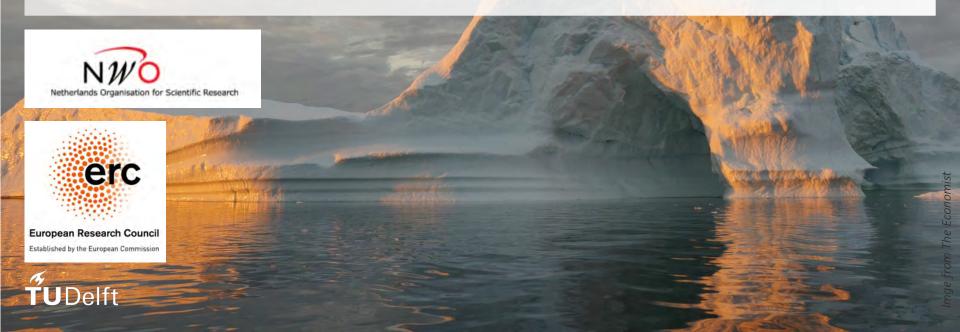
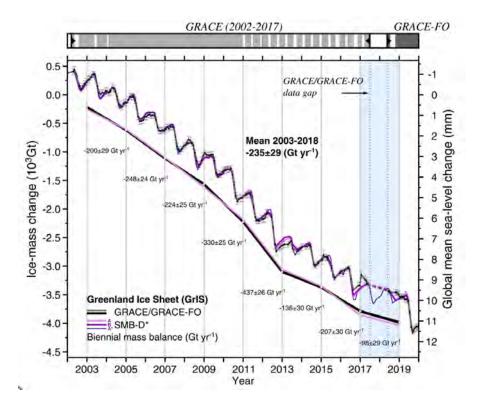
Future complete deglaciation of the GrIS in CESM-CISM2 under high greenhouse gas forcing

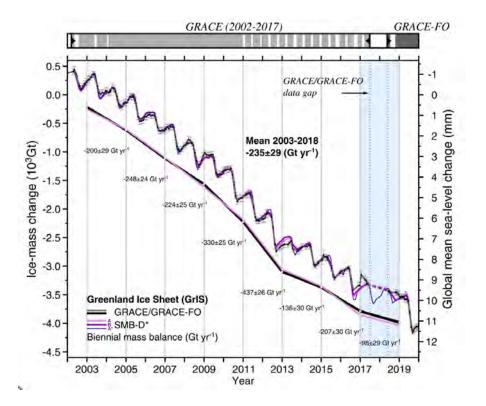
Raymond Sellevold, **Miren Vizcaino**, Michele Petrini, Sotiria Georgiou Department of Geoscience and Remote Sensing, TU Delft





- The GrIS has been losing mass since the early 90s, current contributing to sea level rise with 0.7 mm/yr
- Record mass loss in 2019

Sasgen, Wouters et al., Communications Earth & Environment (2020)

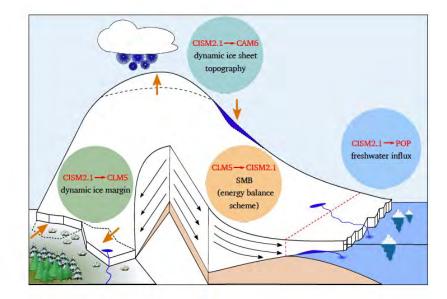


- The GrIS has been losing mass since the early 90s, current contributing to sea level rise with 0.7 mm/yr
- Record mass loss in 2019
- How (fast) will the ice sheet melt under scenarios of high greenhouse gas forcing?

Sasgen, Wouters et al., Communications Earth & Environment (2020)

CESM2-CISM2:

- **Coupling** of ice dynamics, SMB & climate
- Advanced, interactive SMB simulation: in the land component of CESM2 with explicit albedo & refreezing calculation
- Relatively **advanced ice flow** (HO approx.) simulation for a coupled simulation
- **High resolution** (4km/1 deg) for a coupled ice/climate simulation



