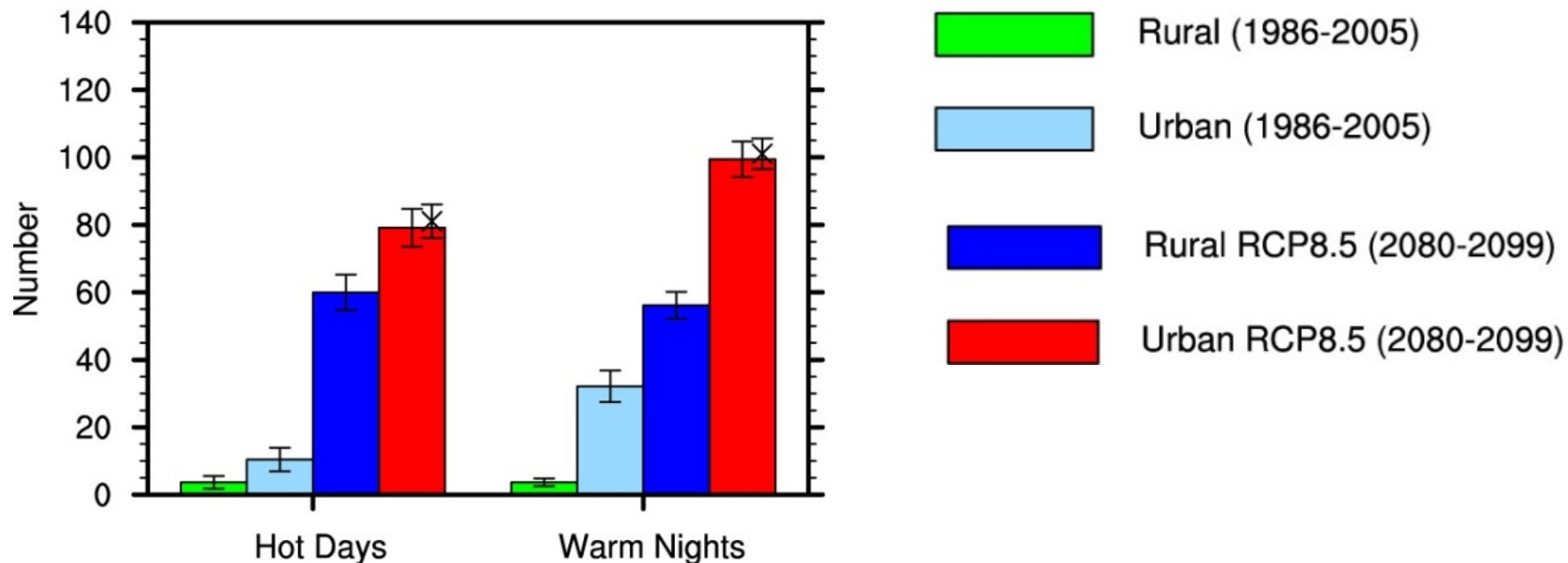




Changes in hot days and warm nights – RCP8.5

Hot days (warm nights) – Number of days per year that daily TMAX (TMIN) exceeds 99th percentile of present day Rural daily TMAX (TMIN)

New York

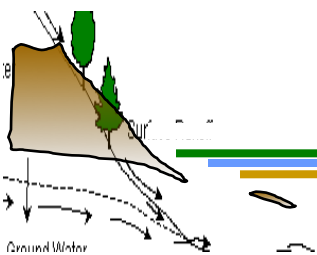


Present-day climate

Cities have more hot days and warm nights than rural land

21st century climate change

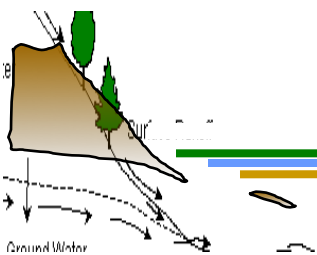
Cities increase more in hot days and warm nights than does rural land



The role of CLM in CESM:

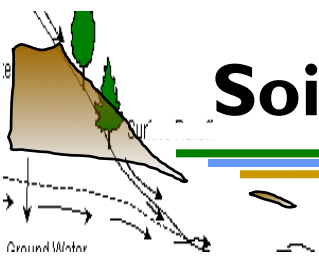
Land to Atmosphere

¹ Latent heat flux	$\lambda_{vap} E_v + \lambda E_g$	W m^{-2}
Sensible heat flux	$H_v + H_g$	W m^{-2}
Water vapor flux	$E_v + E_g$	mm s^{-1}
Zonal momentum flux	τ_x	$\text{kg m}^{-1} \text{s}^{-2}$
Meridional momentum flux	τ_y	$\text{kg m}^{-1} \text{s}^{-2}$
Emitted longwave radiation	$L \uparrow$	W m^{-2}
Direct beam visible albedo	$I \uparrow_{vis}^{\mu}$	-
Direct beam near-infrared albedo	$I \uparrow_{nir}^{\mu}$	-
Diffuse visible albedo	$I \uparrow_{vis}$	-
Diffuse near-infrared albedo	$I \uparrow_{nir}$	-
Absorbed solar radiation	\vec{S}	W m^{-2}
Radiative temperature	T_{rad}	K
Temperature at 2 meter height	T_{2m}	K
Specific humidity at 2 meter height	q_{2m}	kg kg^{-1}
Snow water equivalent	W_{sno}	m
Aerodynamic resistance	r_{am}	s m^{-1}
Friction velocity	u_*	m s^{-1}
² Dust flux	F_j	$\text{kg m}^{-2} \text{s}^{-1}$
Net ecosystem exchange	NEE	$\text{kgCO}_2 \text{m}^{-2} \text{s}^{-1}$



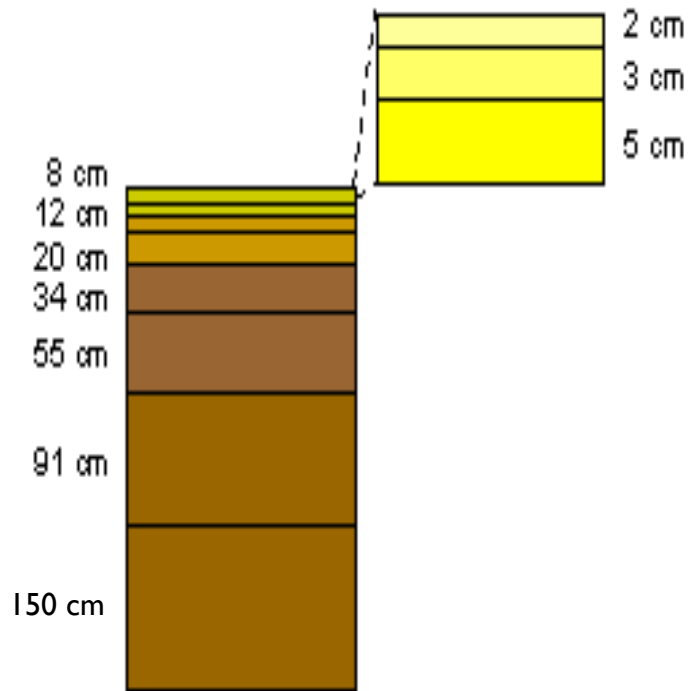
The role of CLM in CESM: Atmosphere to Land

