

The Sensitivity of the Global Mean Climate to Parameterized Momentum Flux in an Experimental Version of CAM6-CLUBB(X)

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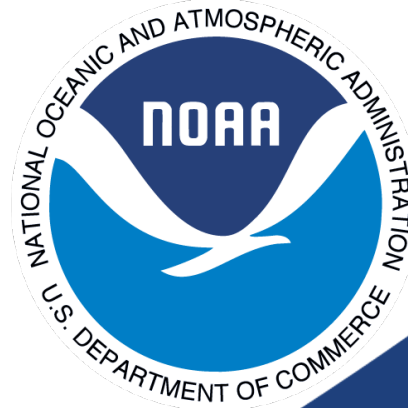
George H. Bryan

CESM Workshop

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PennState



CLUBBX's prognostic momentum flux formulation includes several tunable parameters

$$\frac{\overline{\partial u'w'}}{\partial t} = -\overline{w} \frac{\overline{\partial u'w'}}{\partial z} - \frac{1}{\rho} \frac{\partial \overline{\rho w'^2 u'}}{\partial z} - (1 - C_{uu_shr}) \overline{w'^2} \frac{\partial \overline{u}}{\partial z} - (1 - C_7) \overline{u'w'} \frac{\partial \overline{w}}{\partial z} + (1 - C_7) \frac{g}{\theta_{vs}} \overline{u'\theta'_v} - \frac{C_6}{\tau} \overline{u'w'} - \epsilon_{uw}$$



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**Advection of momentum flux
by the mean vertical wind**

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Turbulent advection of momentum flux



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Turbulent production
from updrafts and downdrafts

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Turbulent production
by existing momentum flux

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Buoyant production of momentum flux



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Return-to-isotropy adjustment term

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Term representing other dissipating processes

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- Allows for upgradient fluxes



CLUBB's prognostic momentum flux formulation includes several tunable parameters

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- Allows for upgradient fluxes
- C_6 , C_7 , and C_{uu_shr} are all tunable parameters in CLUBB



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- **The τ term in the return-to-isotropy adjustment can also be tuned using the new regime-specific formulation**

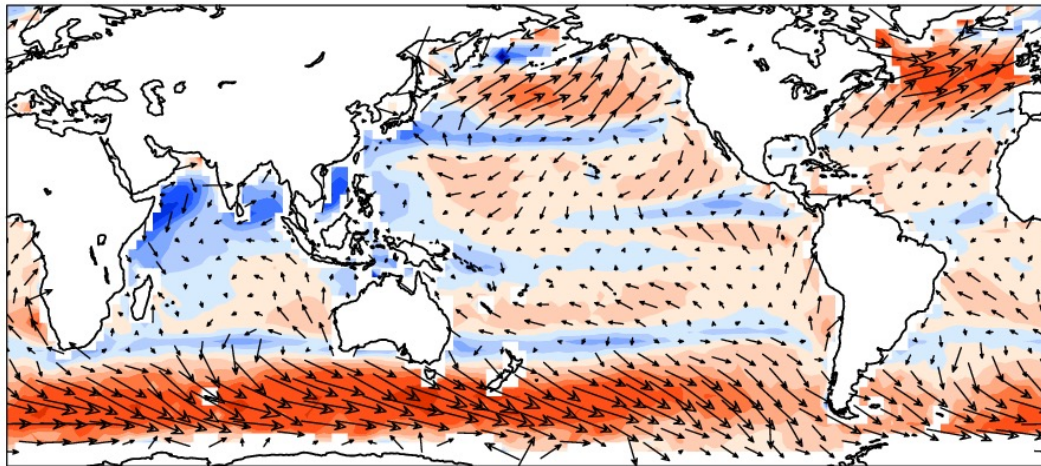


We target several notable mean state biases in CAM6-CLUBBX

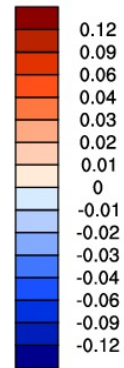


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Baseline Annual Surface Stress Bias



MIN = -0.20 MAX = 0.08



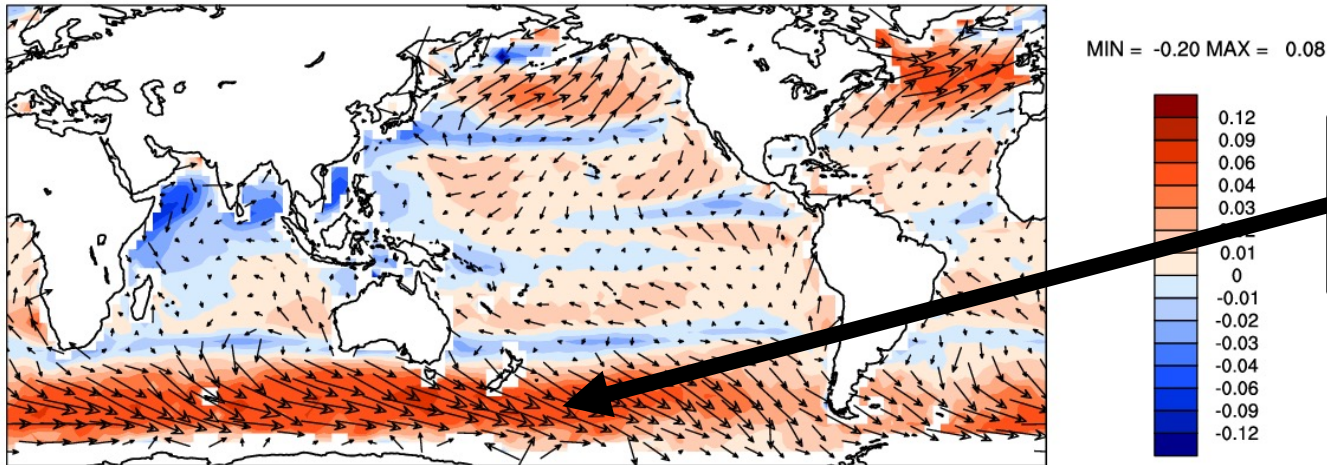
10-year simulations with
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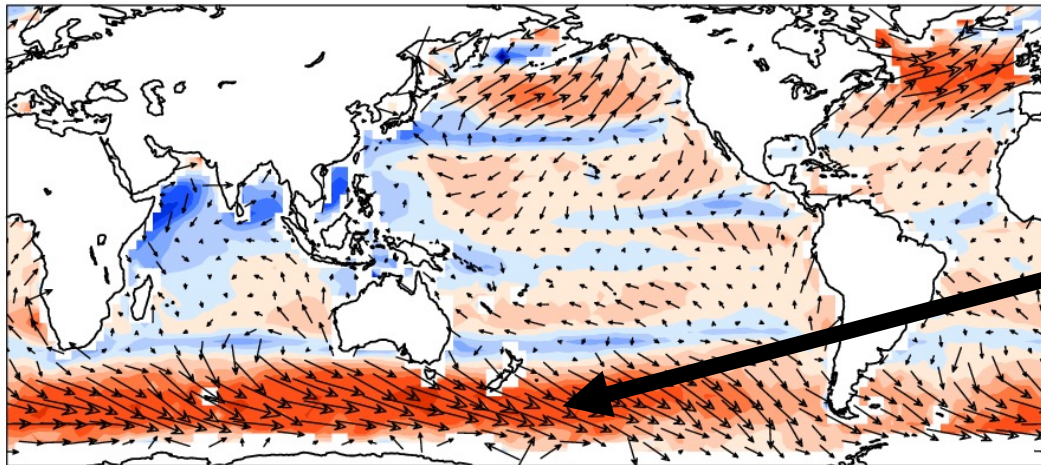
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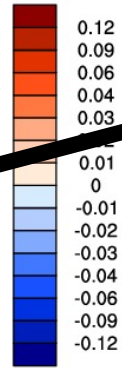
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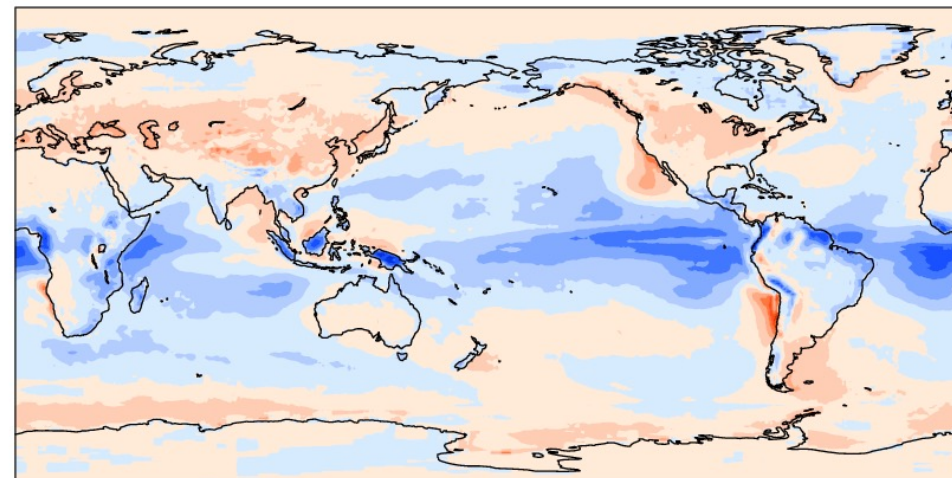
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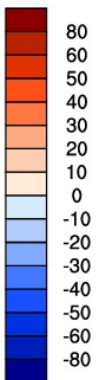
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Baseline Annual SWCF Bias

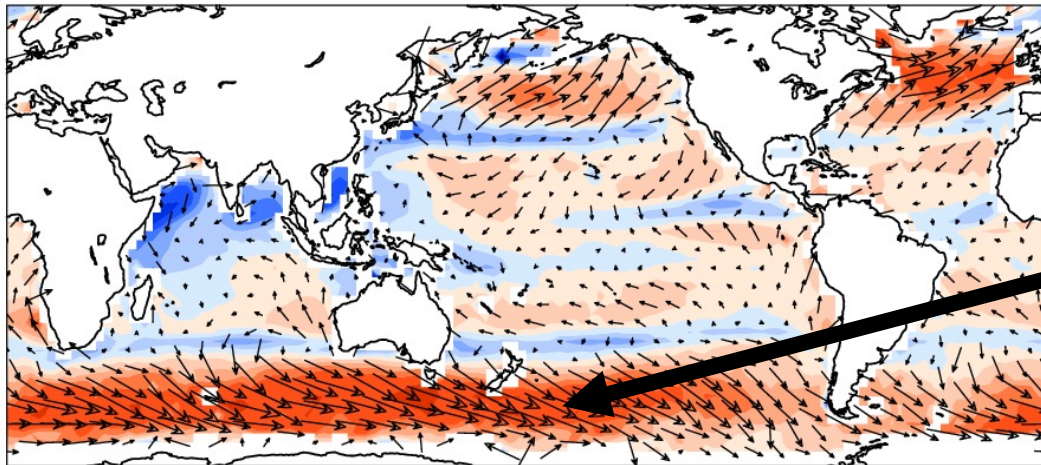


Min = -71.60 Max = 64.74

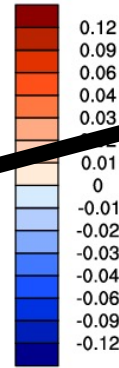


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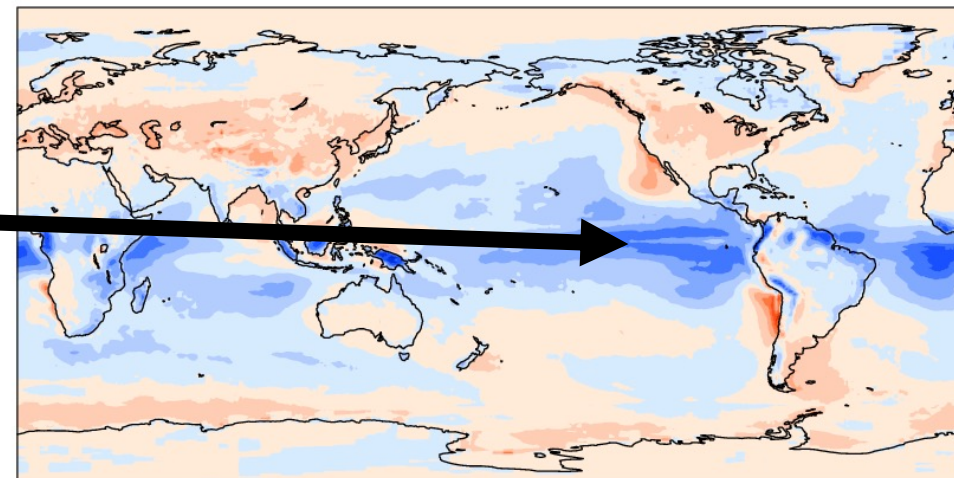