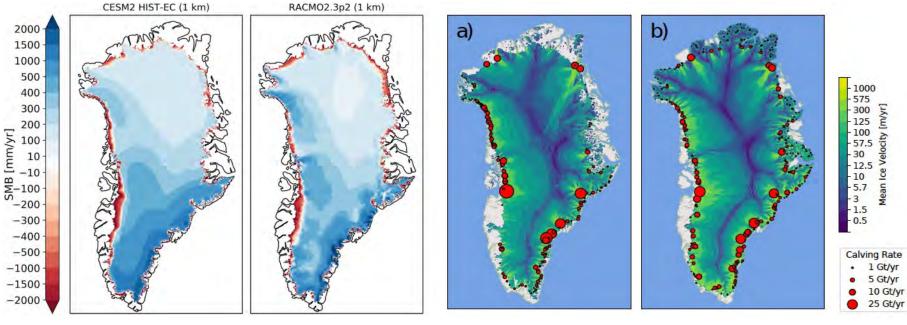
Conclusions

Muntjewerf et al, JAMES, 2021

Method

Results

Conclusions



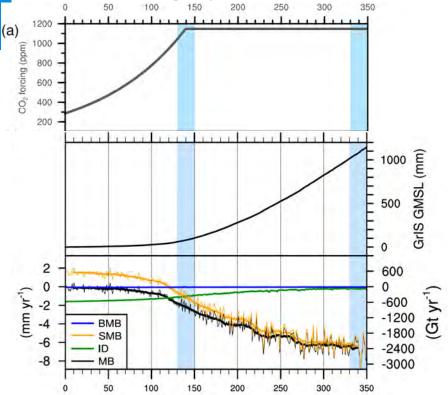
Observations (Joughin et al.; Enderlin et al, 2014)

CESM2-CISM2

Extending 1pct to 4xCO2 simulation (Muntjewerf et al, JAMES, 2021)

Results

Method



Question

Background

Cumulative SLR contribution of
 107 mm (131-150)* & **1140 mm** (331-350)

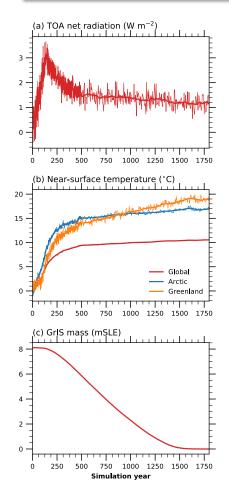
Conclusions

• 20% area reduction

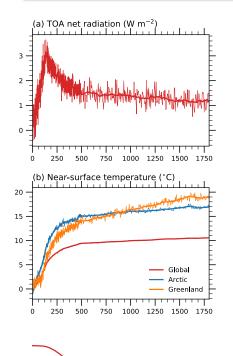
*similar to 2080-2100 under SSP5-8.5 (*Muntjewerf et al, GRL, 2021*

Results

- Climate
- Mass loss rates & pattern
- Precipitation evolution
- Runoff evolution



- 1:5 years acceleration of ice sheet model from year 500
- High warming by end of simulation (>10 K), TOA net radiation > 1 W/m2
- GrIS warming < Artic warming until year 800, when GrIS is substantially smaller



Climate & deglaciation timescale

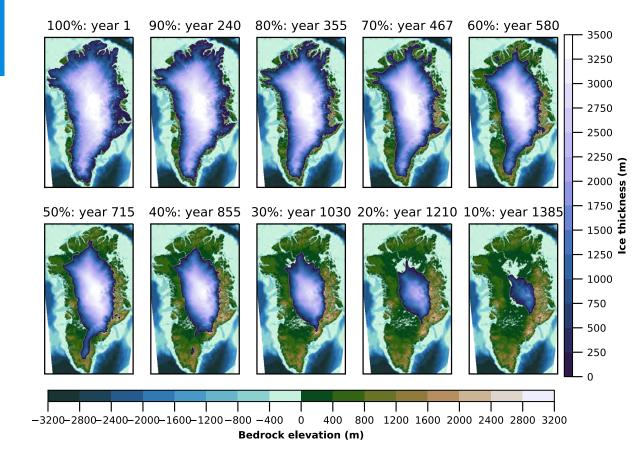
- 1:5 years acceleration of ice sheet model from year 500
- High warming by end of simulation (>10 K), TOA net radiation > 1 W/m2
- GrIS warming < Artic warming until year 800, when GrIS is substantially smaller
- Fast deglaciation in <1,700 years