¹ Reference height	z'_{atm}	m
Zonal wind at z_{atm}	u_{atm}	m s^{-1}
Meridional wind at z_{atm}	v_{atm}	m s^{-1}
Potential temperature	$\overline{\theta_{atm}}$	K
Specific humidity at z_{atm}	q_{atm}	kg kg^{-1}
Pressure at z_{atm}	P_{atm}	Pa
Temperature at z_{atm}	T_{atm}	K
Incident longwave radiation	$L_{atm} \downarrow$	W m^{-2}
² Liquid precipitation	q_{rain}	mm s^{-1}
² Solid precipitation	q_{sno}	mm s^{-1}
Incident direct beam visible solar radiation	$S_{atm} \downarrow^{\mu}_{vis}$	W m^{-2}
Incident direct beam near-infrared solar radiation	$S_{atm} \downarrow^{\mu}_{nir}$	W m^{-2}
Incident diffuse visible solar radiation	$S_{atm} \downarrow_{vis}$	W m^{-2}
Incident diffuse near-infrared solar radiation	$S_{atm} \downarrow_{nir}$	W m^{-2}
Carbon dioxide (CO ₂) concentration	c_a	ppmv
³ Aerosol deposition rate	D_{sp}	$\text{kg m}^{-2} \text{s}^{-1}$
⁴ Nitrogen deposition rate	NF_{ndep_sminn}	$\text{g (N) m}^{-2} \text{yr}^{-1}$
⁵ Lightning frequency	I_l	$\text{flash km}^2 \text{hr}^{-1}$



Soil Texture – thermal/hydrologic parameters

Soil parameters are derived from sand / clay percentage and soil organic matter content which is specified geographically and by soil level

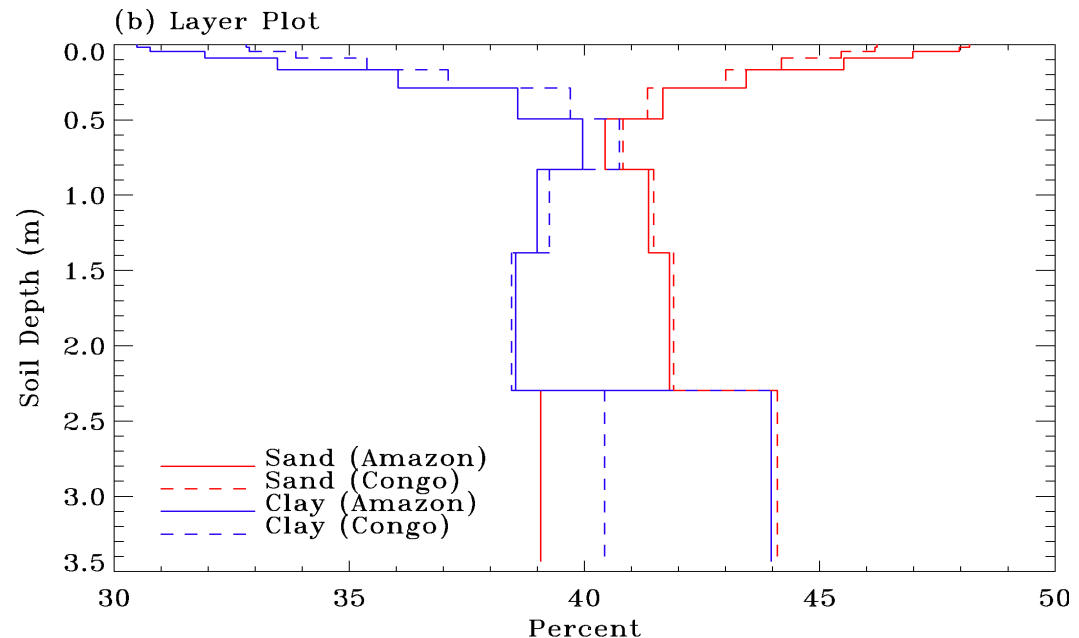


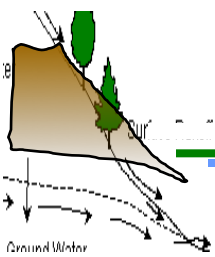
Soil profile

10 soil levels (~3.8m)

5 bedrock levels (~42m)

- Soil moisture concentration at saturation
- Soil moisture concentration at wilting point
- Hydraulic conductivity at saturation
- Saturated soil suction
- Thermal conductivity
- Thermal capacity



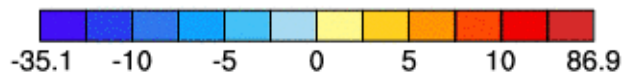
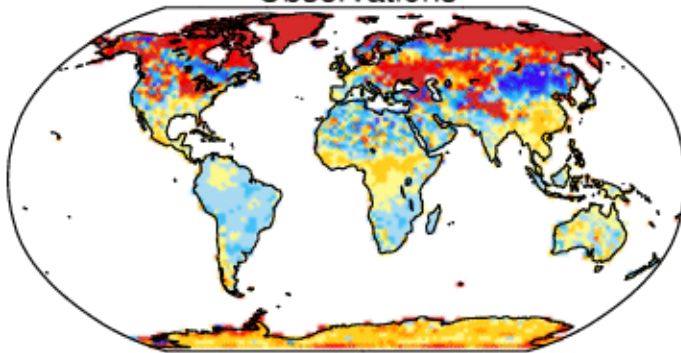


Modeling surface albedo

DJF ASA (% reflected)