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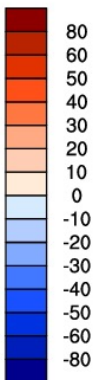
10-year simulations with baseline CLUBBX parameter settings

Negative baseline SWCF bias due to high bias in low cloud fraction

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- From one run to the next, change the value of only one input parameter
- Analyze the difference in model output between runs
- Repeat for **10** different initial combinations of input parameter values



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- Sensitivities calculated at t=72 hours (Qian et al. 2018, JGR-A)

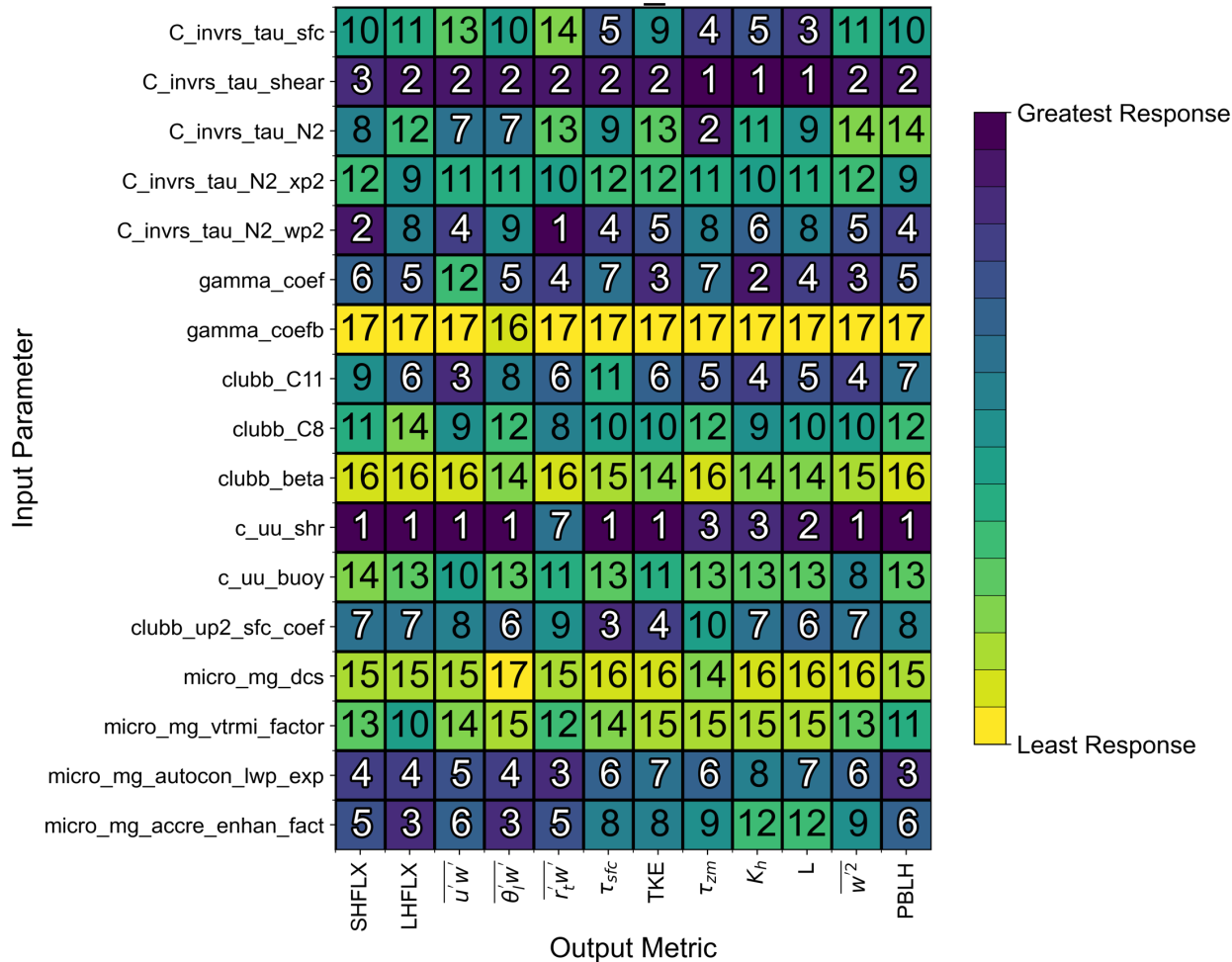


MOAT identifies several input parameters in the momentum flux budget that influence surface stress



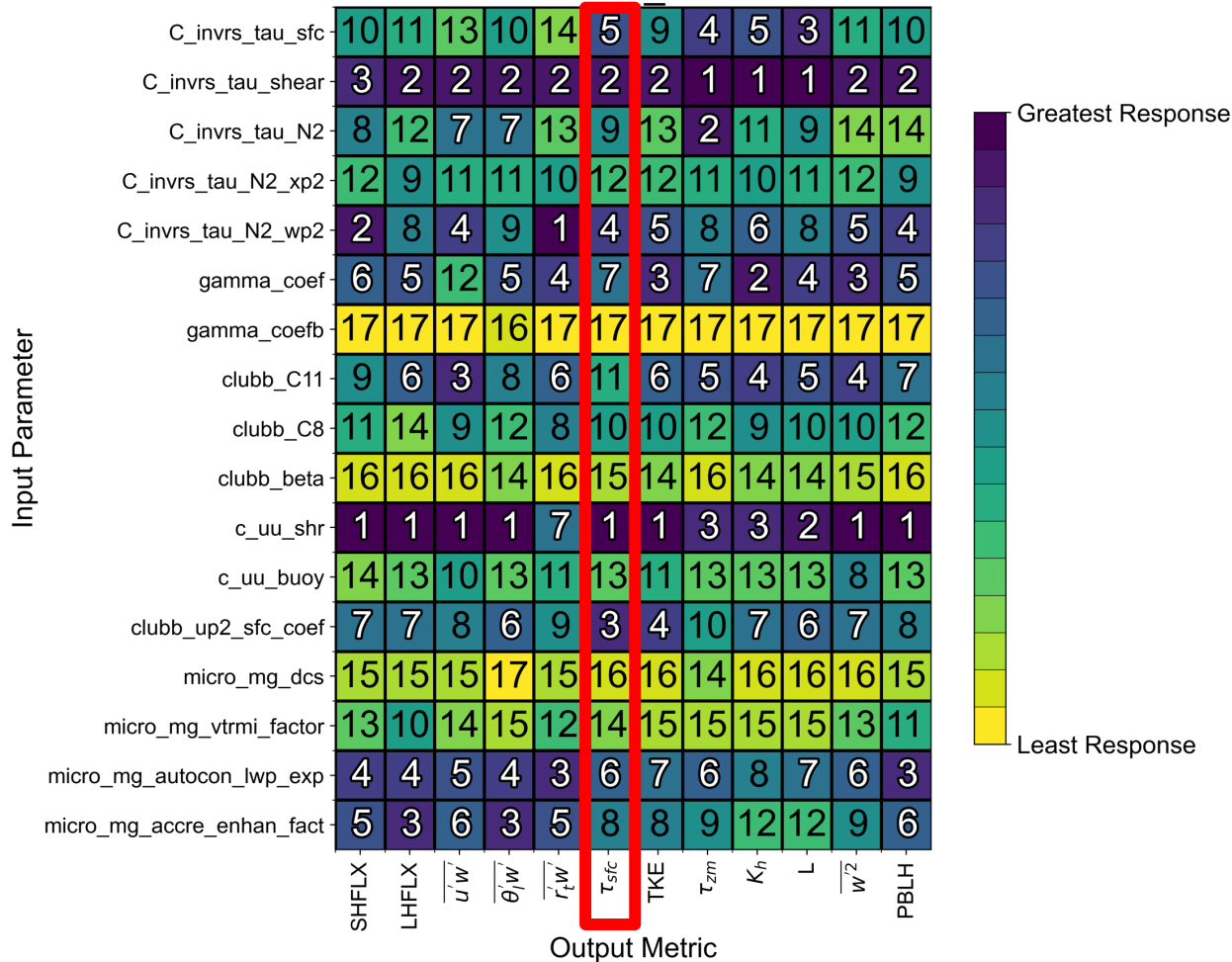
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μ^* Ranking mm3_mjst_20
southern_ocean



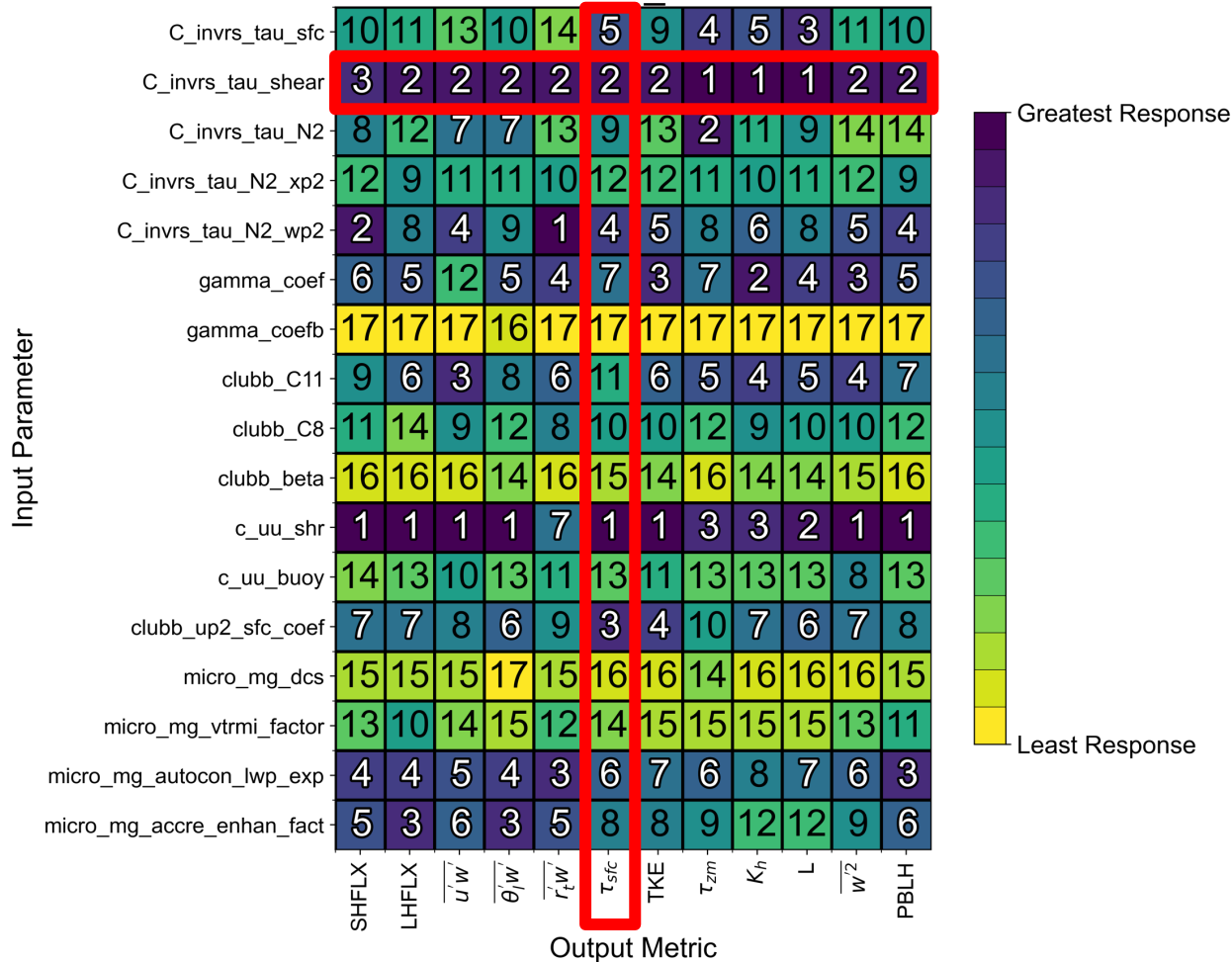
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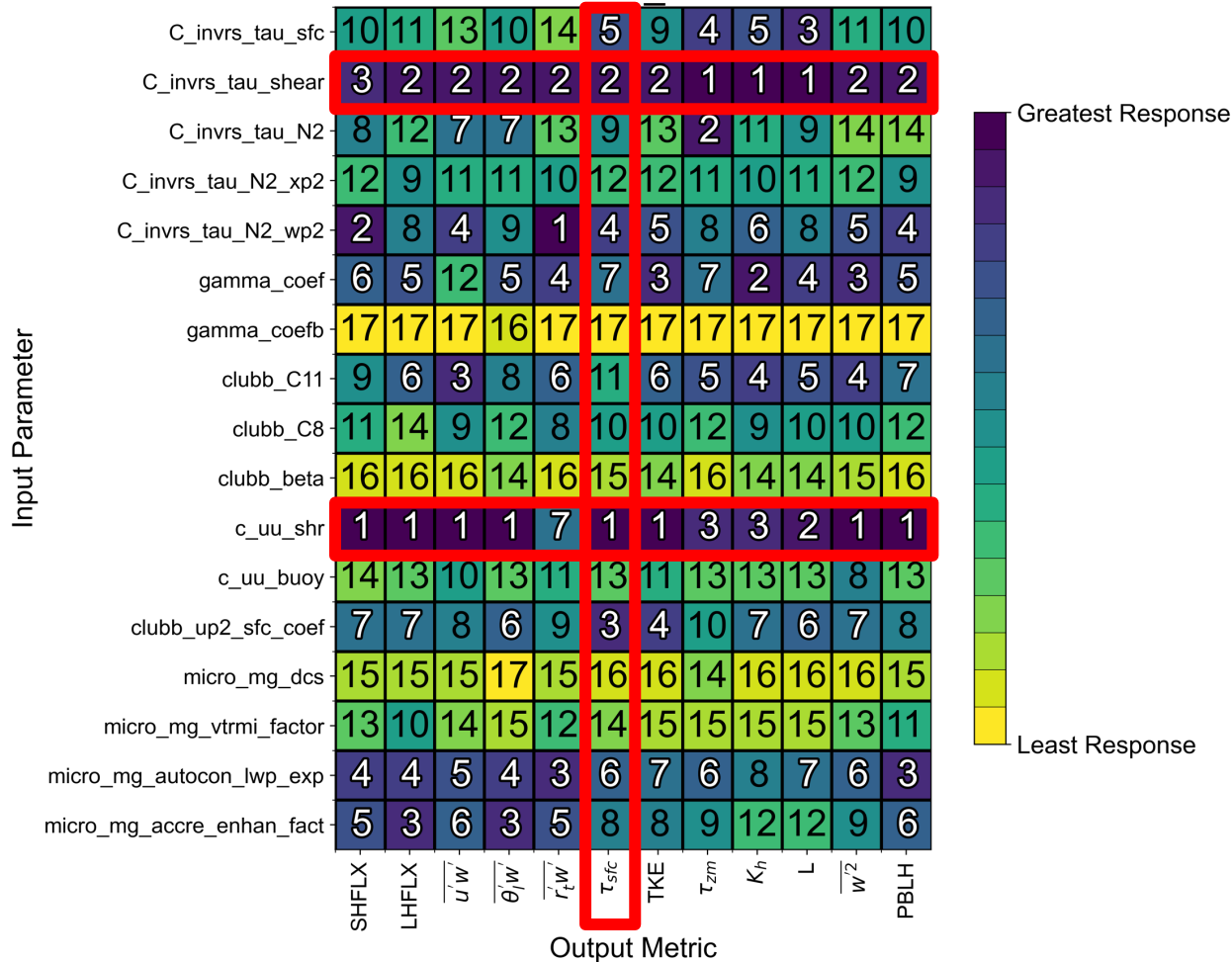
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$$\frac{C_6}{\tau} \overline{u'w'}$$