Global Climate

Mass loss

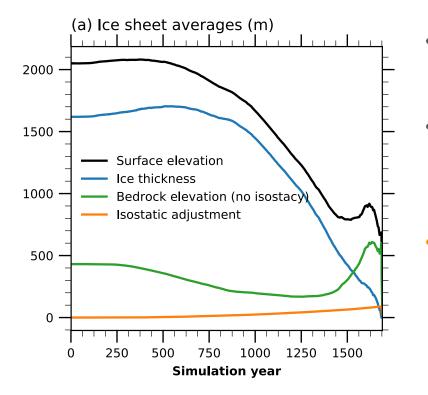
Precipitation

Melt & runoff



- 50% area reduction by year 715
- Southern Part deglaciates first

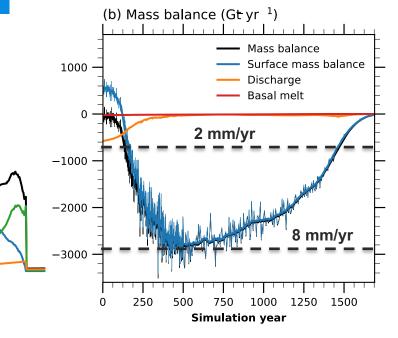
elevation change & contributors



Mean thickness slightly increases 250-600 as relatively thin margin retreats, then decreases Ice sheet retreats from some relatively high bedrock at the margins first, and to high bedrock in the Mid-East in the last centuries Smal dontributi isostasy (+10)

Melt & runoff

Mass balance = SMB – Discharge – Basal Melt

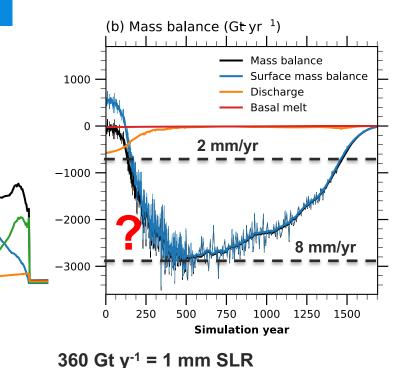


360 Gt y⁻¹ = 1 mm SLR

- SMB decreases from ~600 Gt/yr to 0 within the first 110 years
- Ice discharge decreases (negative contribution to sea level rise) in response to marginal thinning and retreat from increased melt
- By 140, sea level contribution is 2 mm/yr
- This increases rapidly in the following two centuries, reaching a maximum contribution of 8 mm/yr

Melt & runoff

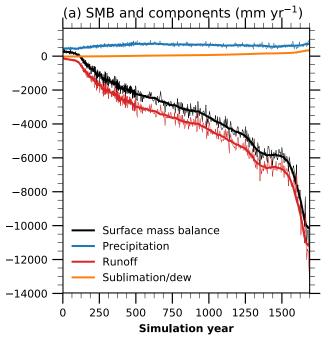
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What causes this strong acceleration?

SMB = Precipitation - Runoff - Sublimation



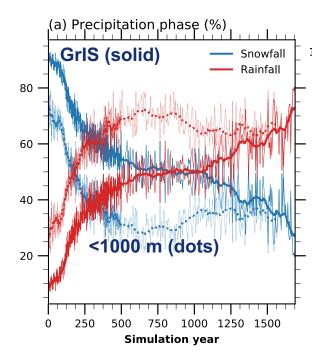
Evolving GrIS mean values

•	We normalize by area	(kg m ⁻² yr ⁻¹)
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- SMB becomes negative by year 110
- Rupoff dominates SMB evolution, with larger increases 120-500 and at end of simulation
- A period without increase 1300-1500

Component	Pre-industrial	Years 131–150 Absolute		Years 331–350 Absolute	Anomaly
			Anomaly		
SMB [4 km]	585 [85]	-367 [166]	-952	-2259 [357]	-2844
SMB [1°]	544 [103]	-521 [217]	-1065	-2589 [442]	-3133
Precipitation	846 [83]	986 [97]	140	1122 [97]	276
Snowfall	780 [80]	750 [74]	-30*	683 [71]	-97
Rain	72 [12]	235 [38]	163	439 [59]	367
Refreezing	223 [54]	693 [73]	470	534 [43]	311
Melt	415 [92]	1,914 [251]	1499	3,804 [443]	3389
Sublimation	45 [4]	50 [6]	5	3 [11]	-42

