

Climate applications of VR-CESM: from glaciers to convection (using CAM-SE)

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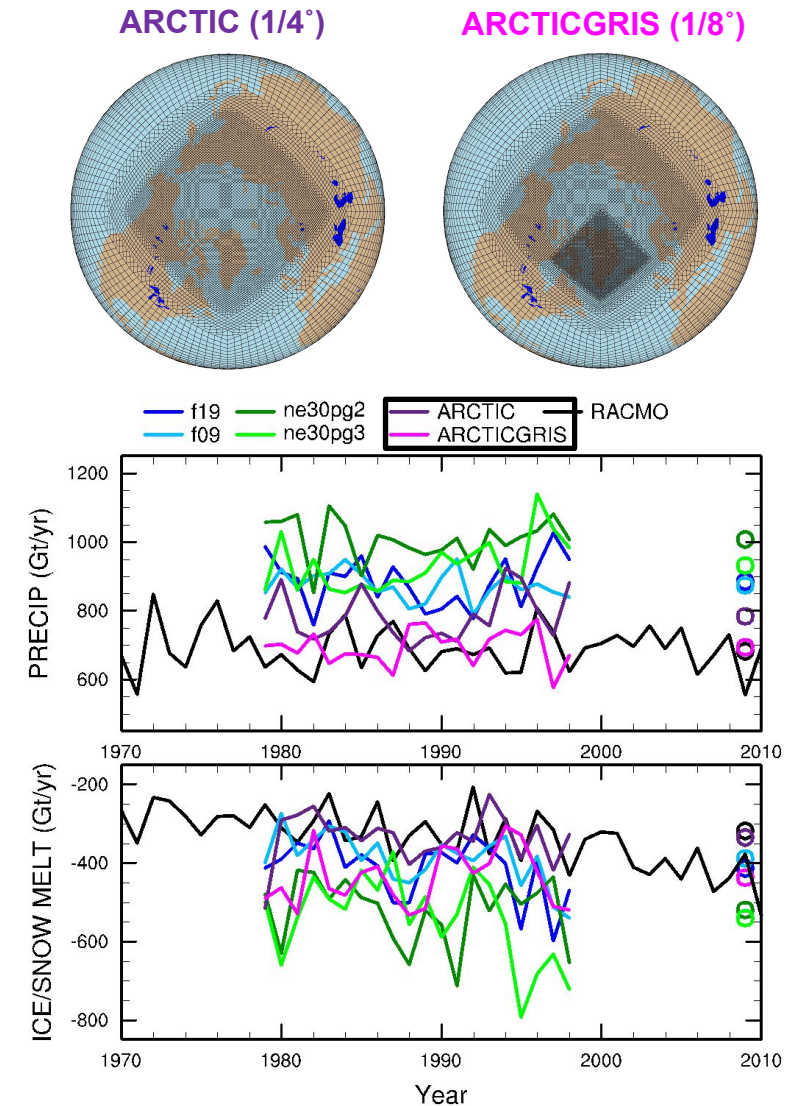
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Greenland Surface Mass Balance: AMIP Experiments

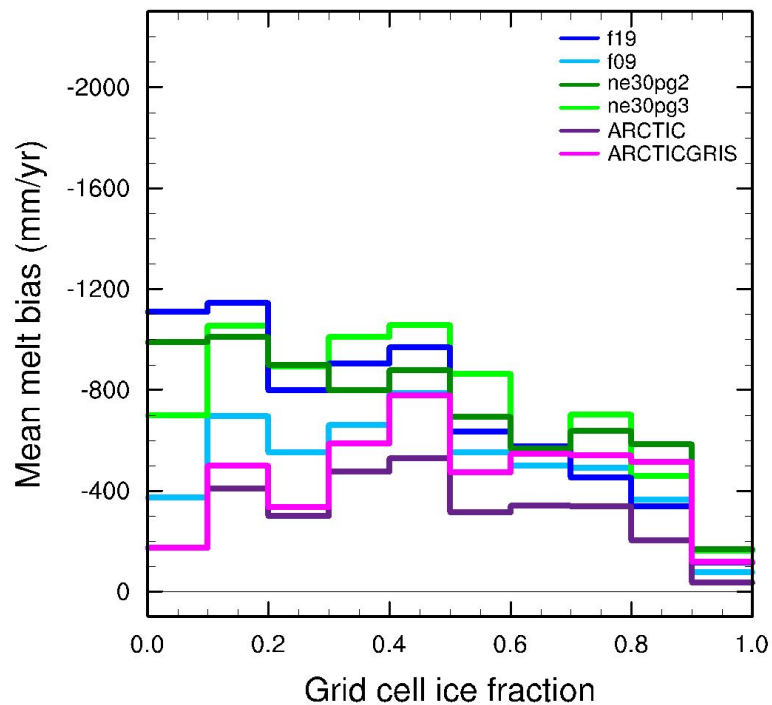
The Greenland Ice Sheet (GrIS) is an important component of the Earth System, but it is challenging to resolve at 1°

- Large improvements in simulated SMB at $\frac{1}{4}^\circ$ (van Kampenhout et al. 2019)
- Began supporting out-of-the-box functionality for two Arctic grids in CESM2.2 (Herrington et al. 2022)
- Coupled to POP2/CISM2 and run over idealized warming scenarios (Ziqi Yin's talk Wed. at 11am)



Problems with mixed land surface types: Greenland

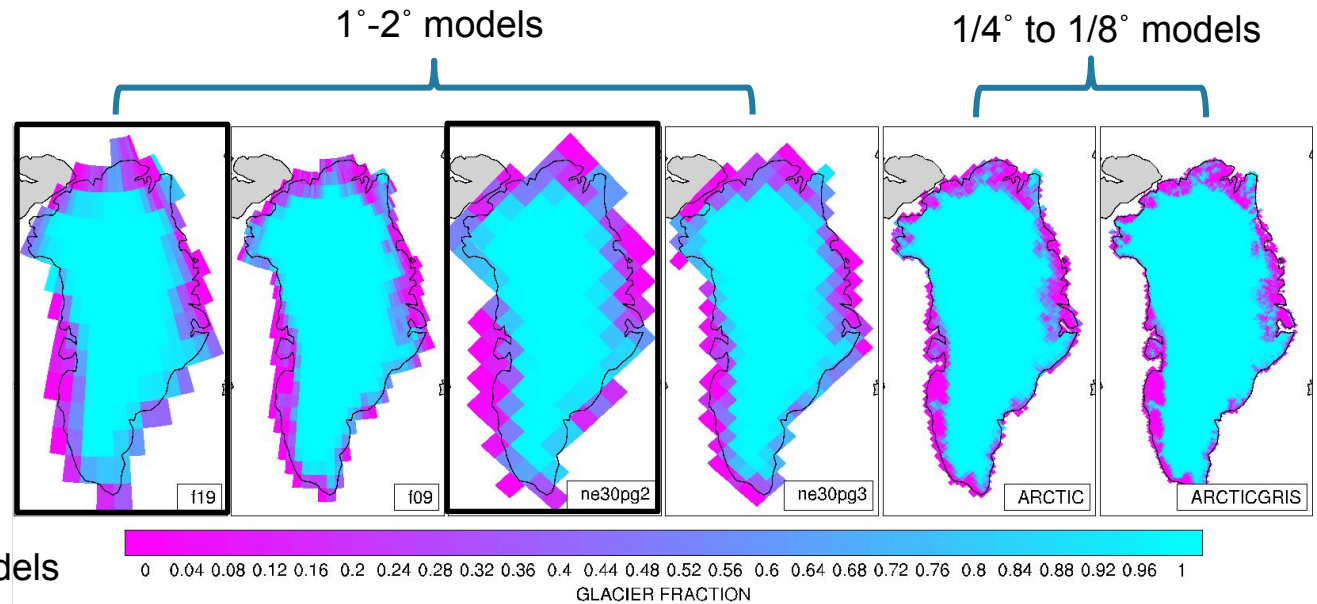
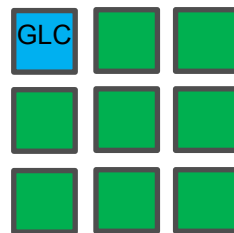
- Map climo melt bias map in all grids, to the coarsest grids (f19/ne30pg2)
- Bin melt bias by grid cell ice fraction