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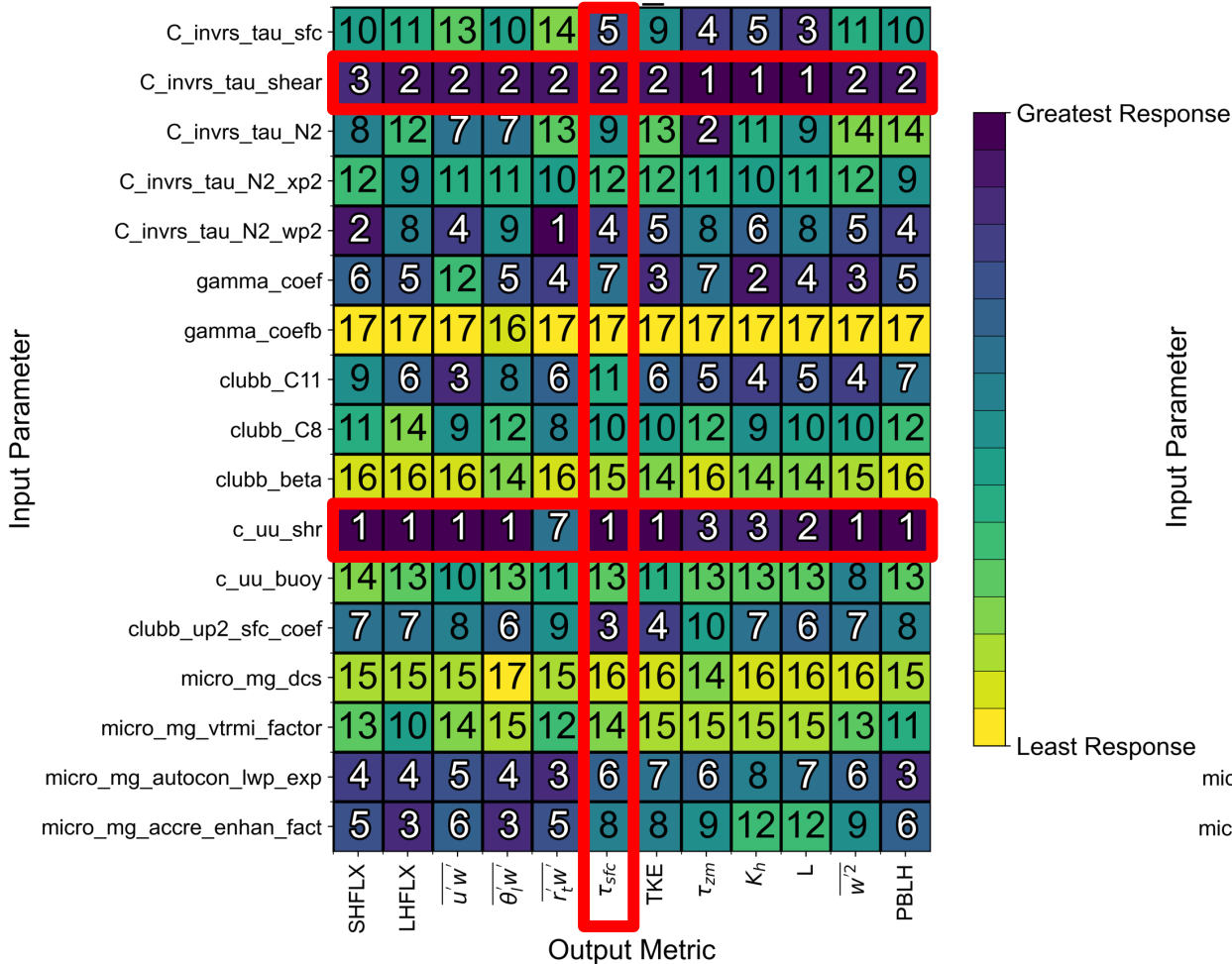
$$\frac{C_6}{\tau} \overline{u'w'}$$

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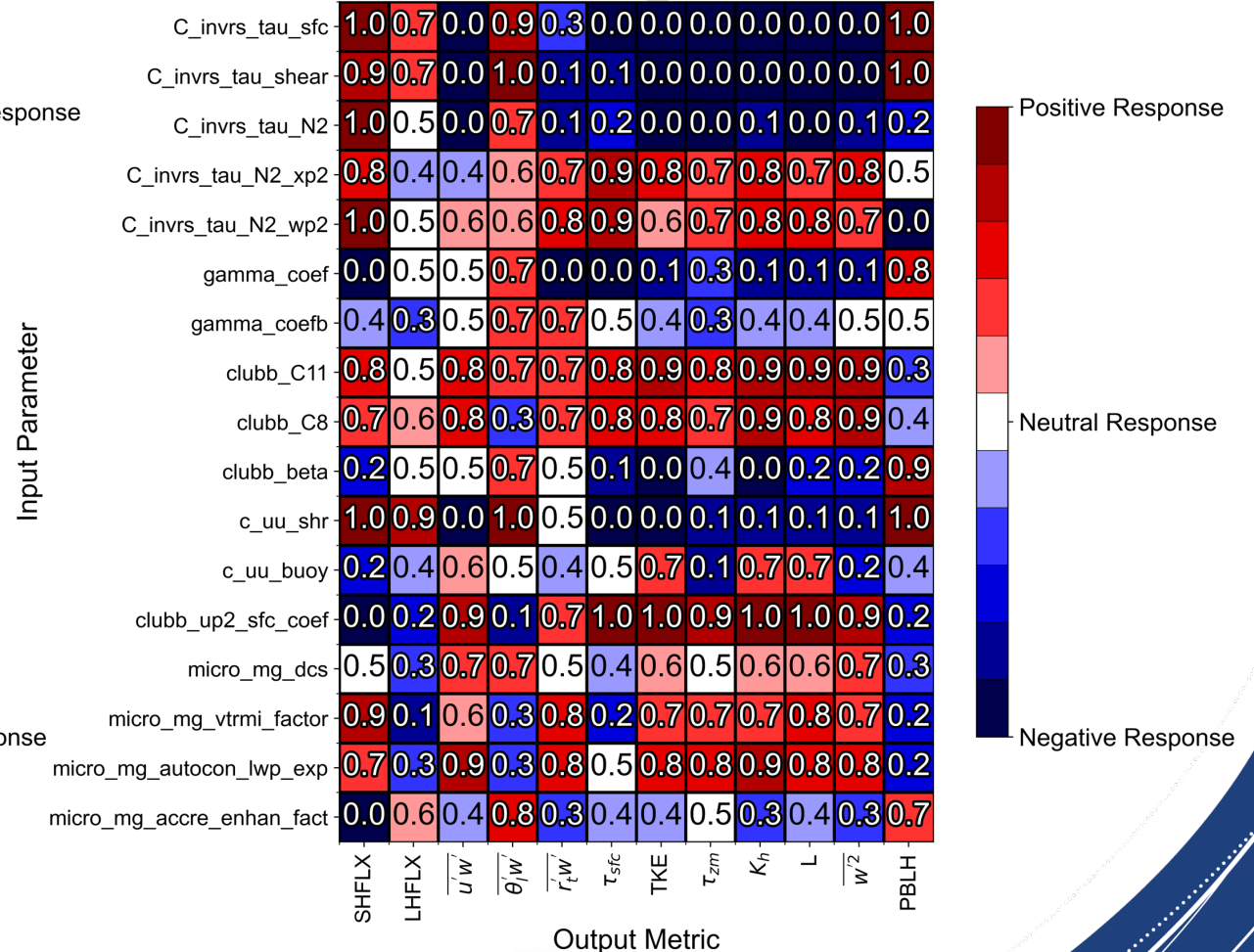