Melt & runoff

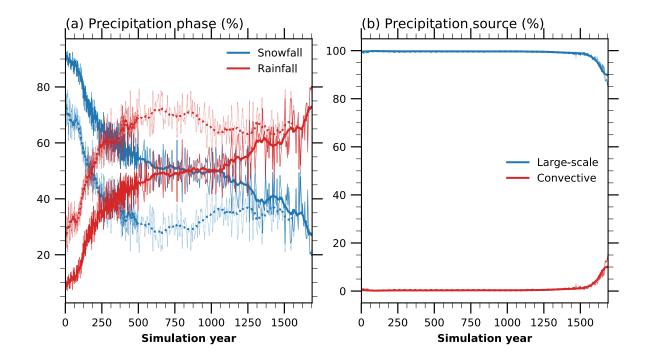


- From 90% snowfall to 30% by end of simulation
 - 50% by 850



Precipitation

Melt & runoff

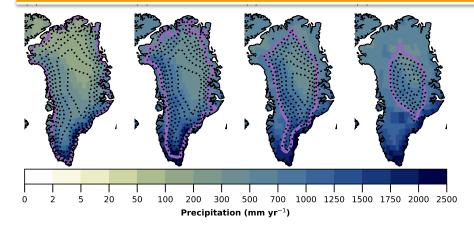


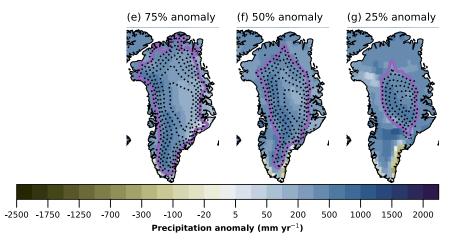
Global Climate

Mass loss

Precipitation

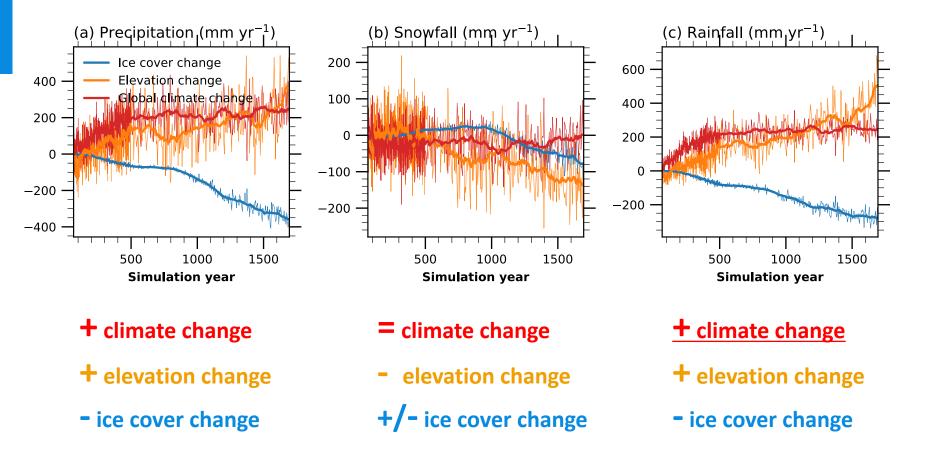






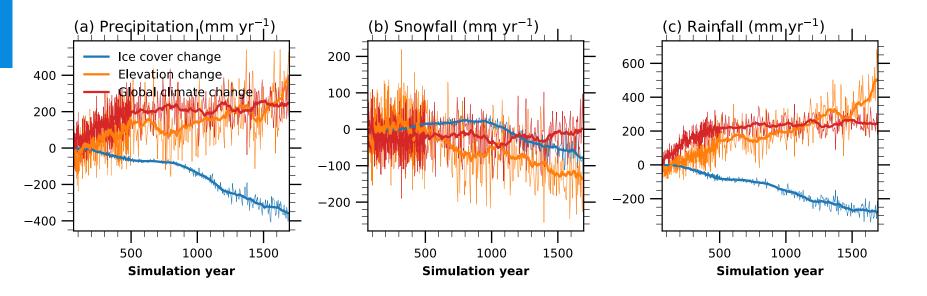
- Precipitation increases in response to climate change
- Precipitation maxima follows
 Western margin
- Precipitation reductions over deglaciated areas in the SE