1°-2° models



1/4° to 1/8° models

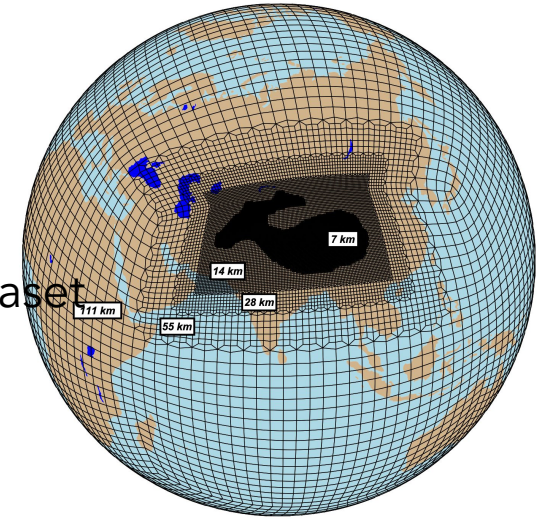


- Smaller grid cell ice fractions have larger (more negative) melt biases
- 1/4° and 1/8° alleviates this dependence of melt bias on ice fraction, because it largely resolves the ice margin

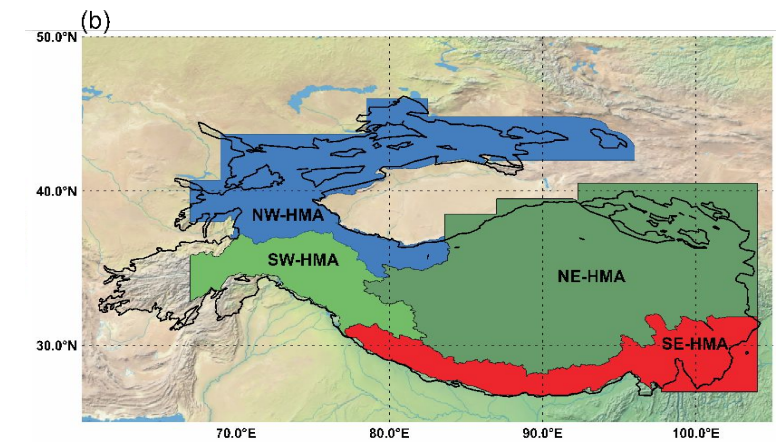
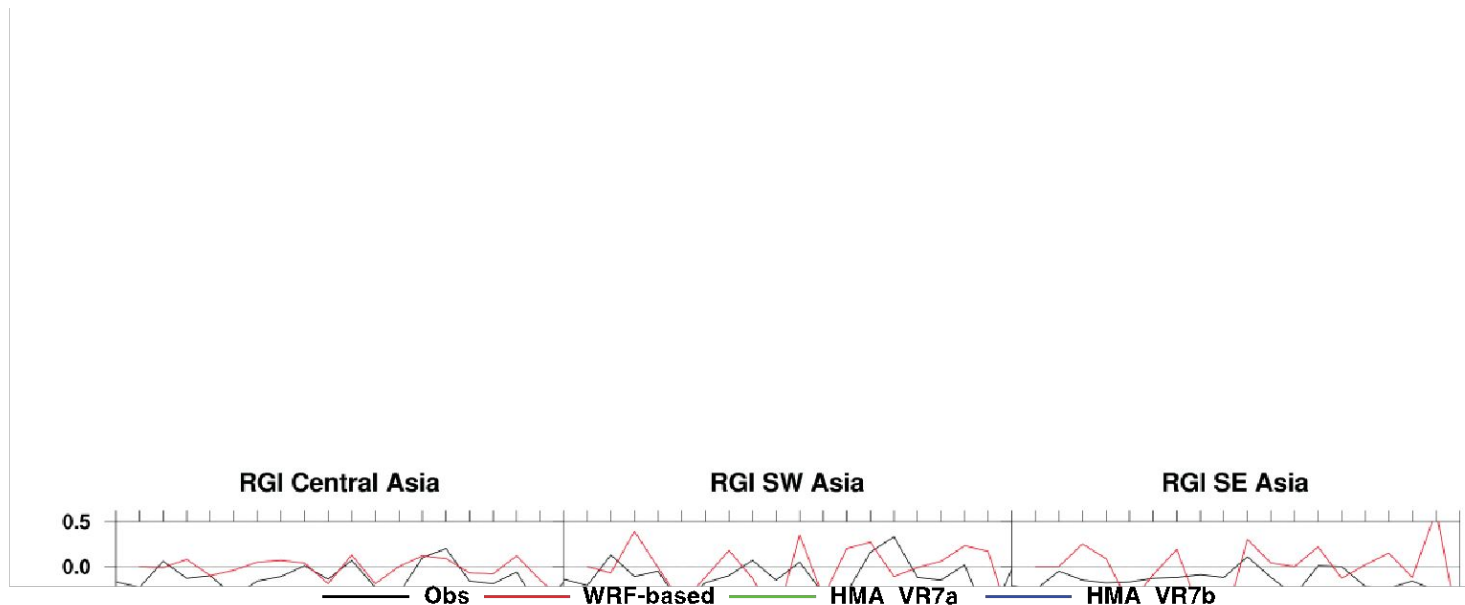
Simulating glacier SMB over High Mountain Asia (HMA)

Two coupled land-atmosphere simulations for the period 1979-1998 (HMA_VR7a,b)