CLM45SP_CRUNCEP

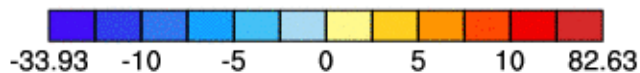
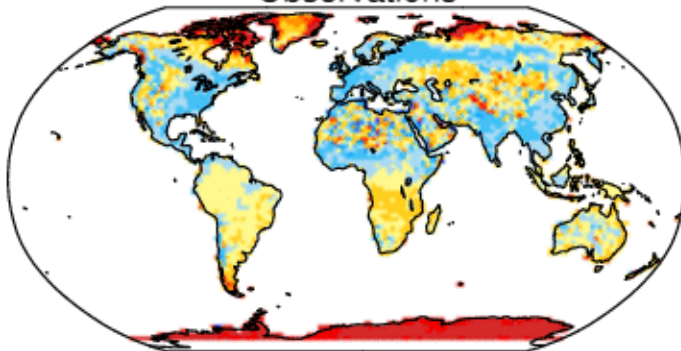
- Observations



JJA ASA (% reflected)

CLM45SP_CRUNCEP

- Observations



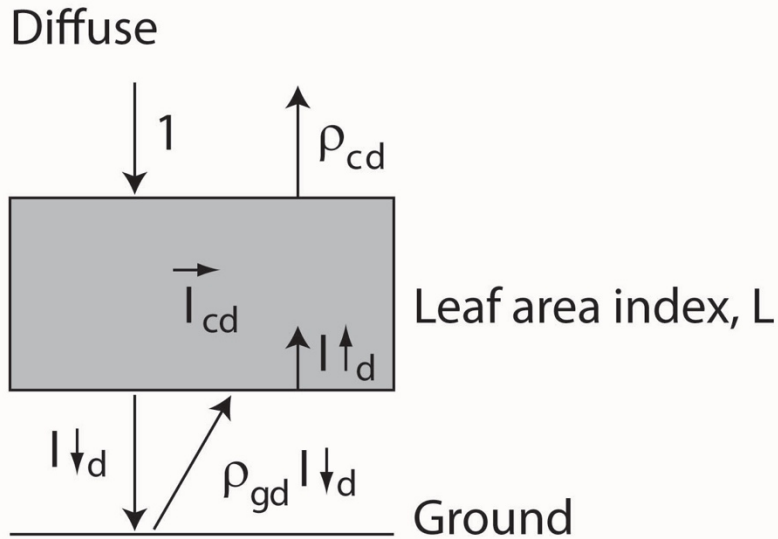
Surface albedo a function of

- Vegetation cover and type
- Snow cover
- Snow age
- Soil moisture
- Soil color
- Solar zenith angle
- Amount of direct vs diffuse solar radiation
- Amount of visible vs IR solar radiation

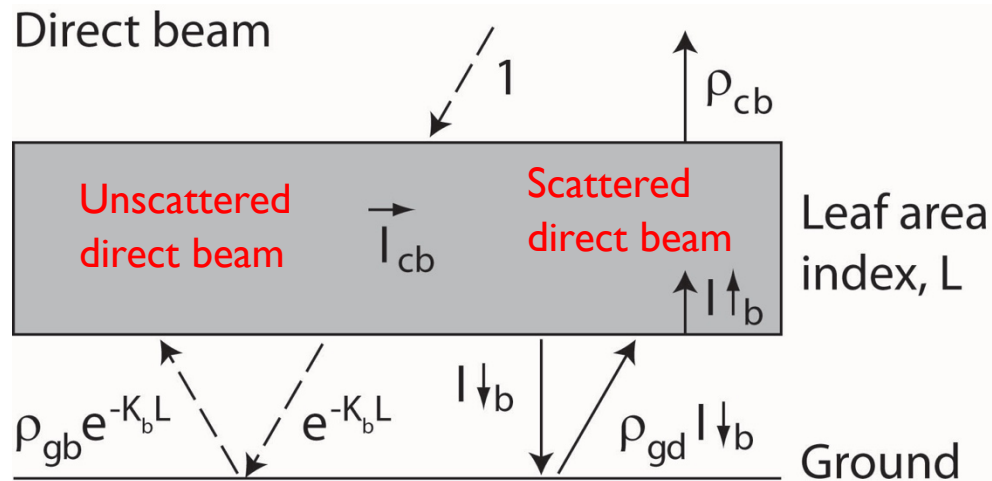
Note: MODIS albedo biased low for snow at high zenith angle

(Wang and Zender, 2010)

Two-stream radiative transfer



Radiative transfer uses the two-stream approximation (Dickinson, Sellers) to determine reflected and absorbed solar radiation