Precipitation



Precipitation

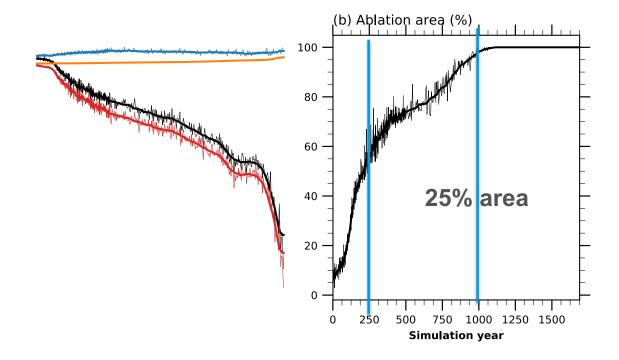
Melt & runoff



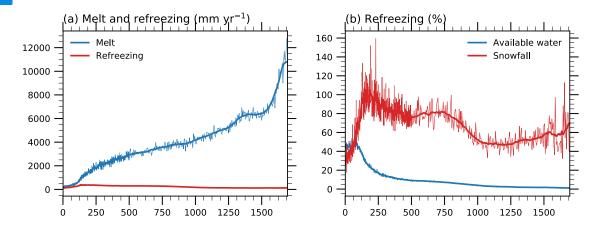
- Snowfall does not increase, and decreases by end of simulation
- Rainfall increases both due to climate and elevation change
- Precipitation partially decreases as the GrIS retreats from North Atlantic area

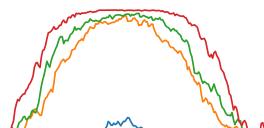
Precipitation

Melt & runoff

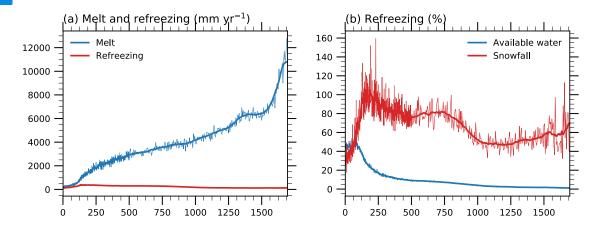


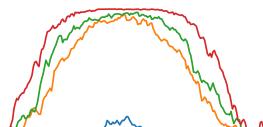
Runoff=Melt+Rain-Refreezing



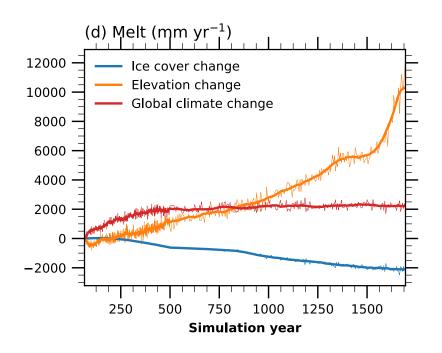


Runoff=Melt+Rain-Refreezing



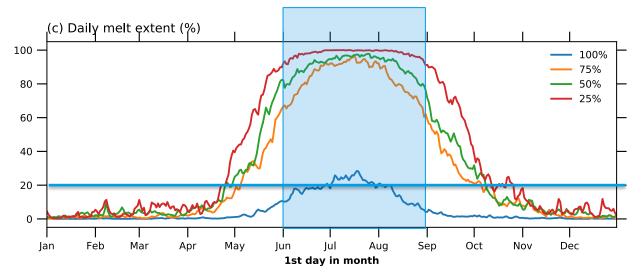






 Elevation contribution equals climate contribution at mid simulation

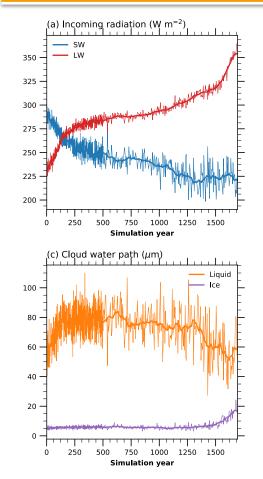




• By 400 most of ice sheet melts in July

A LANDAR AND A LAND A LANDAR AND AND AND ALL

- By 1100 most of ice sheet melts in June-July August
- Melt season extends towards the Fall as deglaciation progresses



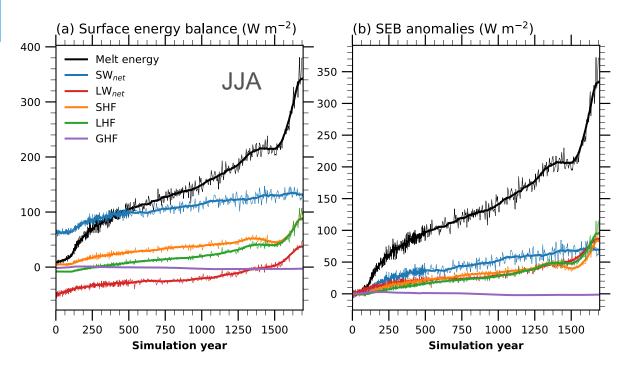
 Incoming SW decreases due to clouds and surface albedo reduction