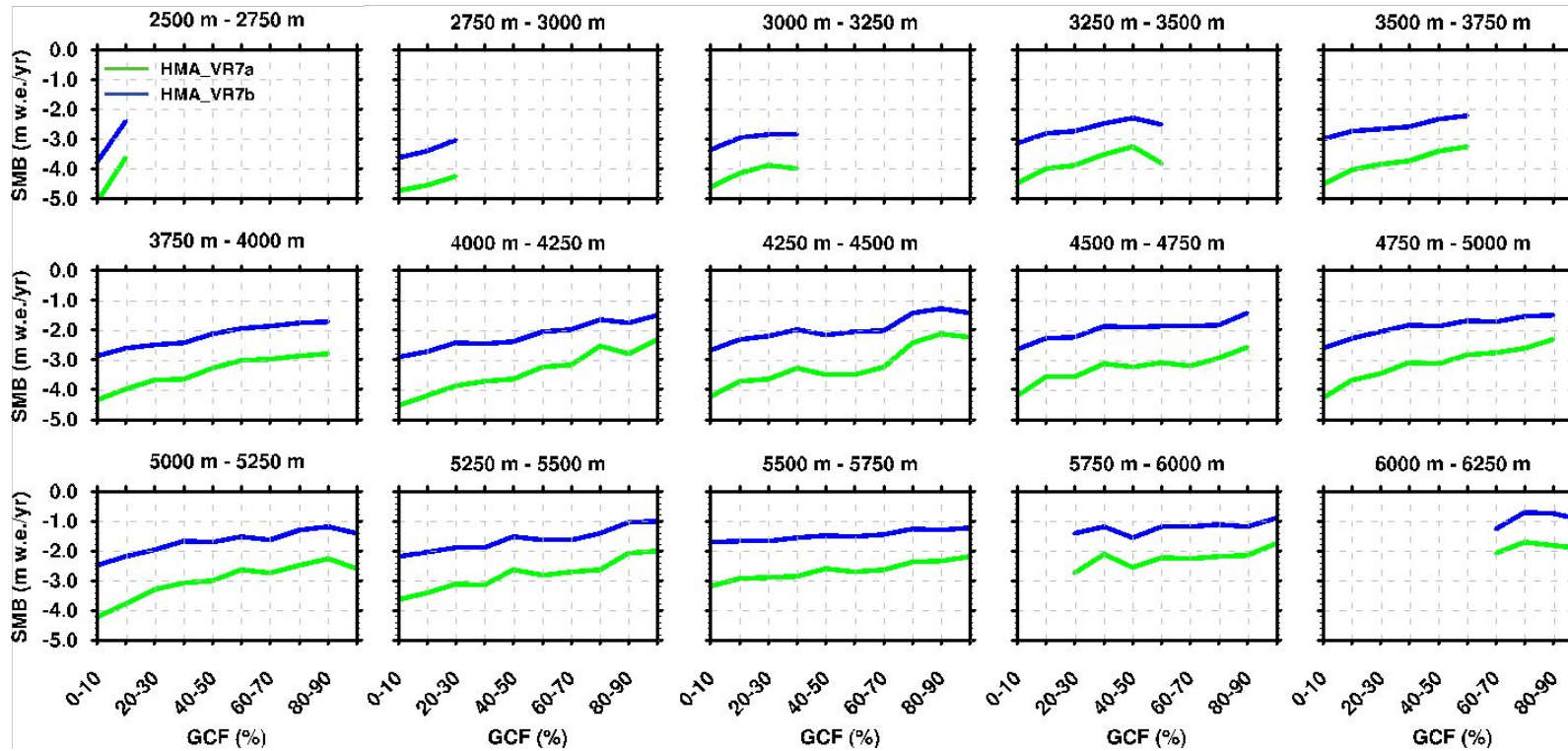
- The second simulation (HMA_VR7b) is performed with updated glacier-cover dataset and includes snow and glacier model modifications.
- SMB is underestimated compared to observations (black)
- In HMA_VR7b simulations (blue), SMB improves, but is still underestimated.



High Mountain Asia VR grid.
Regional refinement up to 7 km.
Model cost: ~90k p. sim. yr.



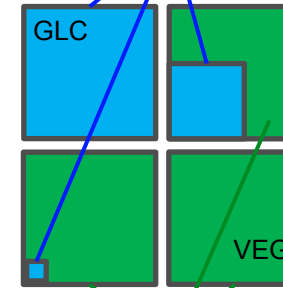
Simulating glacier SMB over High Mountain Asia