Momentum, and sensible heat and evaporation fluxes

Momentum flux

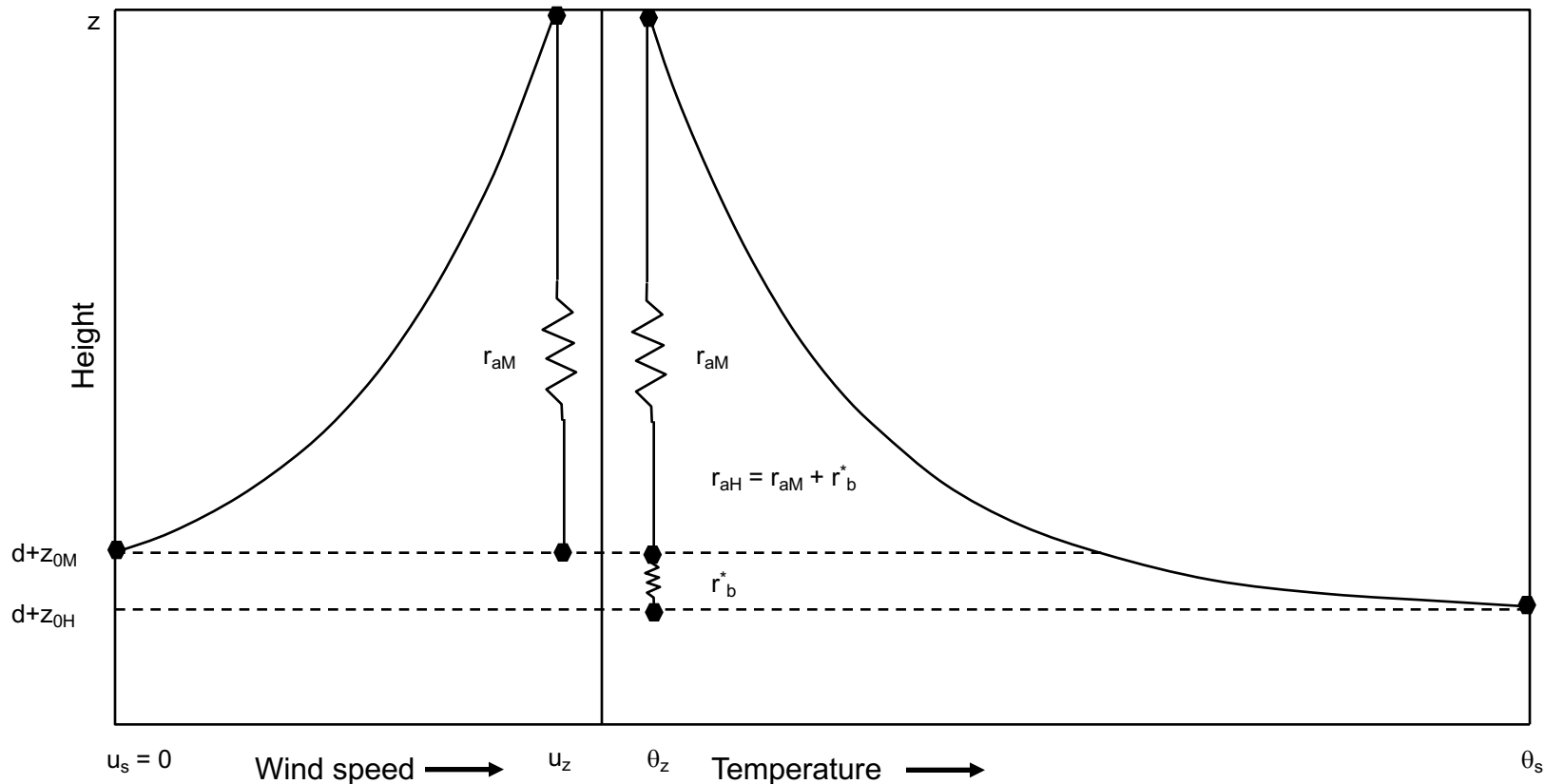
$$u_* u_* = \tau / \rho \quad \text{and} \quad \tau = \rho(u_a - u_s) / r_{aM} = \rho u / r_{aM} \quad \Rightarrow \quad r_{aM} = \frac{1}{k^2 u} \left[\ln \left(\frac{z-d}{z_{0M}} \right) - \psi_m(\zeta) \right]^2$$

Sensible heat flux

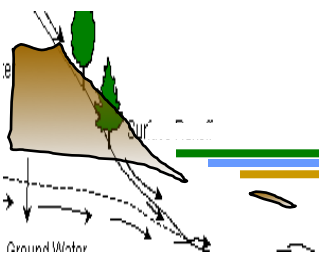
$$\theta_* u_* = -H / (\rho c_p) \quad \text{and} \quad H = -\rho c_p (\theta_a - T_s) / r_{aH} \quad \Rightarrow \quad r_{aH} = \frac{1}{k^2 u} \left[\ln \left(\frac{z-d}{z_{0M}} \right) - \psi_m(\zeta) \right] \left[\ln \left(\frac{z-d}{z_{0H}} \right) - \psi_h(\zeta) \right]$$

Evaporation

$$q_* u_* = -E / \rho \quad \text{and} \quad E = -\rho (q_a - q_s) / r_{aW} \quad \Rightarrow \quad r_{aW} = \frac{1}{k^2 u} \left[\ln \left(\frac{z-d}{z_{0M}} \right) - \psi_m(\zeta) \right] \left[\ln \left(\frac{z-d}{z_{0W}} \right) - \psi_w(\zeta) \right]$$