- LW increases similarly to near-surface temperature, due to climate and elevation change
- Cloud liquid water path increases with climate, but (likely) decreases with elevation and areal change (less maritime areas)
 Increase in ice water path in the last two centuries

Global Climate

Mass loss

Melt=Swnet+Lwnet+SHF+LHF+GHF

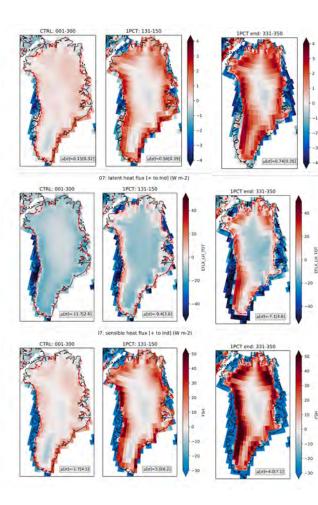


- LW_{net} dominates until 120, then Swnet due to albedo feedback
- SHF & LHF largely increase from 120, and each contribute similarly to Lwnet between 500-1400
- Small decrease in SHF contribution around 1500
- Temperature-related fluxes dominate during the last century

T2m-Tsurface (K)

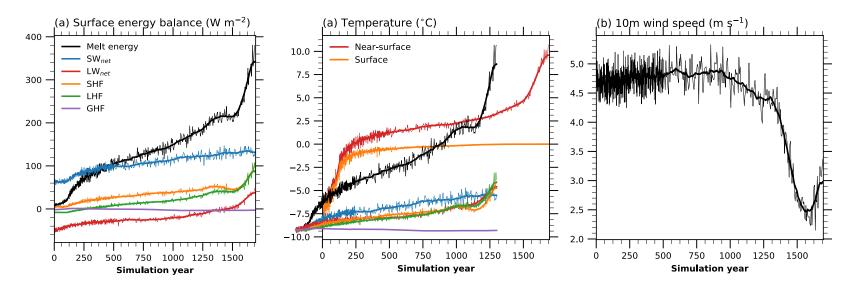
Latent heat flux (Wm⁻²)

Sensible heat flux (Wm⁻²)

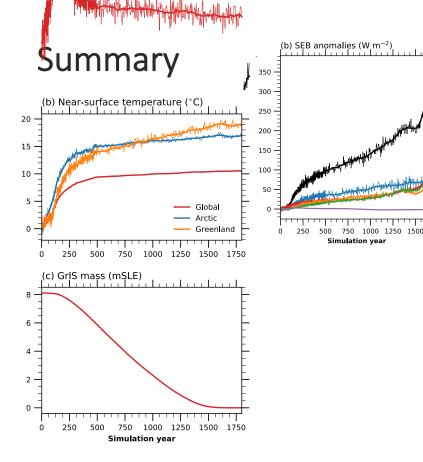


- Strong surface
 temperature inversion
 as ablation areas
 expand
- Results in much increased latent and sensible heat flux

Precipitation



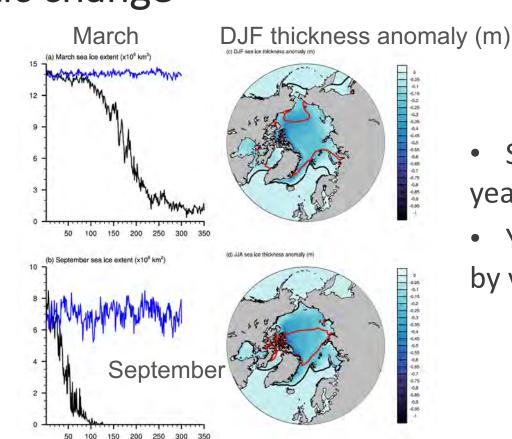
- Large increase in SHF & LHF around 120 relates to large areas reaching melt conditions and associated increasing near-surface temperature inversion
- Decrease in SHF around 1500 related to decreased wind speed