HMA CLM Land Surface +
Elevation Downscaling



Glacier Land Unit:
Elevation Downscaling
36 ECs

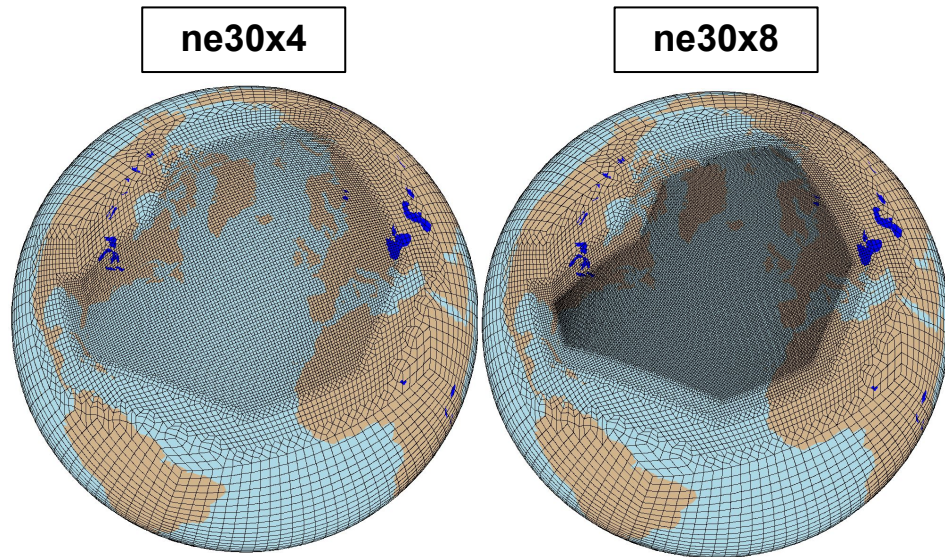


Vegetated Land Unit:
No Elevation Downscaling

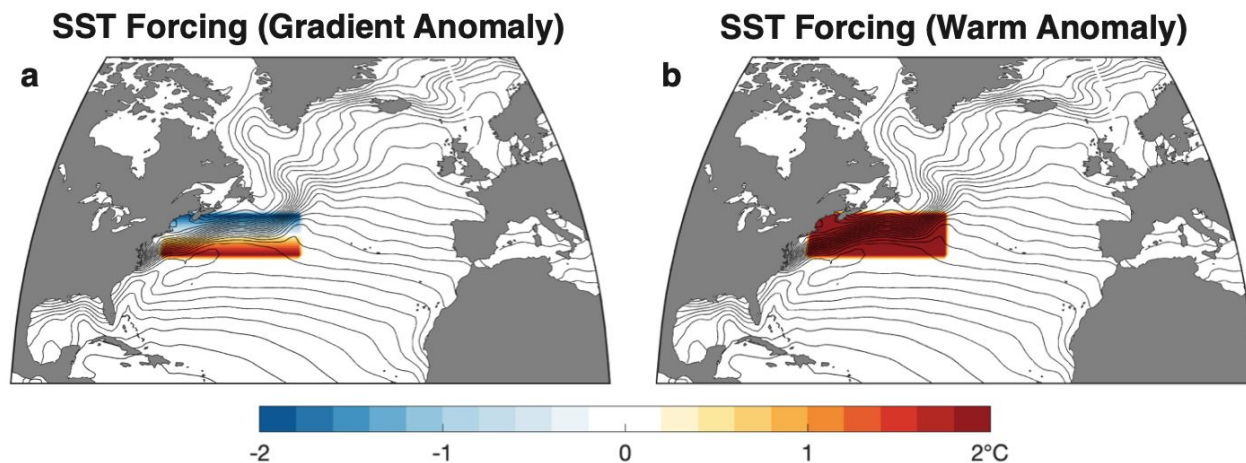
SMB declines with decreasing glacier fraction

- SMB bias partly due to homogeneous land-atmosphere coupling over mixed land surface types and partly due to exclusion of elevation downscaling over vegetation.
- Other explanations SMB bias: reduced cloud cover/warm summer temperature biases.

North Atlantic Grid (1/4° and 1/8°)



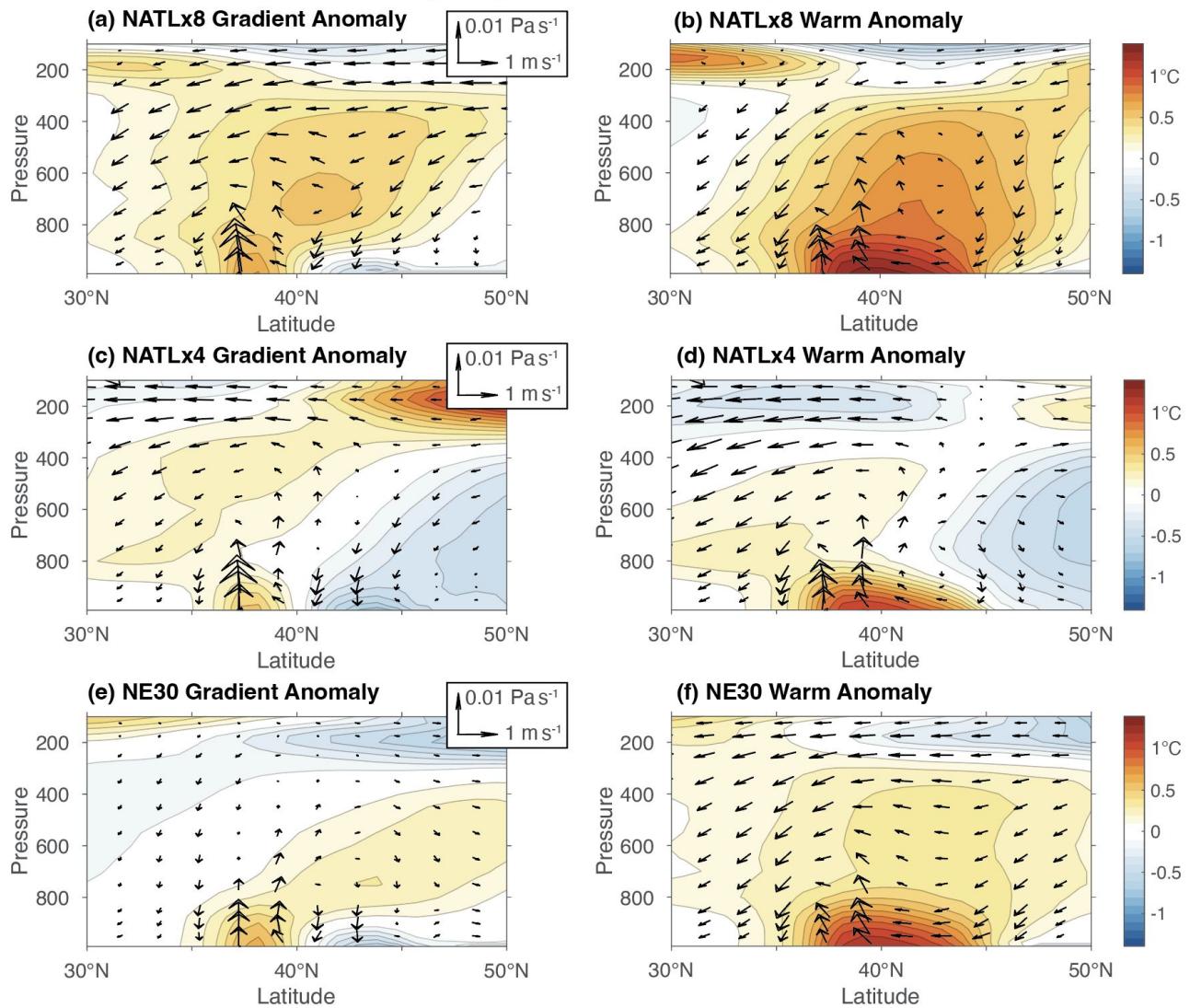
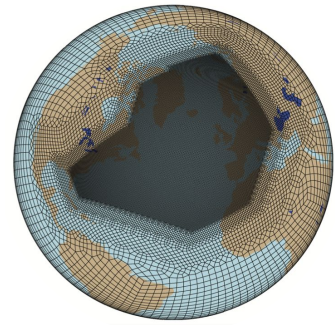
- Reference: Perpetual year 2000-like* seasonally varying (1° resolution) SSTs *1995-2005 climatology
- Experiments: Two different SST anomaly patterns in the Gulf Stream
- Each simulation run with the VR-NATL grid and with a 1° reference grid (NE30)



Results show averages over 30 years (VR) or 50 years (NE30)

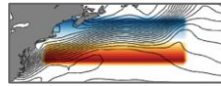
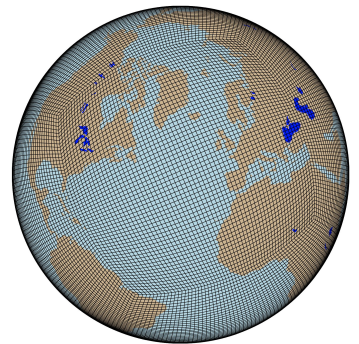
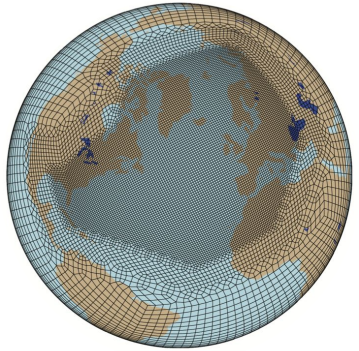
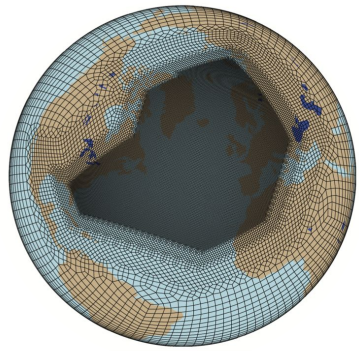
Deeper warm anomaly in forcing region only at 1/8°