Snow/Soil thermodynamics

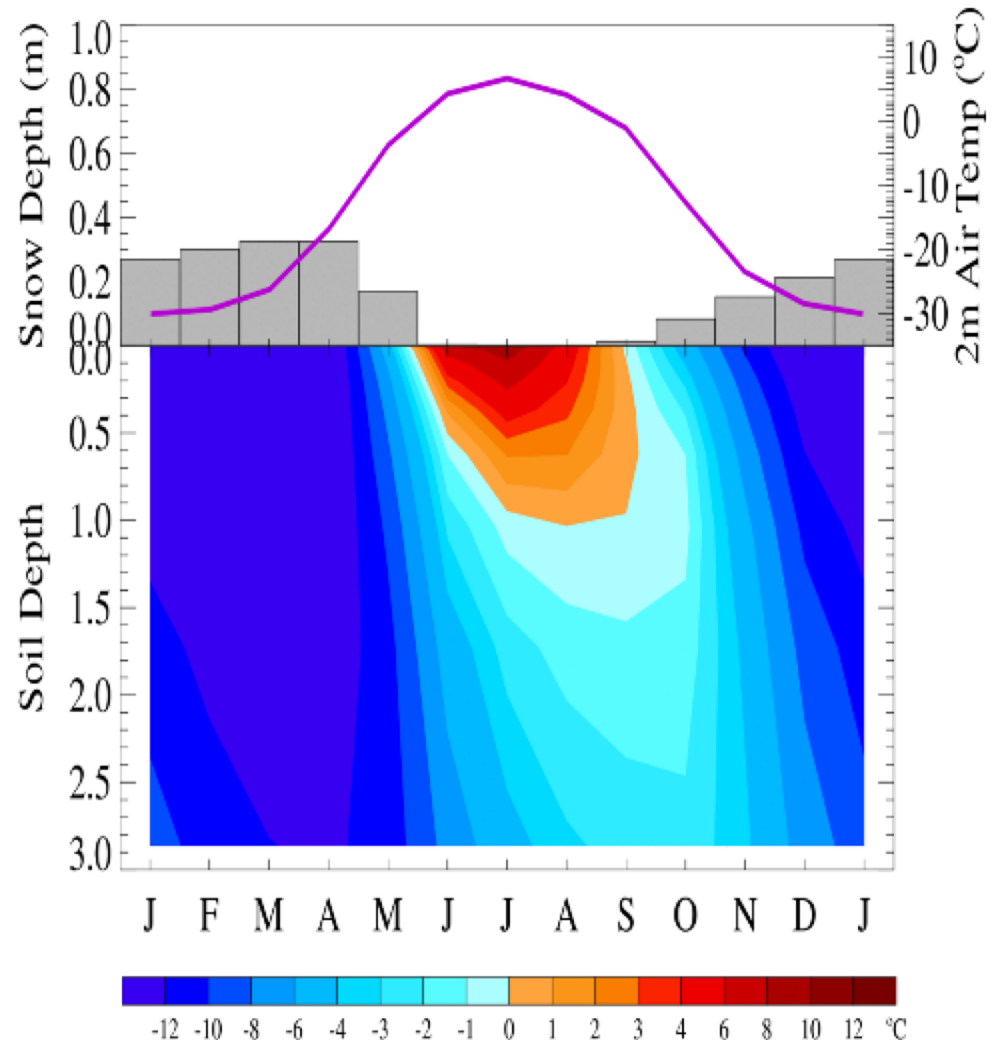


Solve the heat diffusion equation for multi-layer snow and soil model

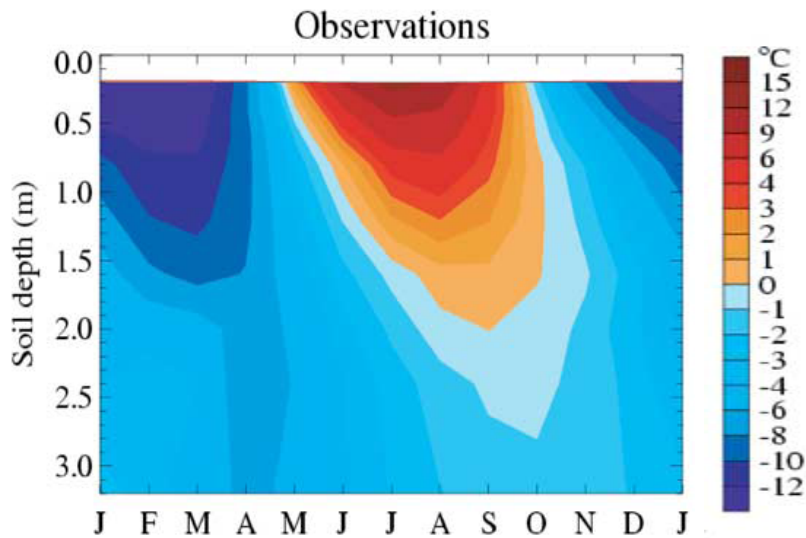
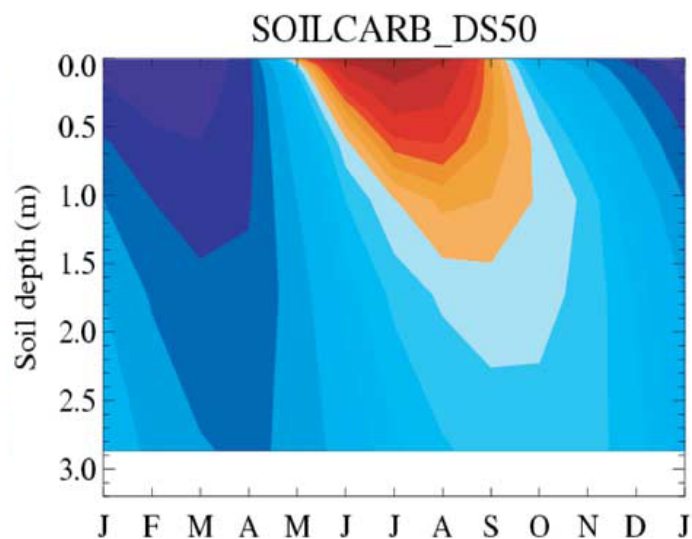
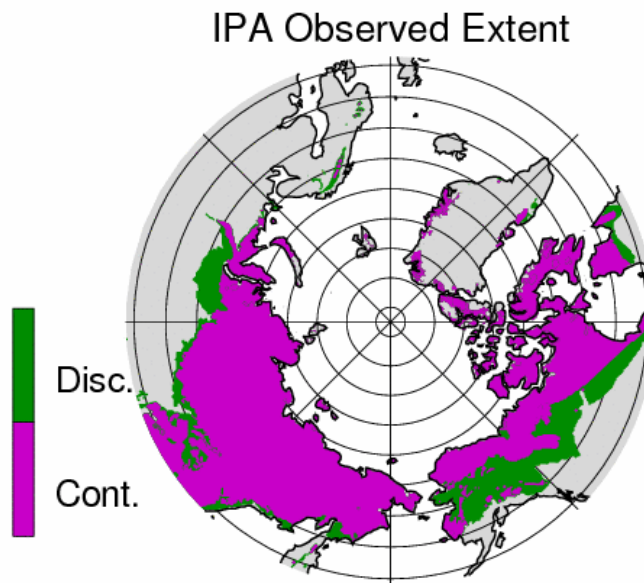
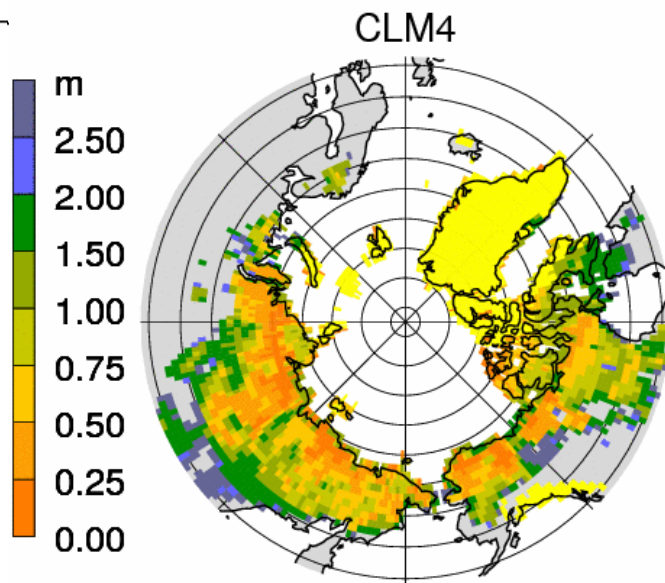
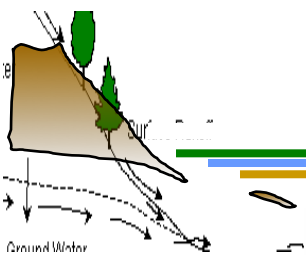
$$C_p \frac{\partial T}{\partial t} = \frac{\partial}{\partial z} \left(K \frac{\partial T}{\partial z} \right)$$

where C_p (heat capacity) and K (thermal conductivity) are functions of:

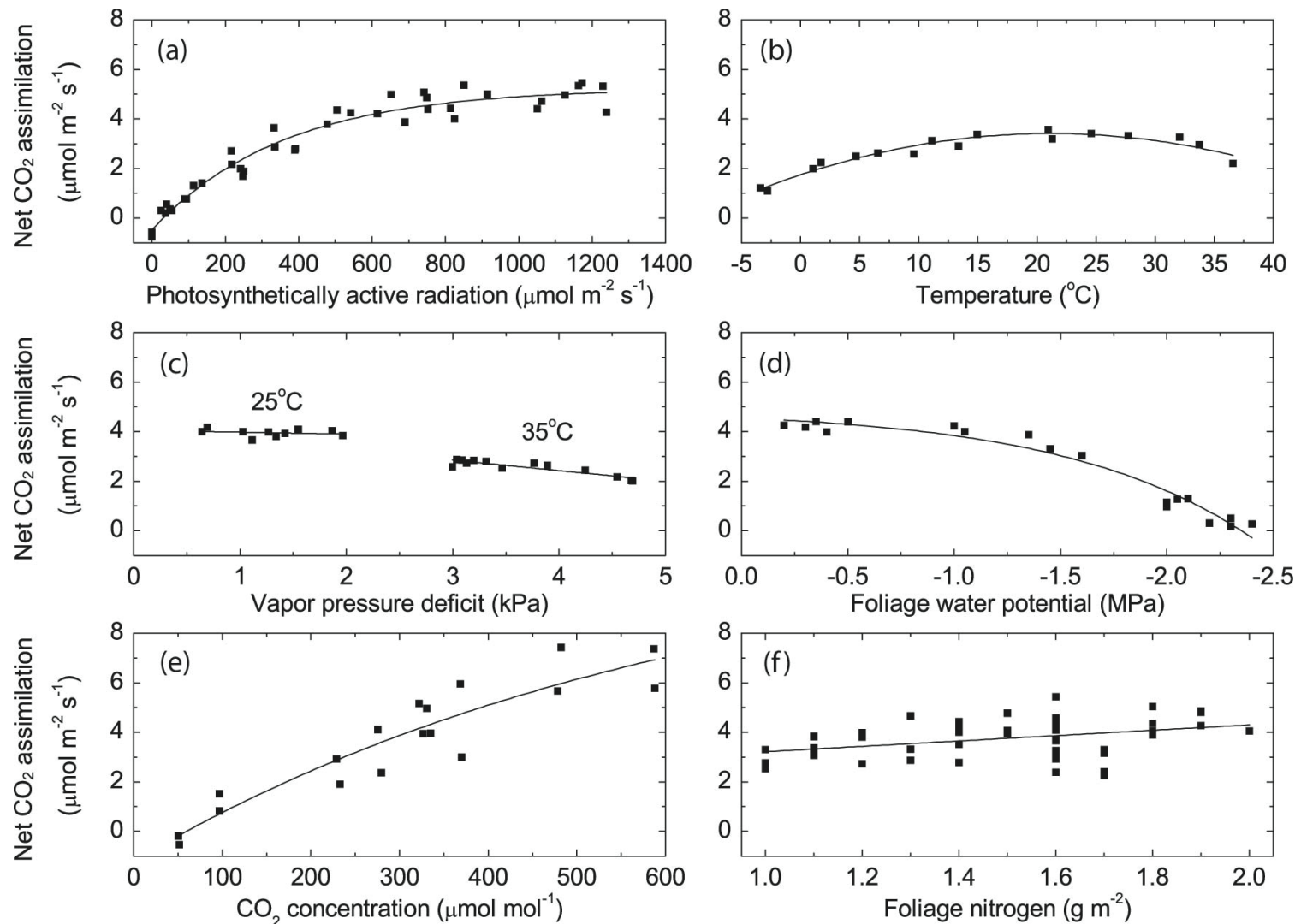
- temperature
- total soil moisture
- soil texture
- ice/liquid content



Modeling Permafrost in CLM



Leaf photosynthesis



Leaf photosynthesis and stomatal conductance

Farquhar photosynthesis model

$$A_n = \min(w_c, w_j, w_p) - R_d$$

w_c is the rubisco-limited rate of photosynthesis, w_j is light-limited rate allowed by RuBP regeneration, w_p is product limited rate of carboxylation

rubisco-limited rate is

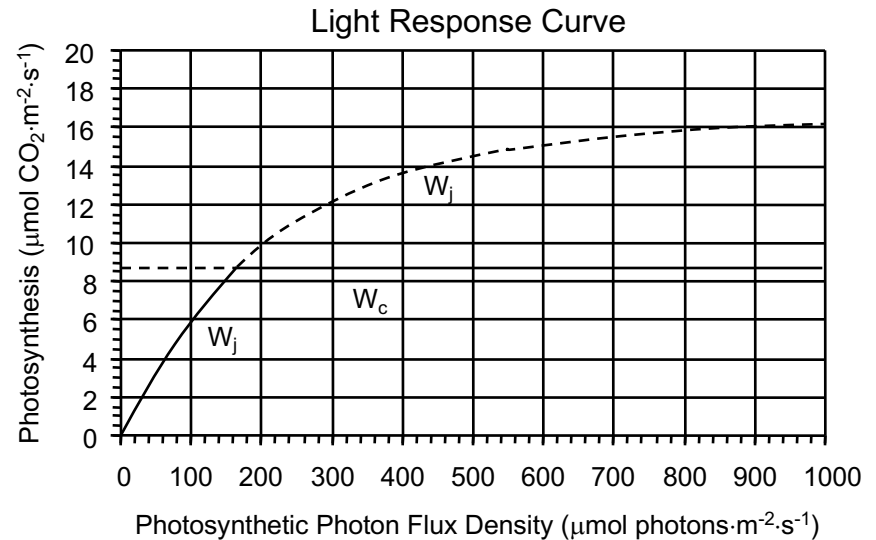
$$w_c = \frac{V_{c\max}(c_i - \Gamma^*)}{c_i + K_c(1 + O_i/K_o)}$$