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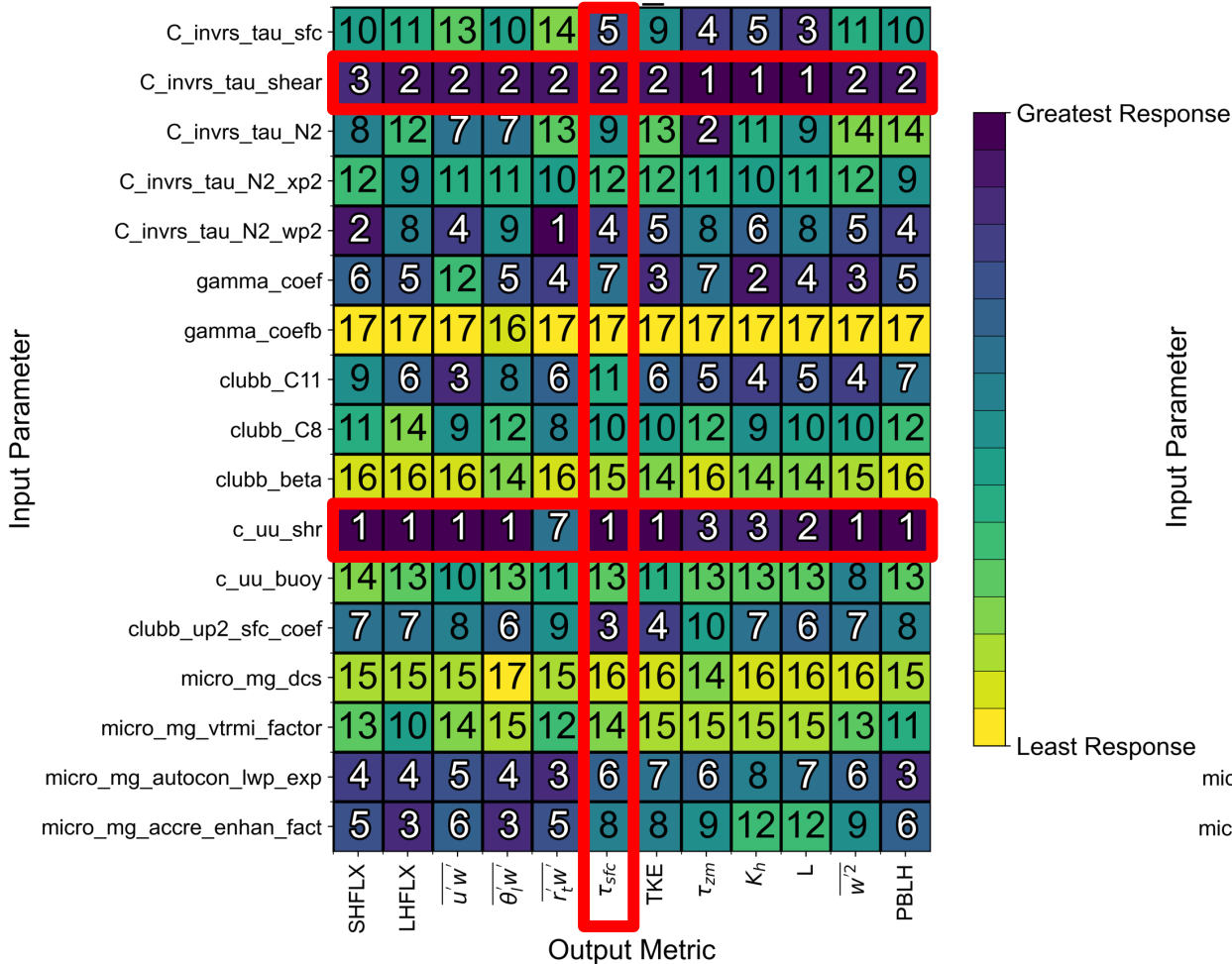


Monotonicity mm3_mjrd_20 southern_ocean

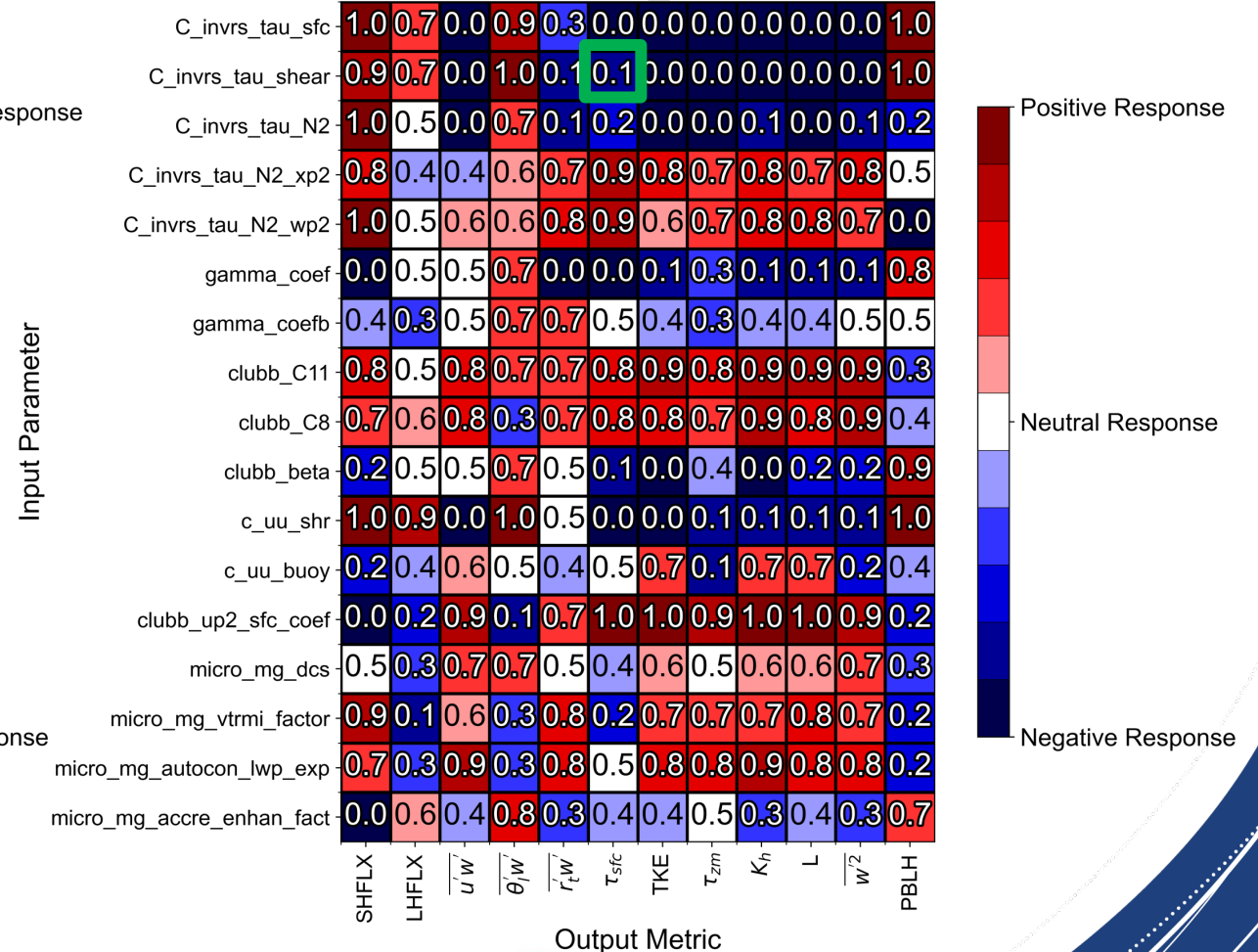


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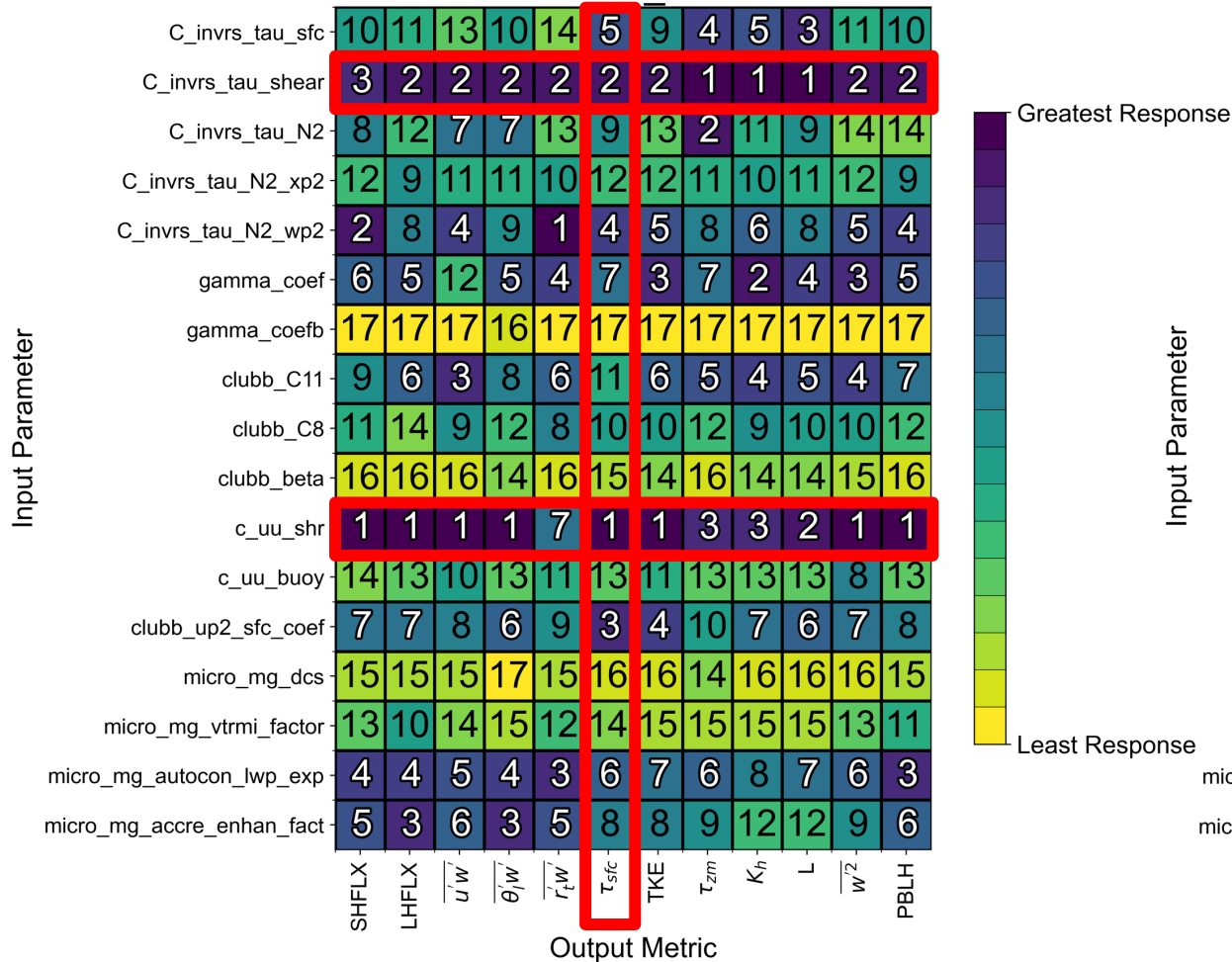


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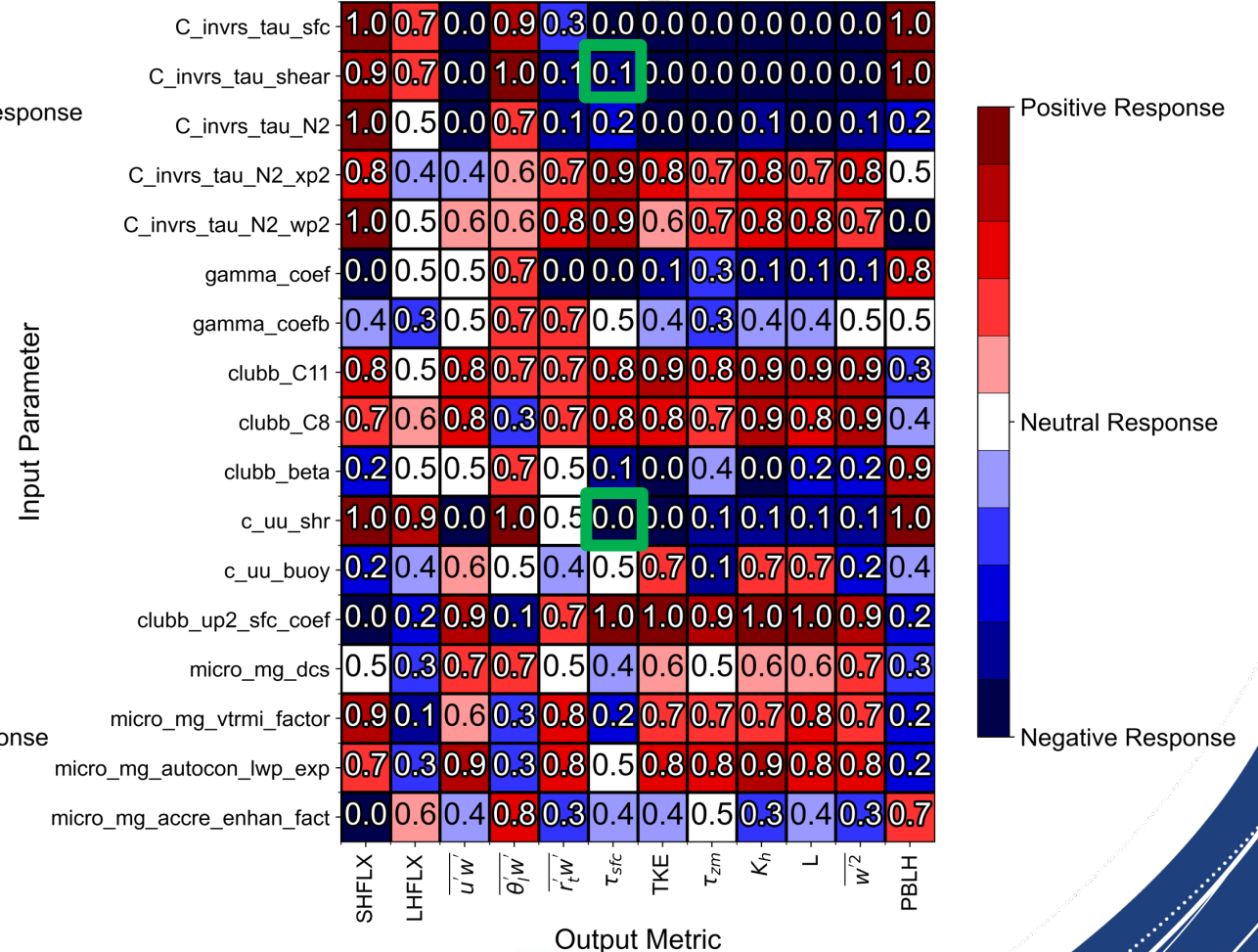