DJF Forcing Longitudes v , ω , and θ Response



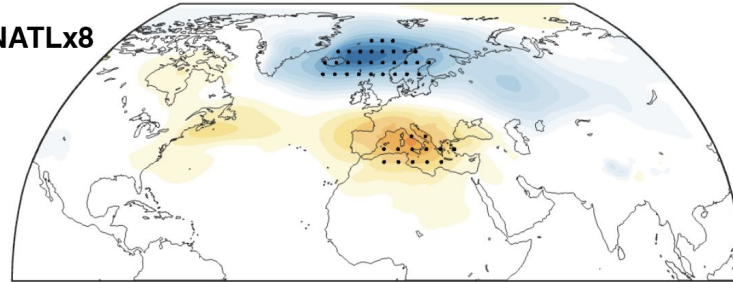
- Updrafts associated with cold-fronts are permitted at 1/4°, but updrafts in warm front aren't permitted occur until 1/8°
- Additional warm front ascent → greater poleward eddy transport

NAO-like large-scale circulation response

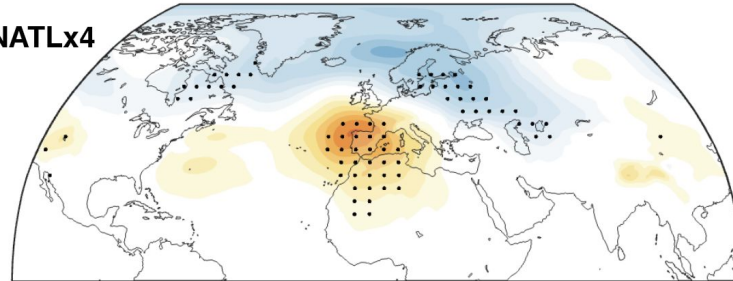


DJF SLP Response (Gradient Anomaly)

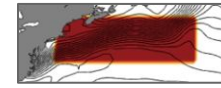
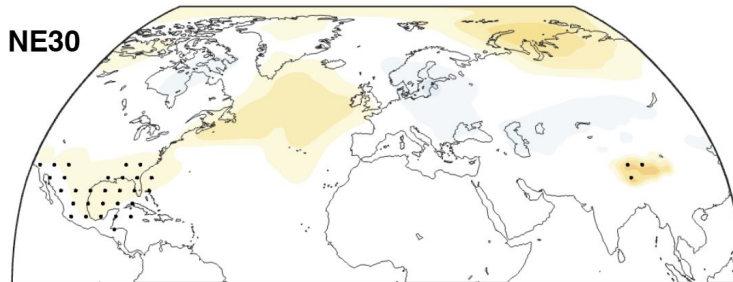
NATLx8



NATLx4

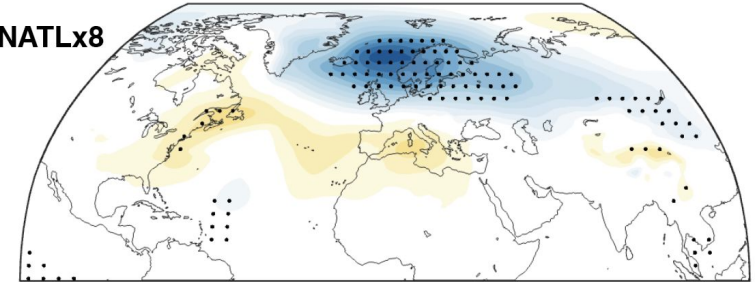


NE30

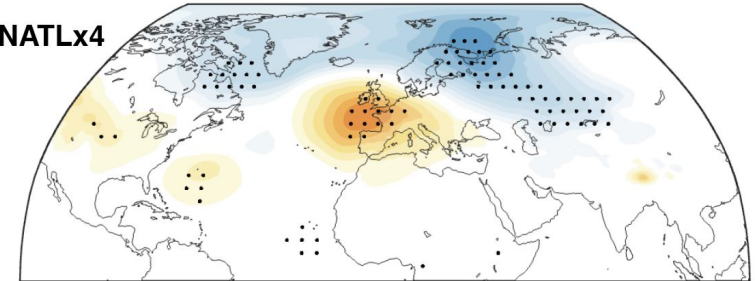


DJF SLP Response (Warm Anomaly)

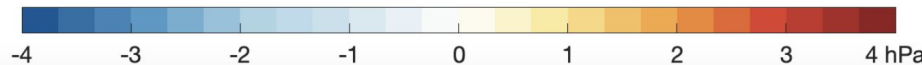
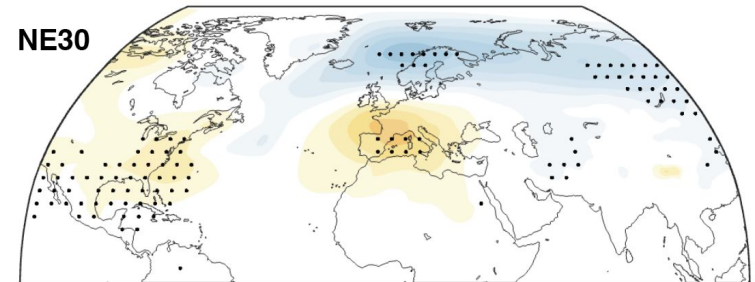
NATLx8