RuBP regeneration-limited rate is

$$w_j = \frac{J(c_i - \Gamma^*)}{4(c_i + 2\Gamma^*)}$$

product-limited rate is

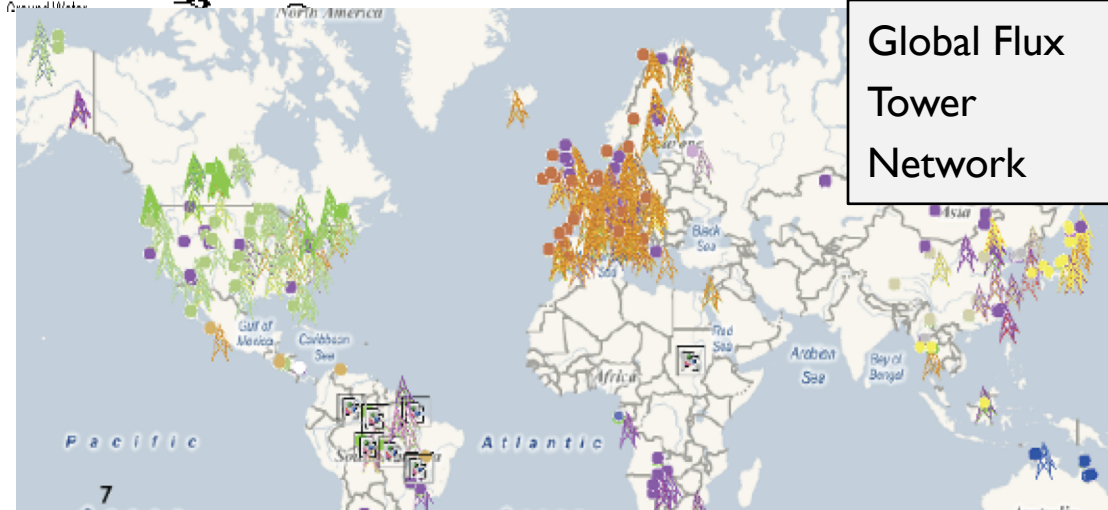
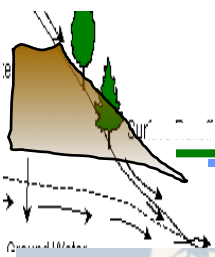
$$w_p = 3T_p$$