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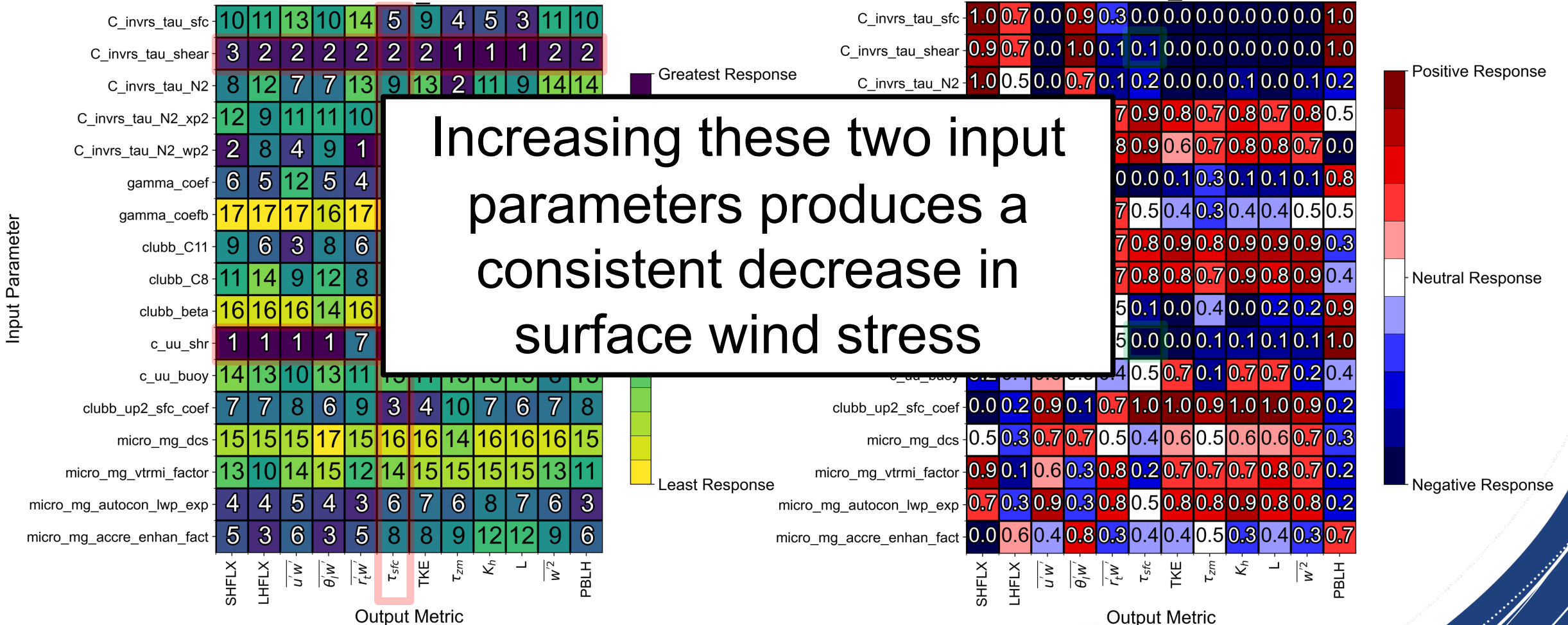
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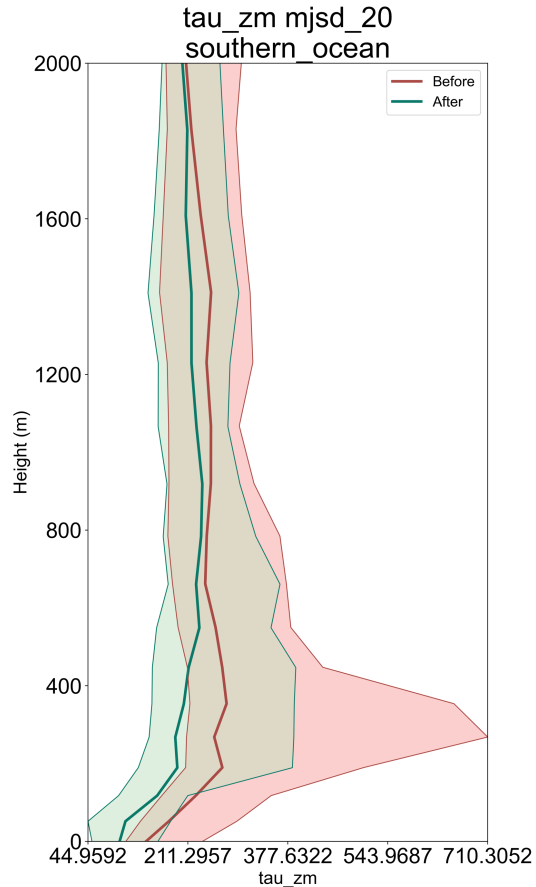
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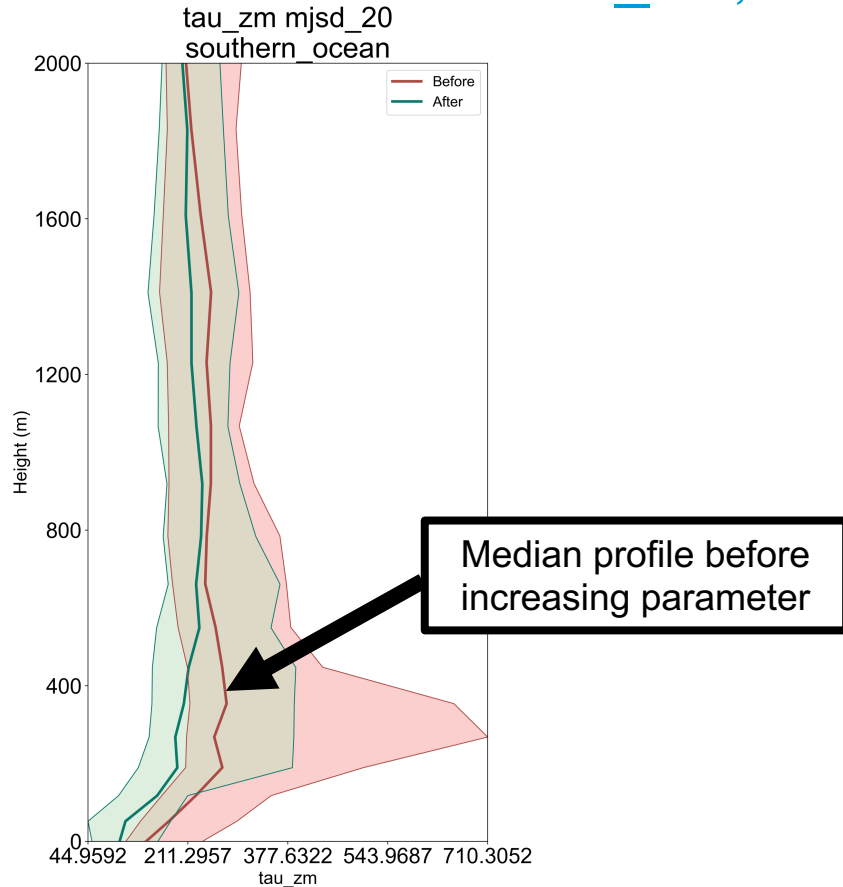


Increasing these two input parameters produces a consistent decrease in surface wind stress

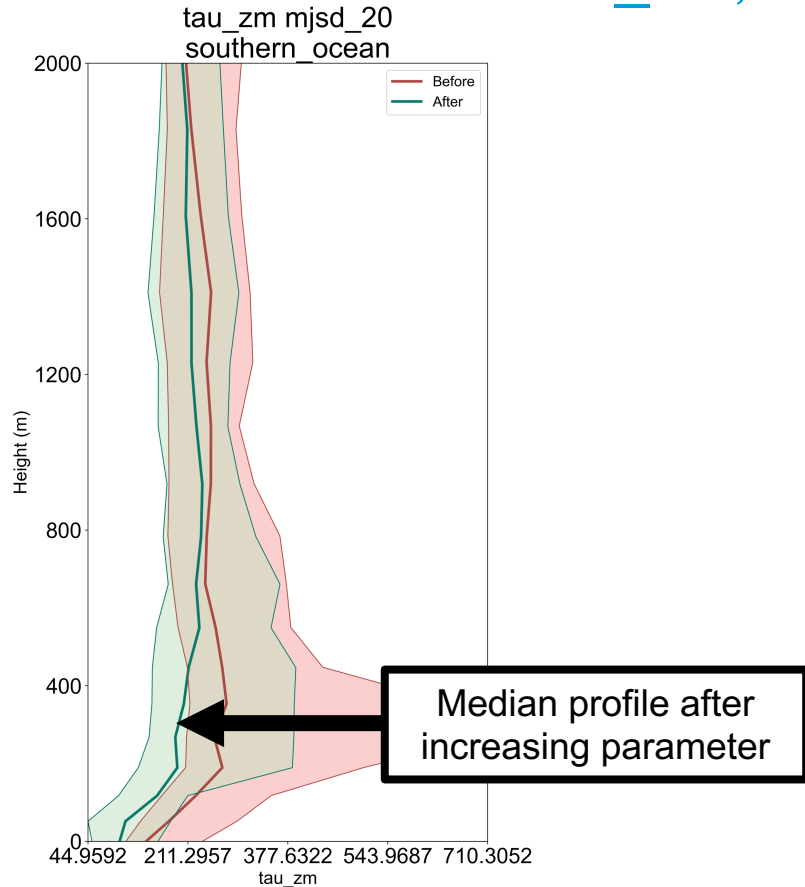
We can track the impact of increasing one of these terms ($C_{invrs_tau, shear}$) through the CLUBBX equations