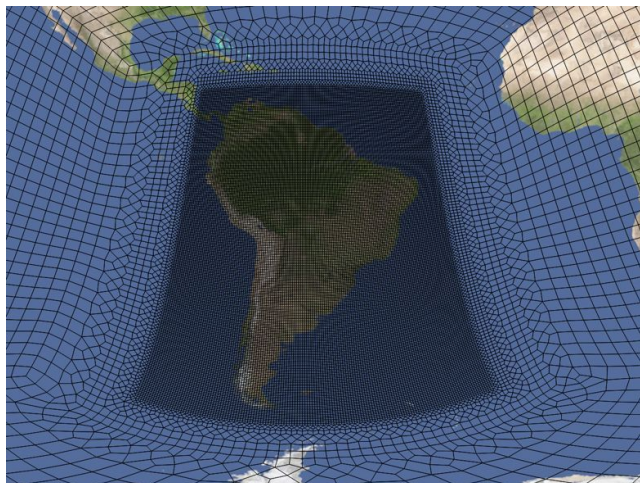
NATLx4



NE30



Variable Resolution: South American (SAM) Domain



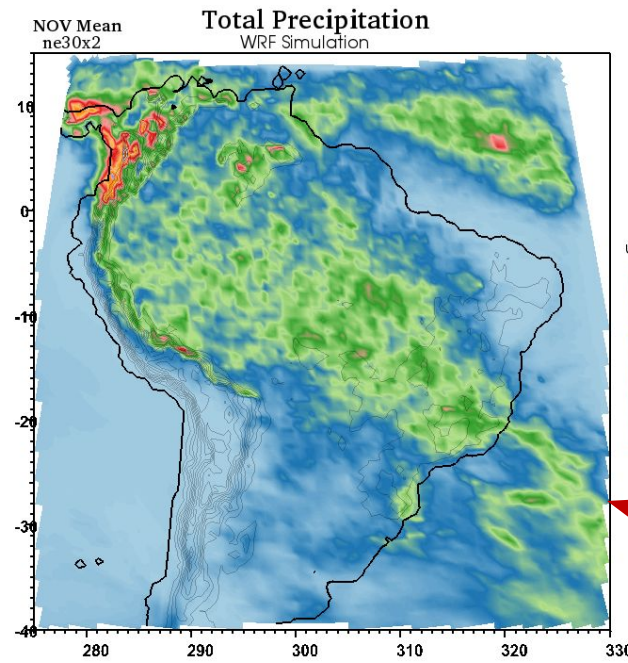
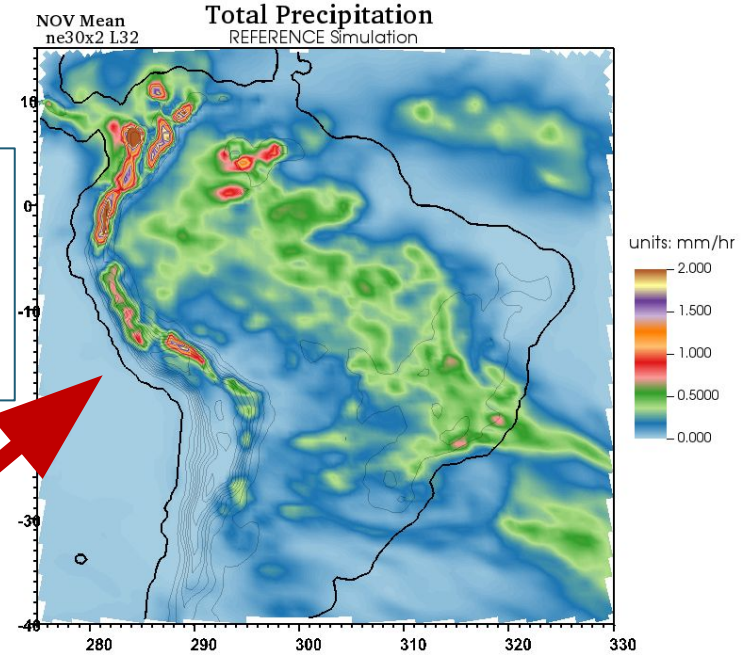
SE Refined resolution	Nominal Grid resolution	Horizontal Grid points
ne30	100km	48,602
ne60	50km	64,766
ne120	25km	116,534
ne240	12km	337,503
ne480	6km	1,167,113

The Domain:
A Hierarchy of grid resolutions coincident with a 4-km WRF domain

For the SE dycore, the erroneous precipitation over the Andes is caused by the applied hyper-viscosity damping.

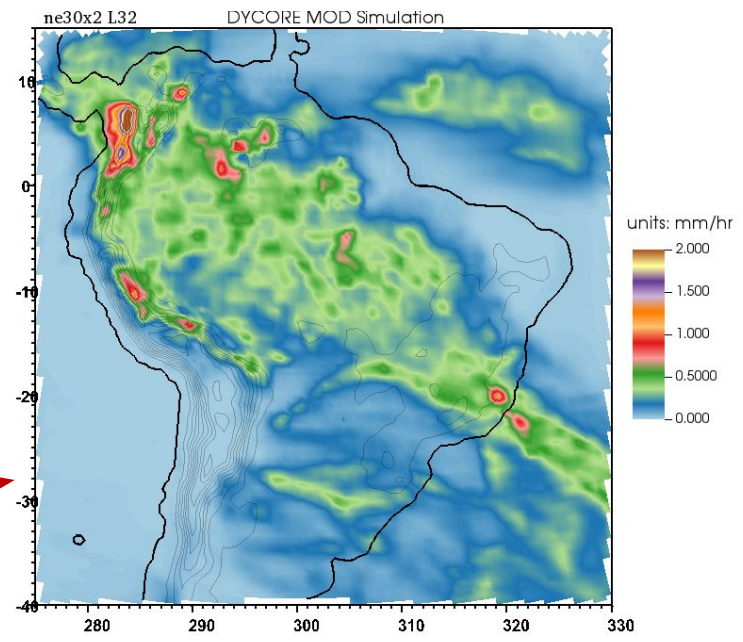
Replacing it with Conservative Damping consistent with terrain following model surfaces:

- The dry halo and heavy precipitation over the ridge has been eliminated.
- The precipitation better resembles the diurnal variations for upscaled WRF results.