Vizcaino et al, in preparation Sellevold et al, in preparation

- 5 K (140), 8 K (350), 10 K (1750)
- Fast deglaciation: <1,700 yrs
- Fastest margin retreat from SW, then North
- Snowfall rates are similar to preindustrial during most of simulation
- Large melt acceleration from + albedo feedback and much increased turbulent heat fluxes
- Refreezing increases with initial melt increase, up to snowfall accumulation rate, then declines
- Large + elevation-melt feedback
- Small melt-isostasy feedback

Additional slides



Arctic change

- Seasonally ice-free by year 90
- Year-around ice-free by year 300

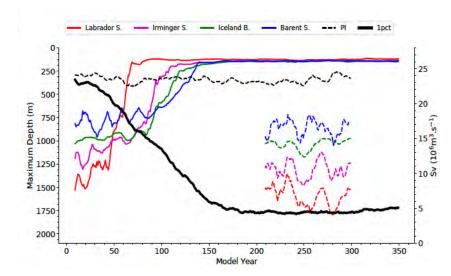
JJA thickness anomaly (m)

30

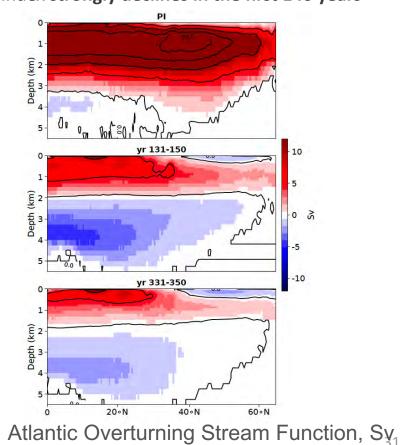
Mixed layer depth declines in all basins

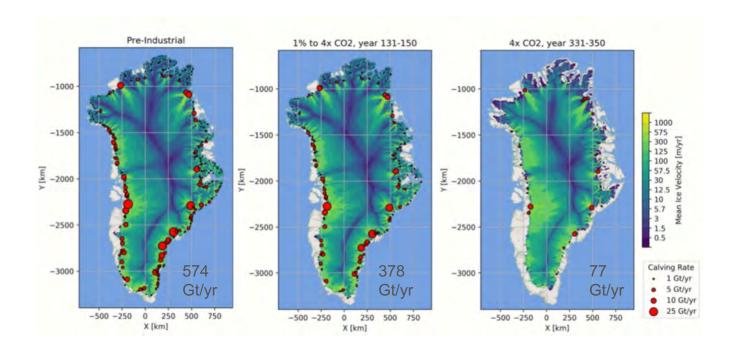
NAMOC weakens

NAMOC index strongly declines in the first 140 years



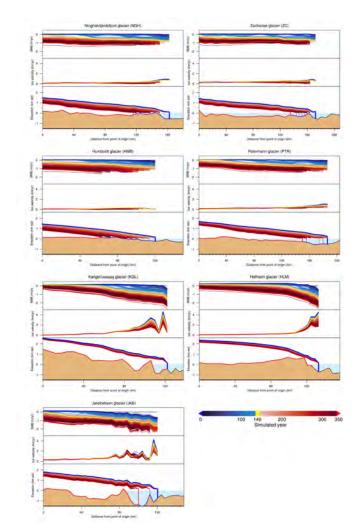
Mixed Layer Depth in individual basins (m), and NAMOC index (Sv, maximum of overturning stream function)

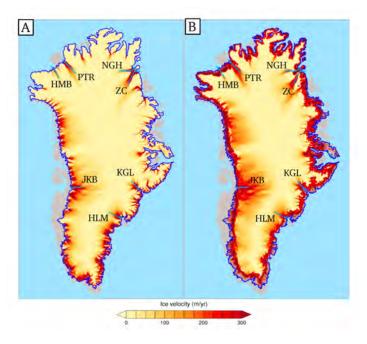




Marginal thinning (-15% volume) & retreat (-20% area) result in an almost land-terminating ice sheet by years 331-350







Basin	Glacier	Terminus position. at year 350 (km)	Start retreat (year)	Transition to land margin (year)
NE	Nioghalvfjerdsfjord	-46	159	N/A
NE	Zachariae	-50	180	N/A
NO	Petermann	-36	246	N/A
NO	Humboldt	-60	184	311
SE	Kangerlussuaq	0	-	N/A
SE	Hellheim	0		N/A
CW	Jakobshavn	-20	271	N/A