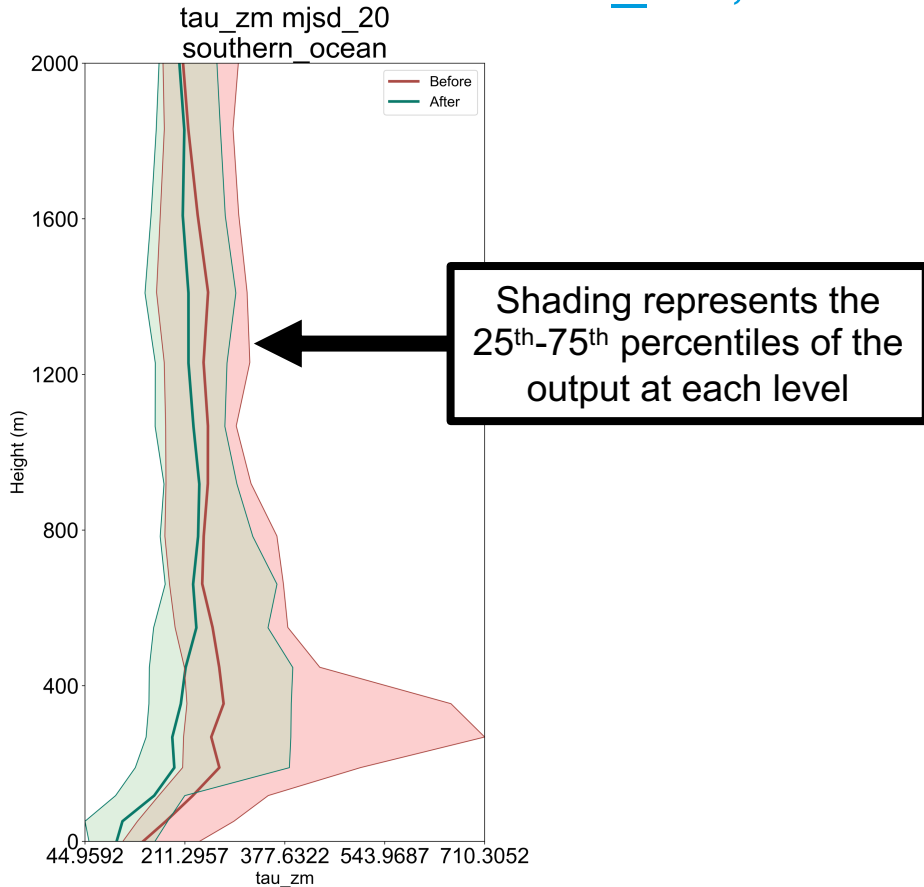
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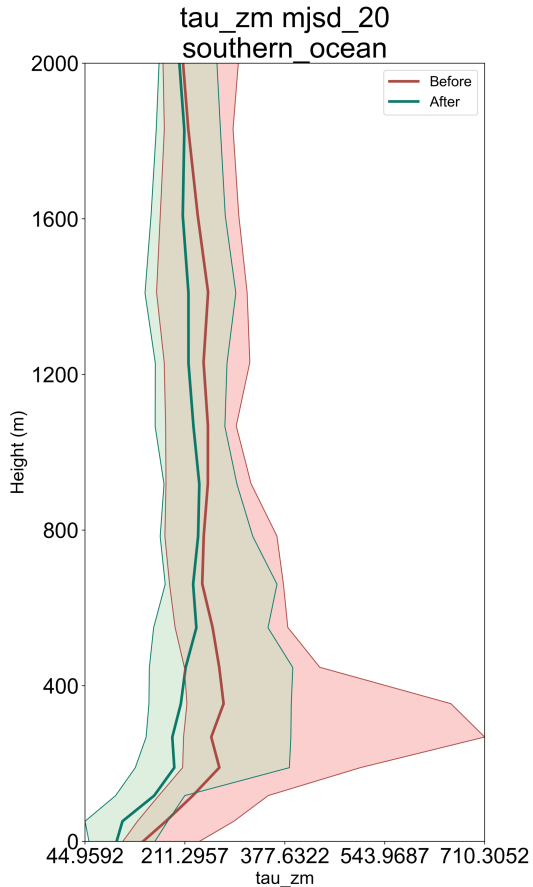
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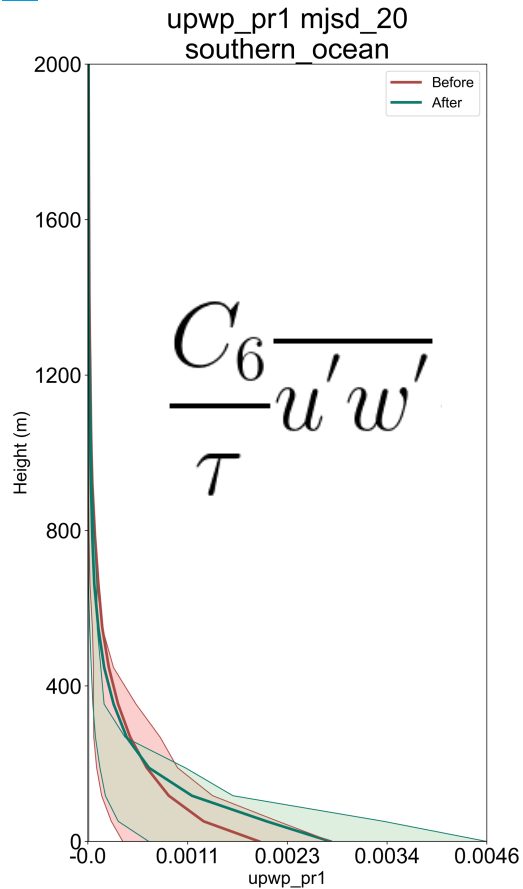
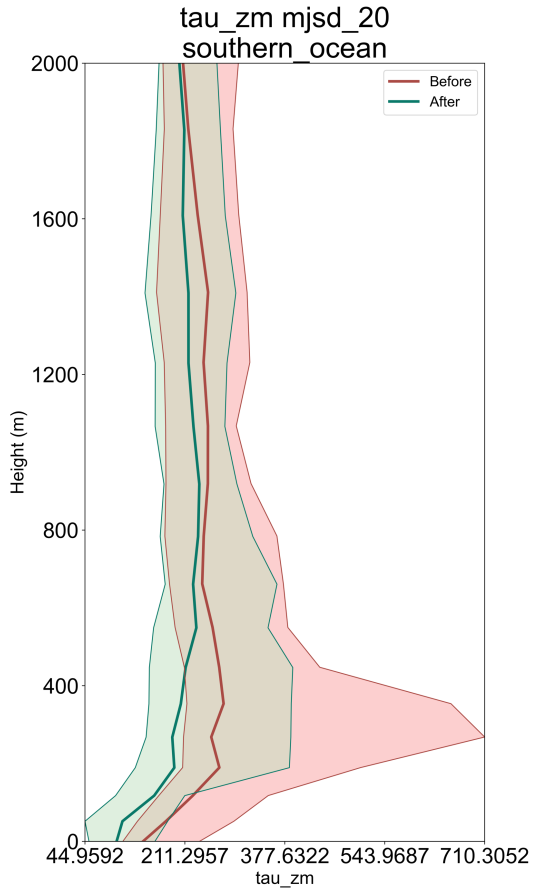
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Increasing parameter increases eddy dissipation and reduces the eddy turnover timescale/turbulent length scale



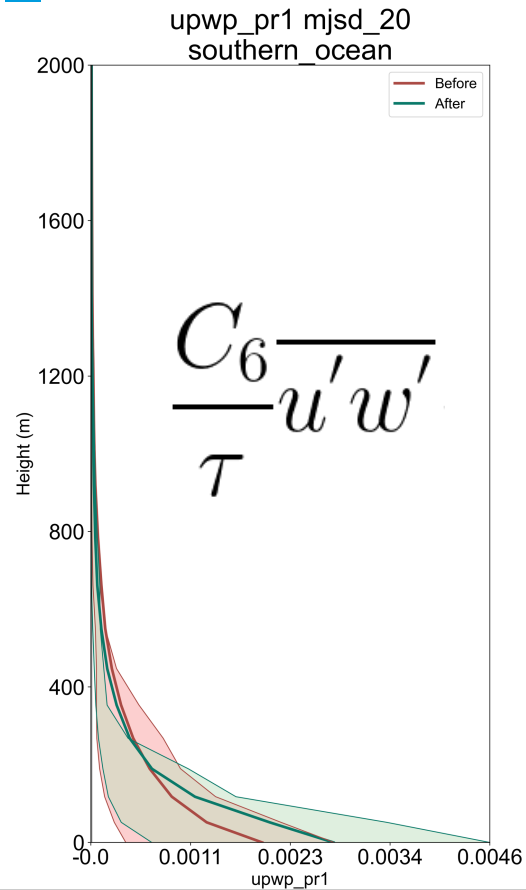
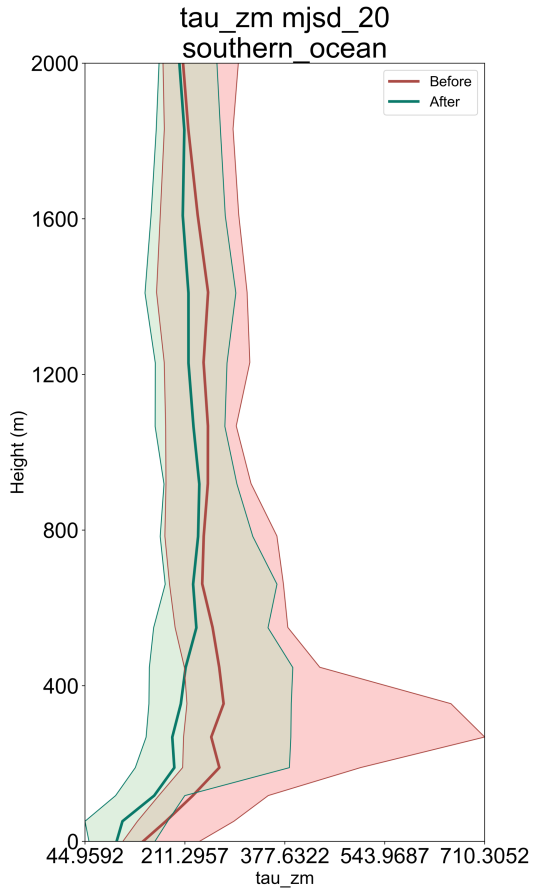
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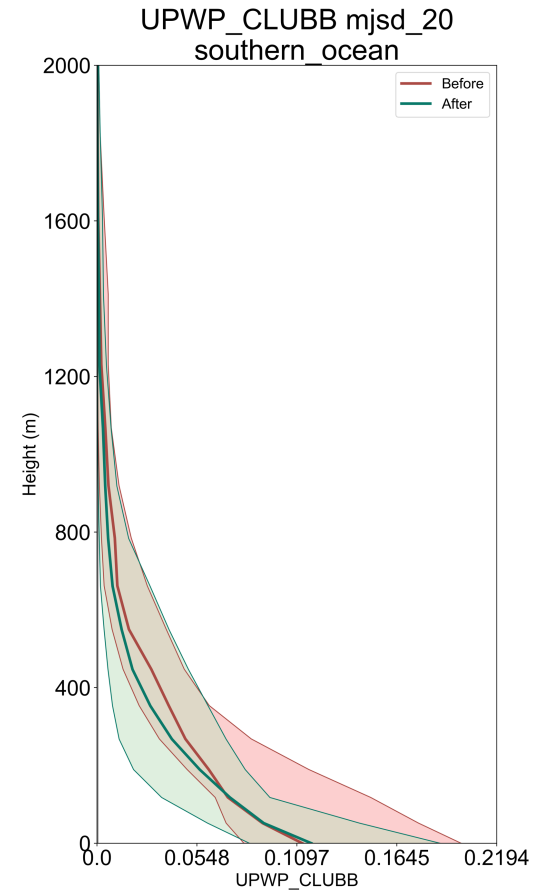
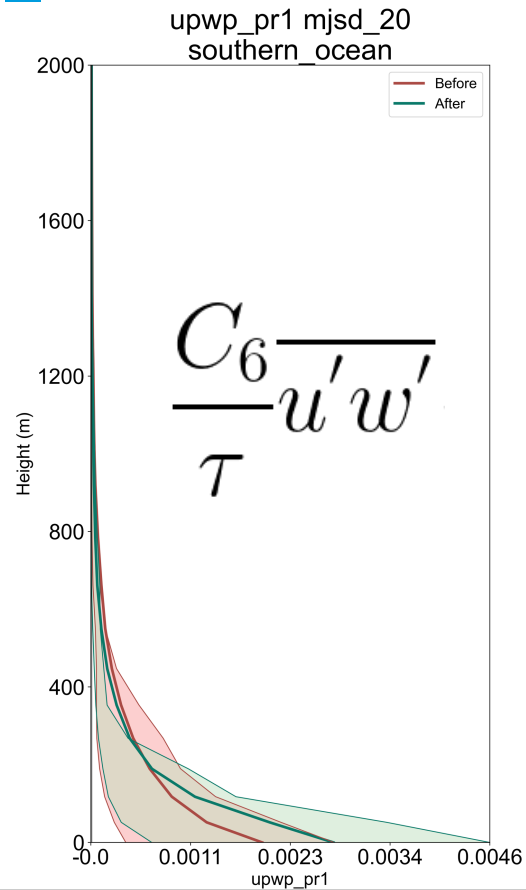
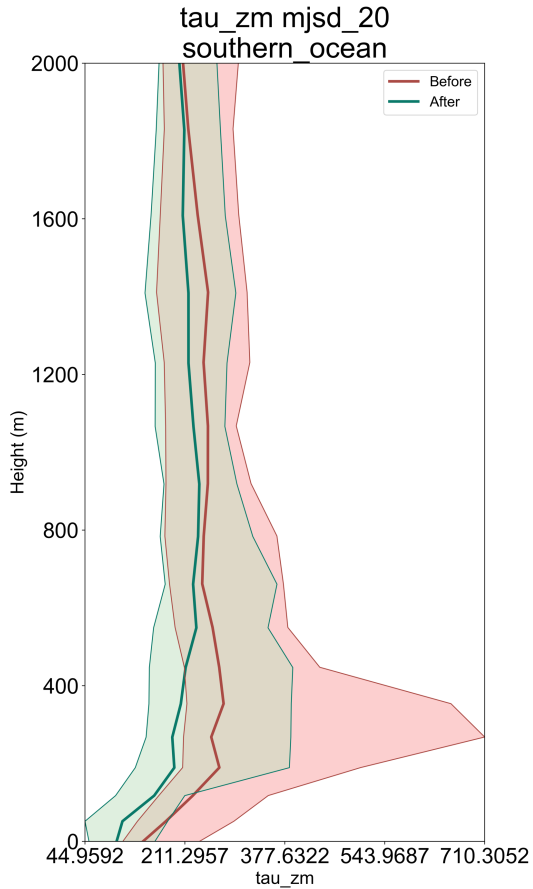