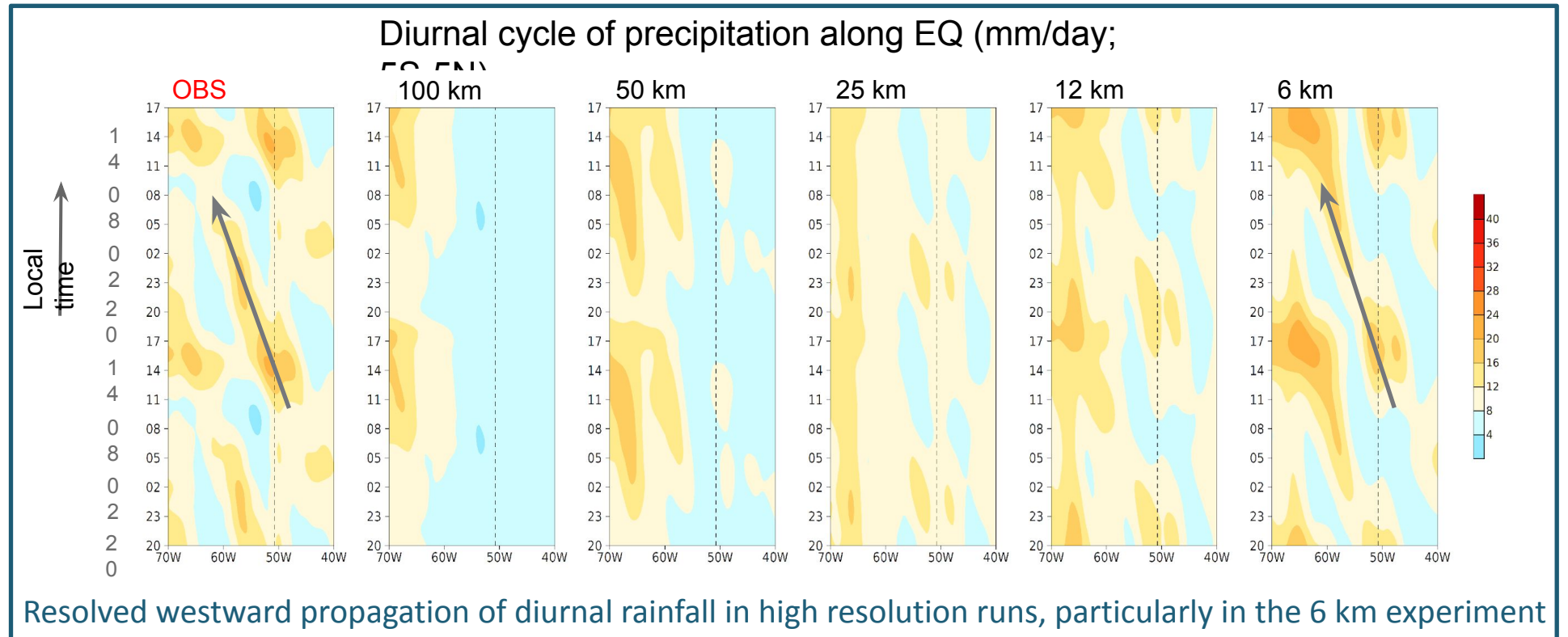
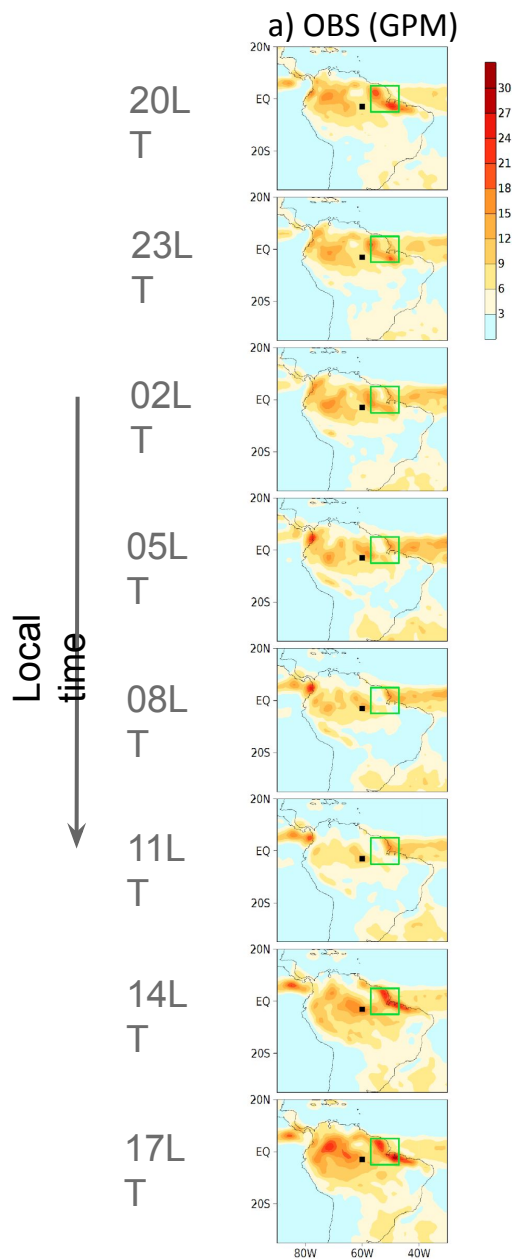
WRF
NOV 2010 Mean for 50km resolution

CESM
NOV 2010 Mean for 50km resolution



Climatological diurnal cycle of rainfall in east Amazon (March-May)

The observed precipitation over the east Amazon is closely associated with diurnal westward propagation rainfall events, but is poorly simulated in CAM6 at 1°



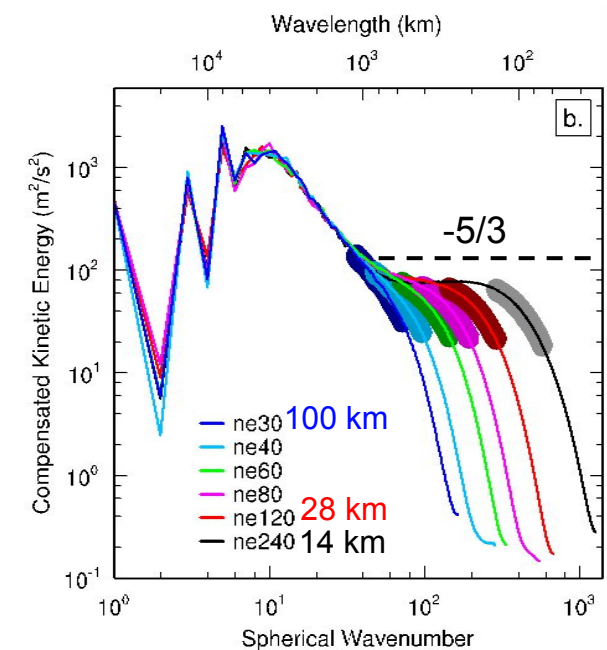
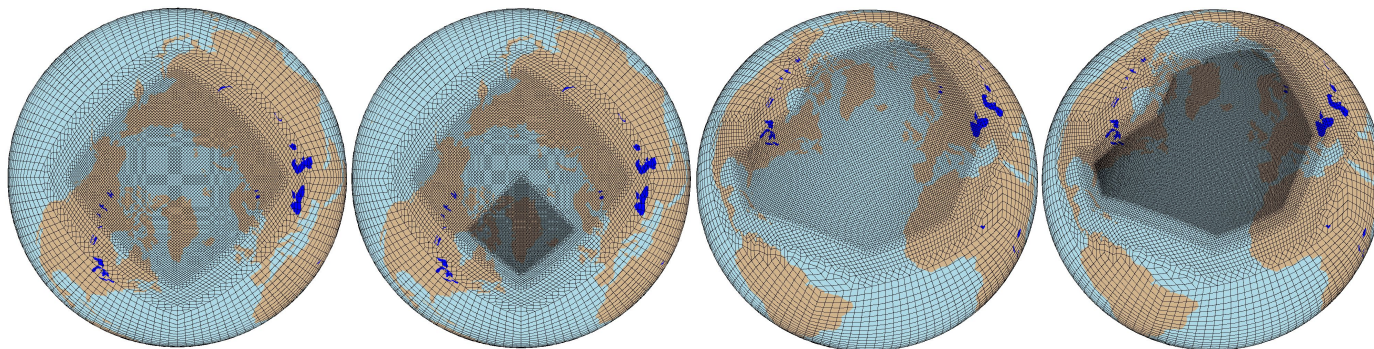
Variable-resolution CAM-SE milestones down to 7 km resolution

More realistic boundary conditions

- 28 km – Realistic topography and orographic precipitation (ice sheets and mountain ranges)
- 28 km – Realistic land surface-type (resolves ice sheet margins, but *not* mountain glaciers)
- 14 km – Realistic surface emissions for air-quality studies (concentrated sources over urban centers)