Ball-Berry stomatal conductance

$$\frac{1}{r_s} = g_s = g_1 \frac{A_n h_s}{c_s / P_{atm}} + g_0 \beta_t$$

Evaluating the model with tower flux data

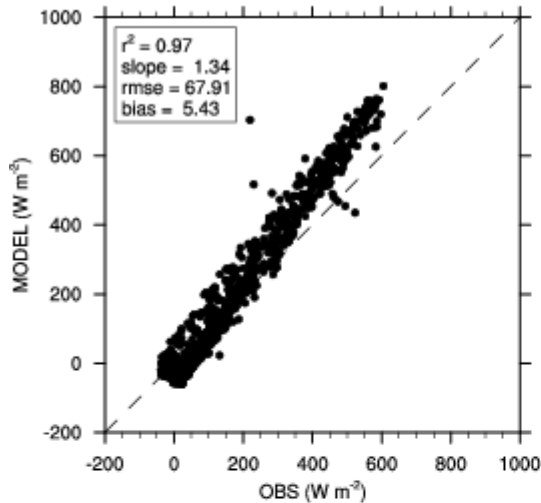


Evaluating CLM4.5 with tower flux data

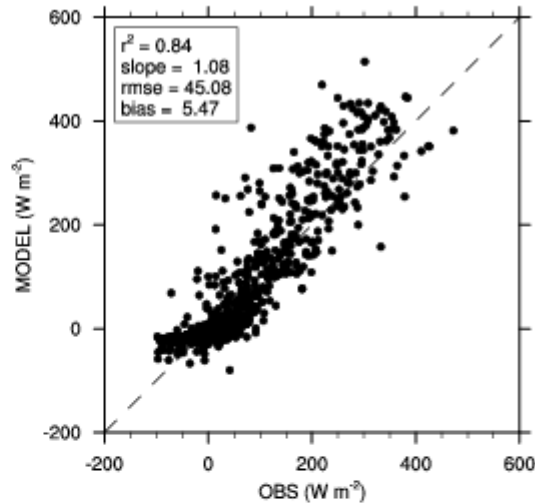
Howland Forest, Maine, July, 1996

AMF_USHo1 CLM451_r111_SP, Observed Fluxes, DOY_183-213_1996

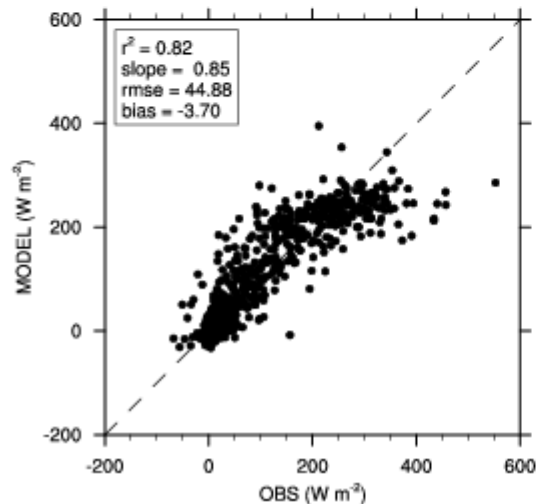
Net Radiation



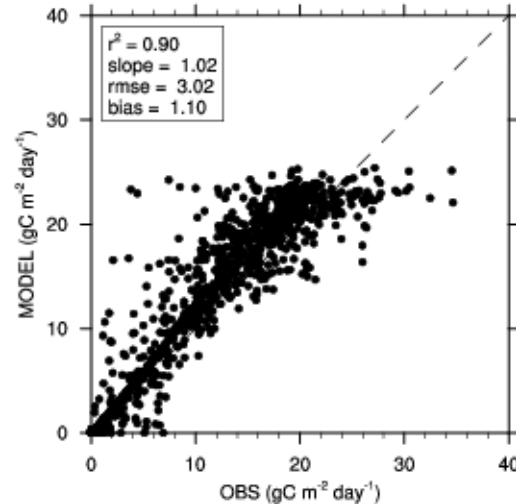
Sensible Heat Flux

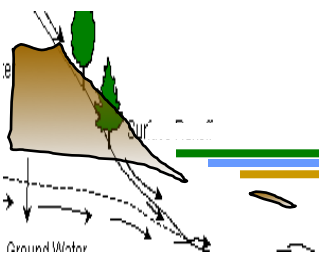


Latent Heat Flux



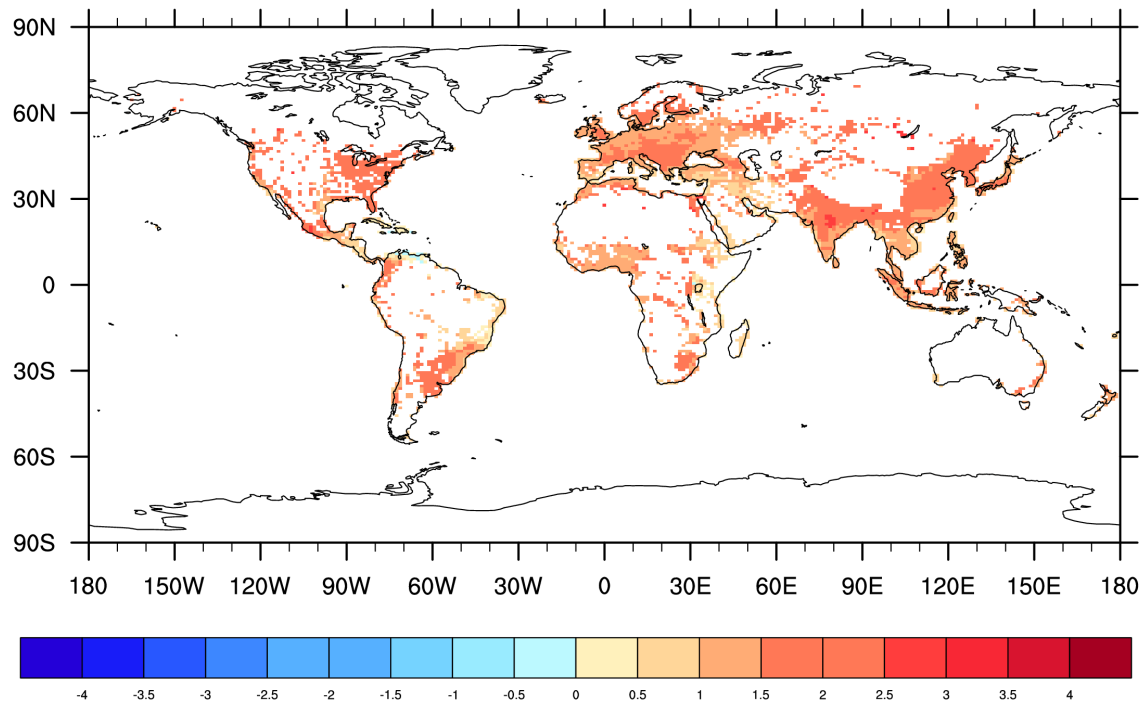
GPP





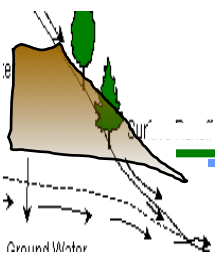
Urban Heat Island in CCSM4

Present day Urban Heat Island (UHI) simulated by CLM
Urban ($^{\circ}\text{C}$)

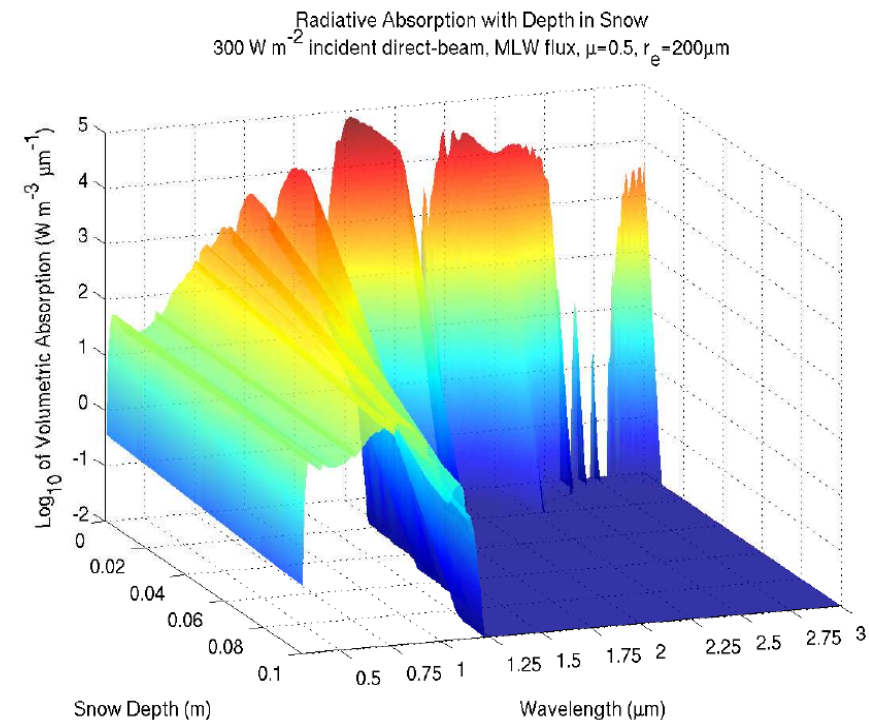
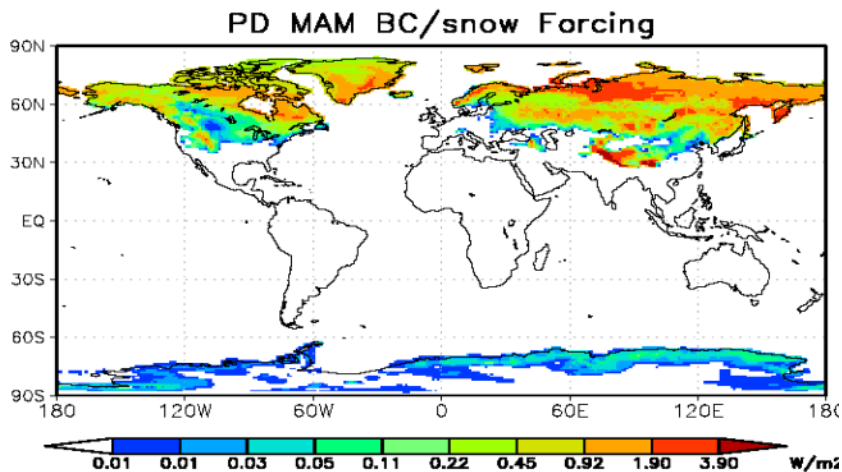


Modeled UHI ranges from near-zero up to 4°C with spatial and seasonal variability controlled by urban to rural contrasts in energy balance.

Snow, Ice, and Aerosol Radiative Model (SNICAR)



- Snow darkening from deposited black carbon, mineral dust, and organic matter
- Vertically-resolved solar heating in the snowpack
- Snow aging (evolution of effective grain size) based on:
 - Snow temperature and temperature gradient
 - Snow density
 - Liquid water content and
 - Melt/freezing cycling



Flanner et al (2007), *JGR*
Flanner and Zender (2006), *JGR*
Flanner and Zender (2005), *GRL*