Increasing parameter increases eddy dissipation and reduces the eddy turnover timescale/turbulent length scale

Increasing parameter increases the inverse of the eddy timescale, thus increasing the return-to-isotropy term



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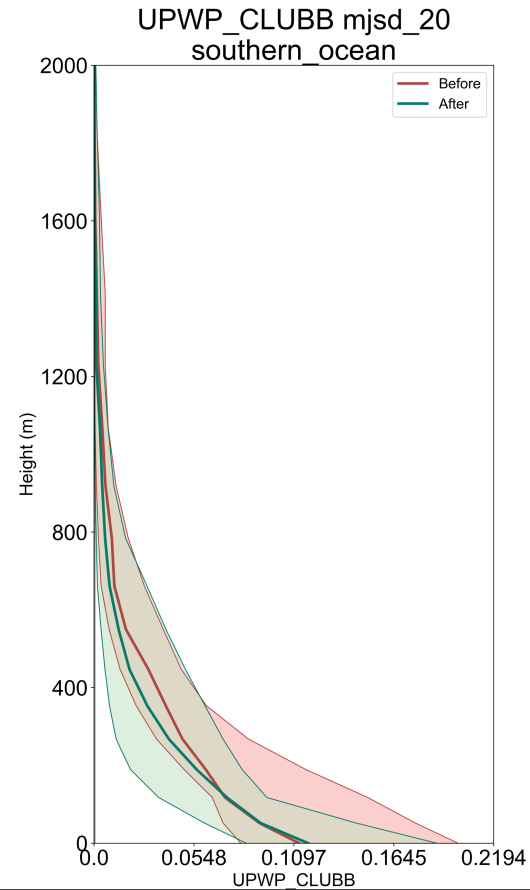
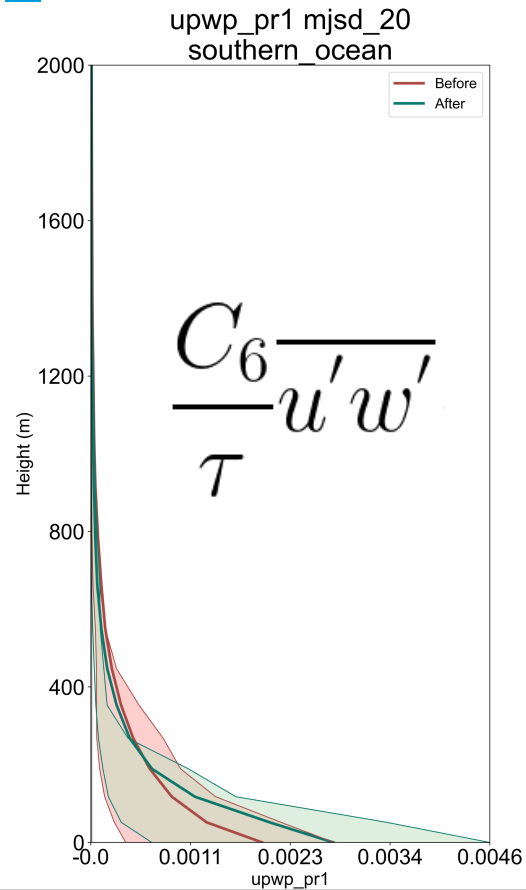
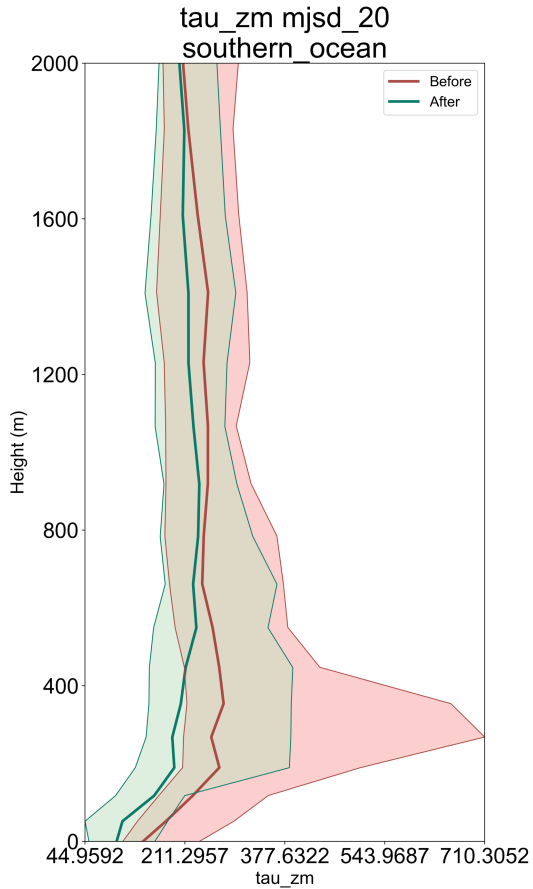


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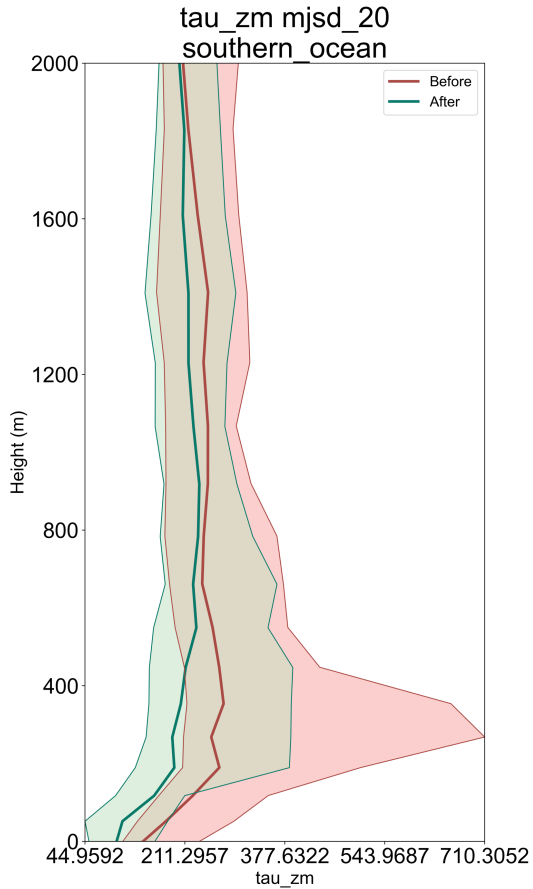
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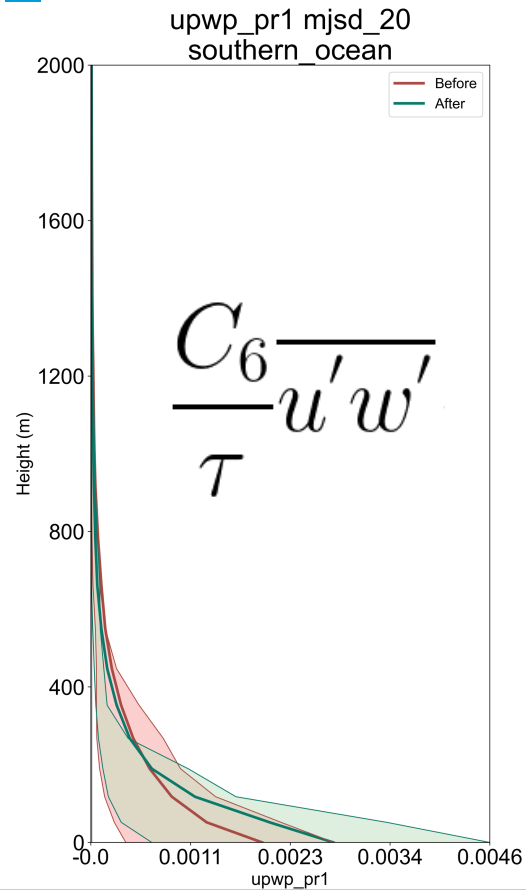
Increasing the magnitude of the return-to-isotropy term reduces the time tendency of vertical momentum flux



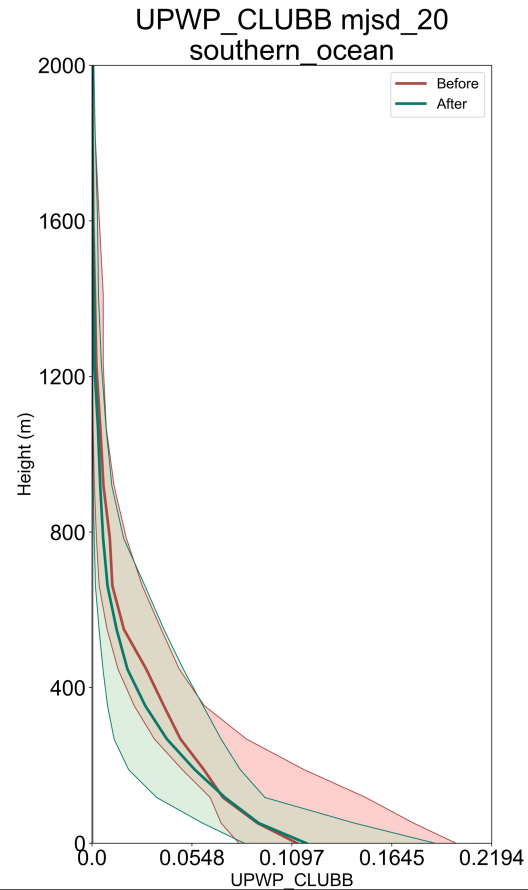
We can track the impact of increasing one of these terms ($C_{invs_tau, shear}$) through the CLUBB equations