Permitting meso-scale motions and convection

- 28 km – updraft permitting – ‘elements’ of TCs, cold-fronts and precipitation extremes
- 14 km – meso-scale permitting – meso-scale ascent in warm fronts
- 7 km – convection permitting – propagating MCSs and the diurnal cycle over land



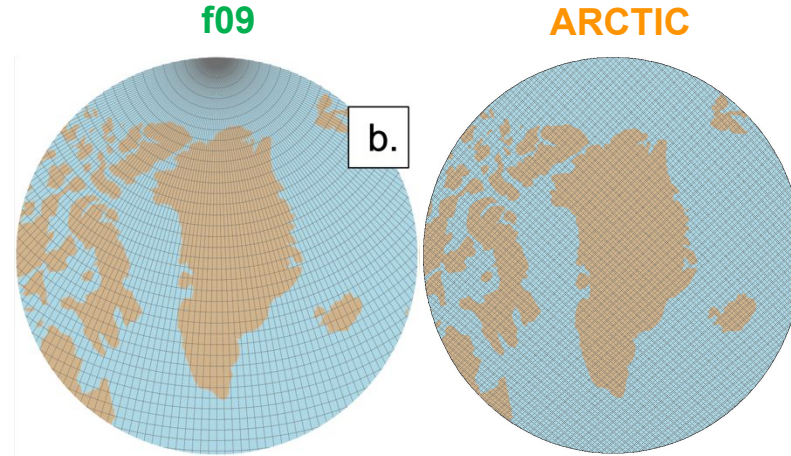
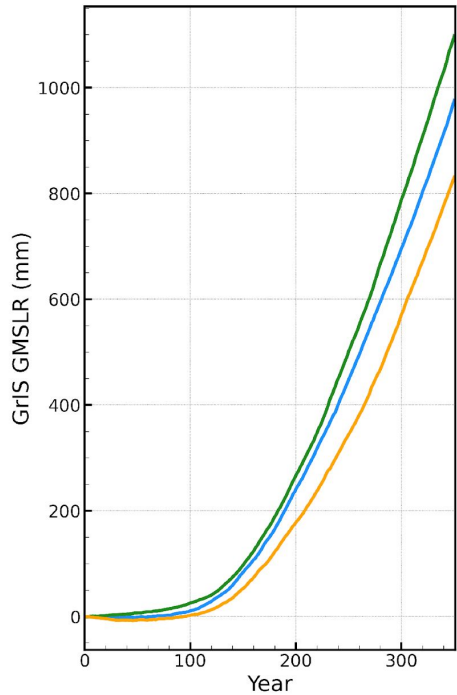
14 km CAM-SE resolves a significant chunk of the meso-scale spectrum

ARCTIC 1/4° VR grid: Idealized Warming Experiment

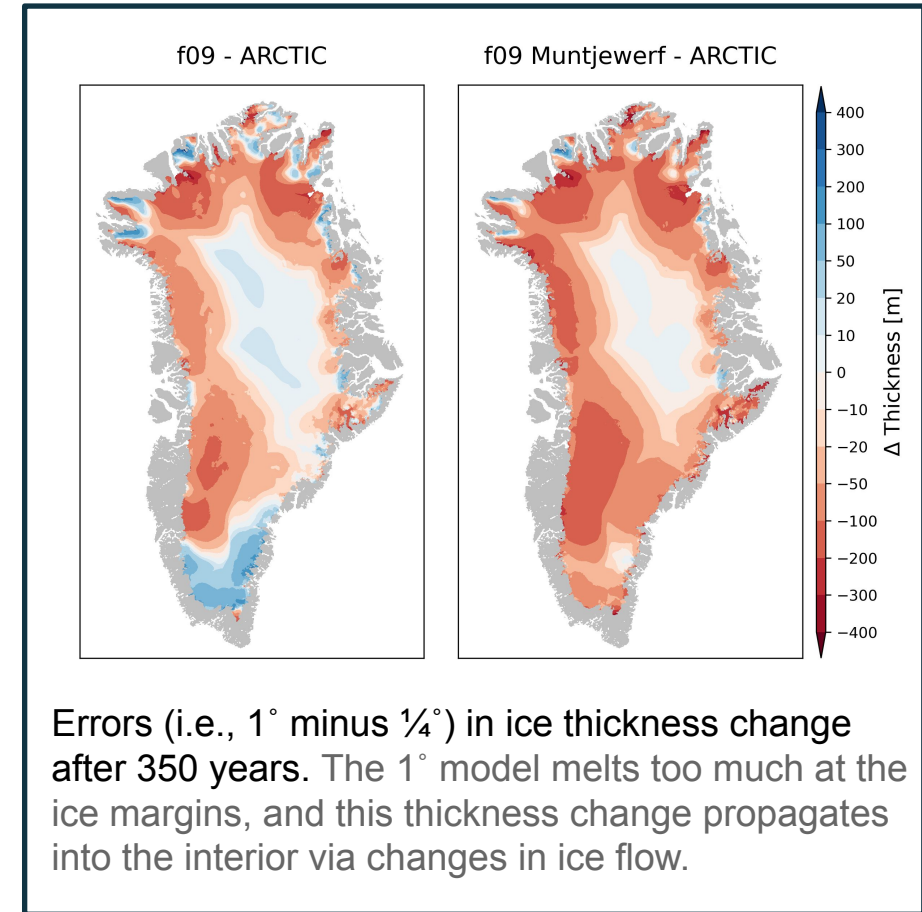
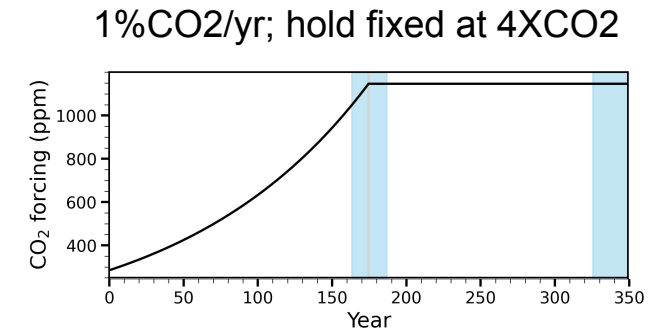
Here, we couple the **ARCTIC** grid to POP2/CISM2, re-tune the model and test:

Is the GrIS response different between 1° (f09) and 1/4° (ARCTIC)?

- **f09** – CESM2.1-CMIP6*
- **f09** – CESM2.1-no hacks**
- **ARCTIC** – CESM2.2



ARCTIC exhibits a smaller increase in melting, which primarily explains its lower mass loss and sea level rise.



Errors (i.e., 1° minus 1/4°) in ice thickness change after 350 years. The 1° model melts too much at the ice margins, and this thickness change propagates into the interior via changes in ice flow.

GrIS sea level contribution in the experiments