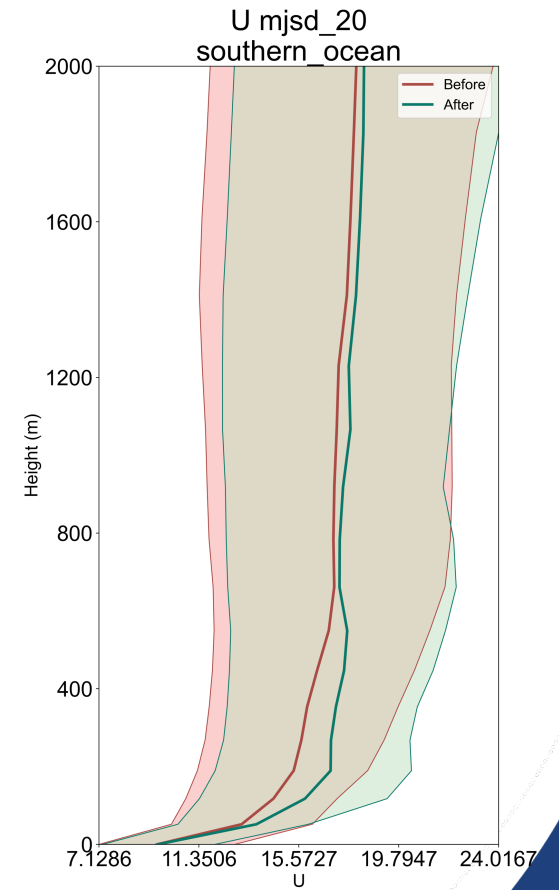
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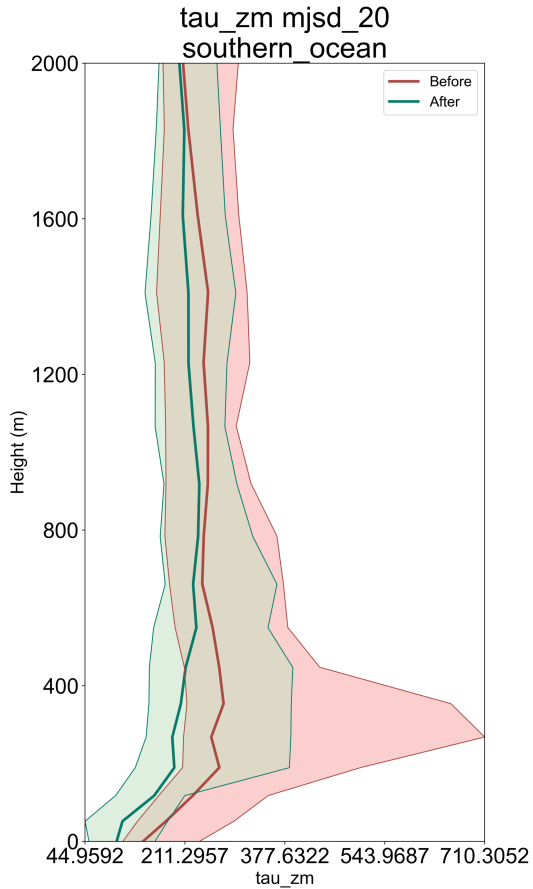
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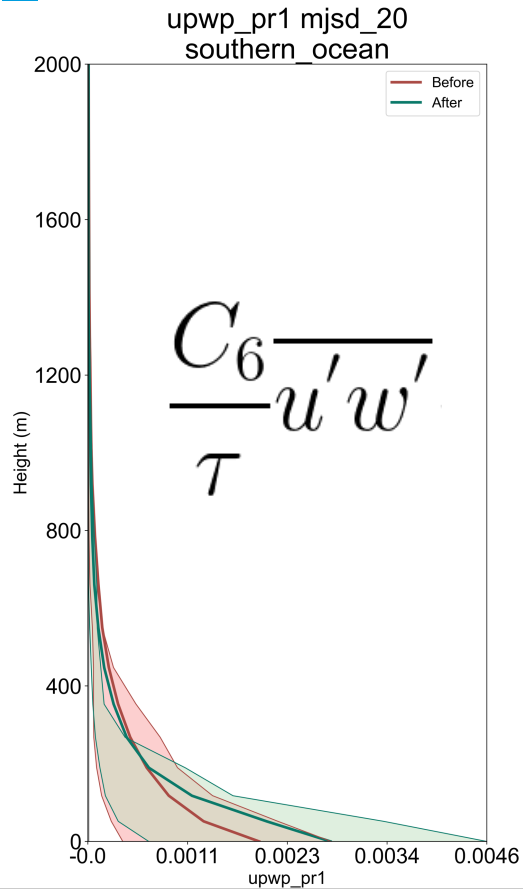
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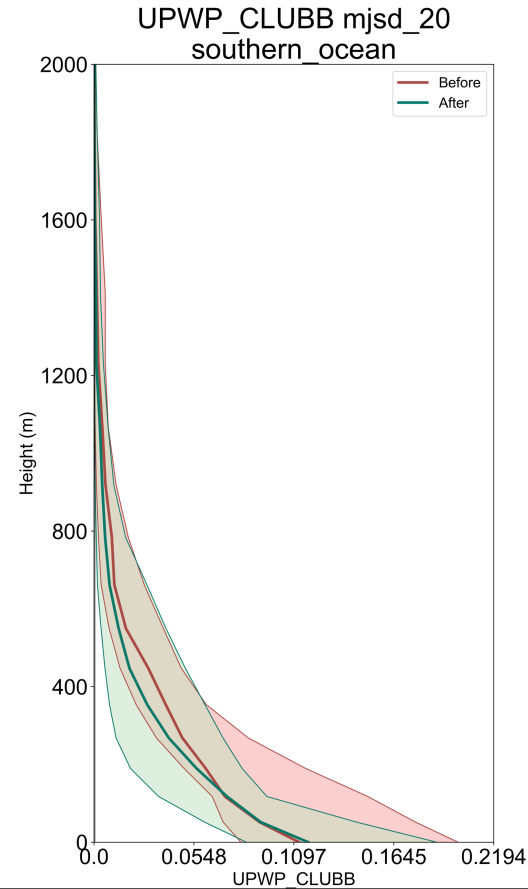
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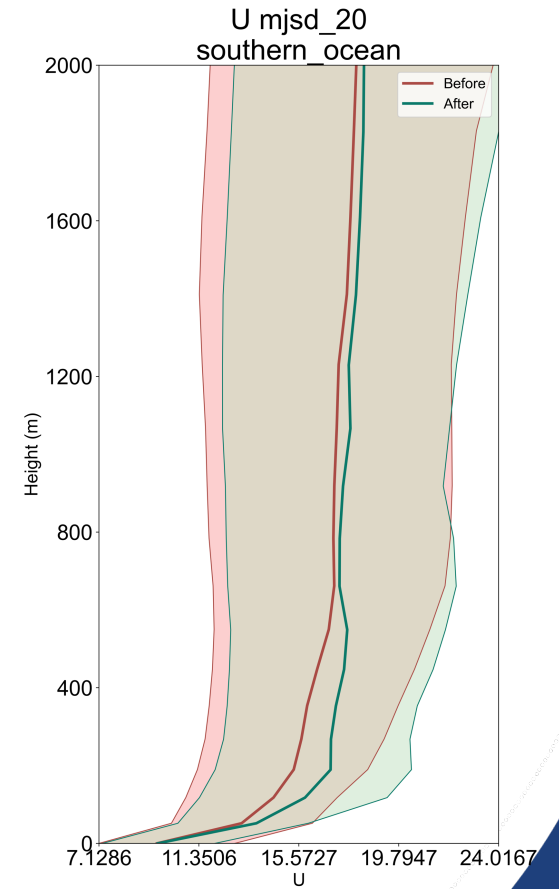
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Increasing the magnitude of the return-to-isotropy term reduces the time tendency of vertical momentum flux



Reducing the magnitude of vertical momentum flux reduces the mixing of stronger winds down to the surface

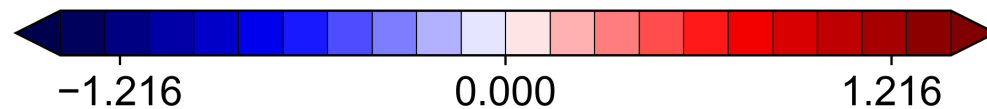
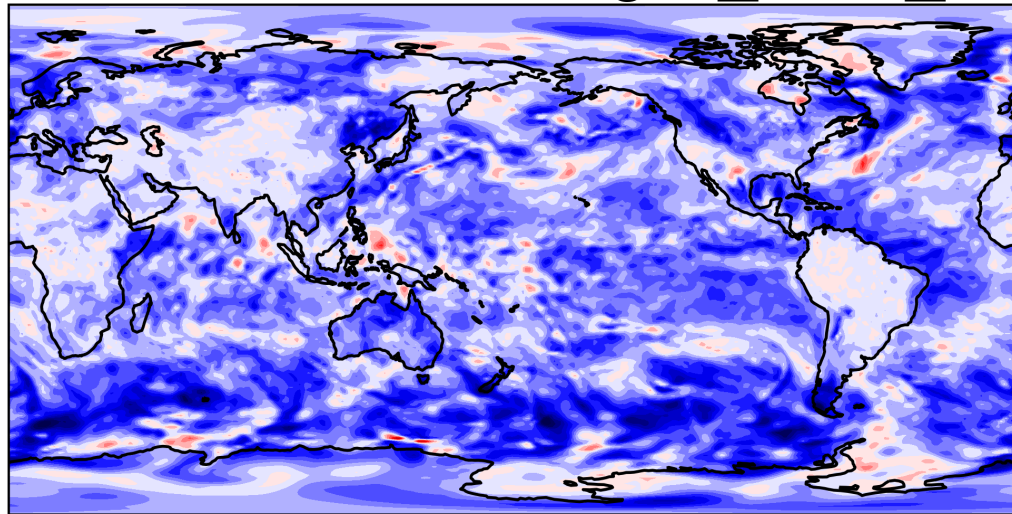


With reduced near-surface wind speeds, surface wind stress is reduced



With reduced near-surface wind speeds, surface wind stress is reduced

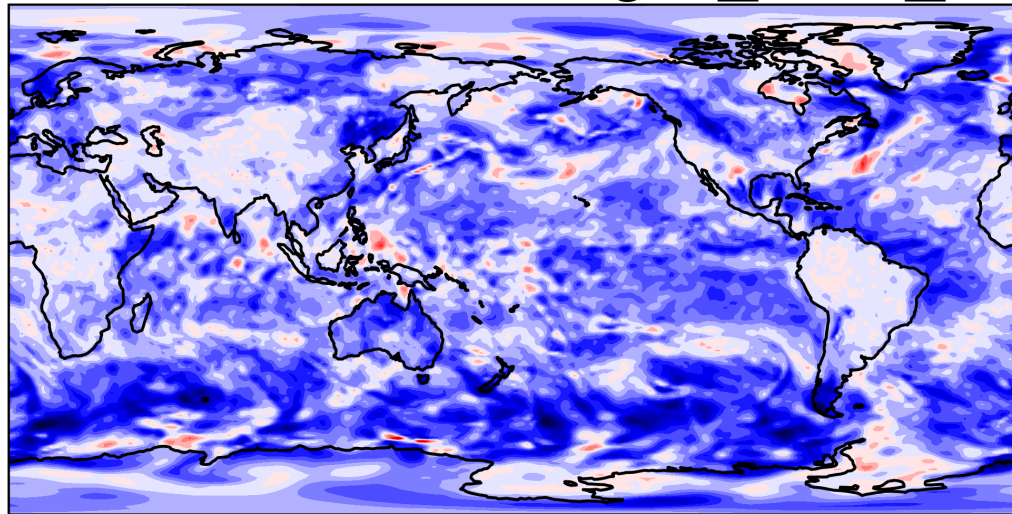
Difference in UBOT
Before and After Increasing $C_{invrs_tau_shear}$



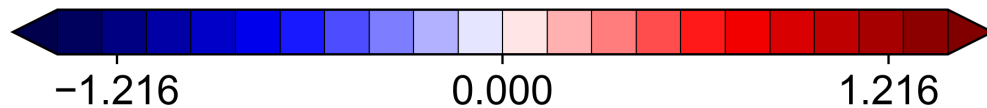
After Avg. = 5.68563 Before Avg. = 5.99972 Diff = -0.31409

With reduced near-surface wind speeds, surface wind stress is reduced

Difference in UBOT Before and After Increasing $C_{invrs_tau_shear}$



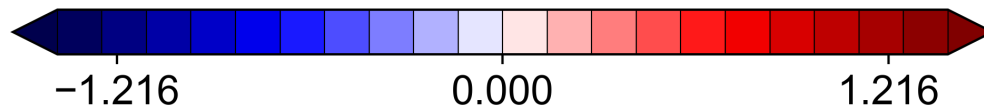
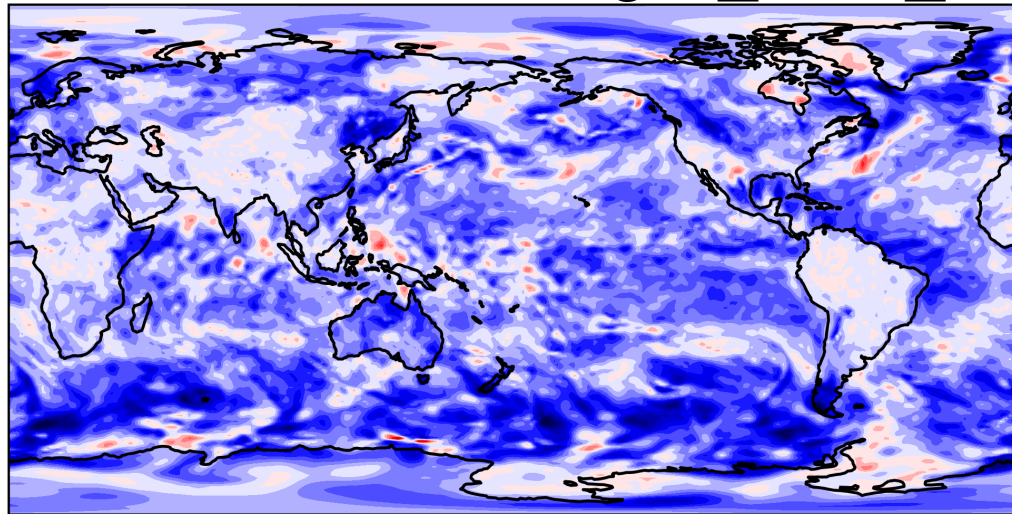
Difference in globally-averaged near-surface wind between "before" and "after" configs



After Avg. = 5.68563 Before Avg. = 5.99972 **Diff = -0.31409**

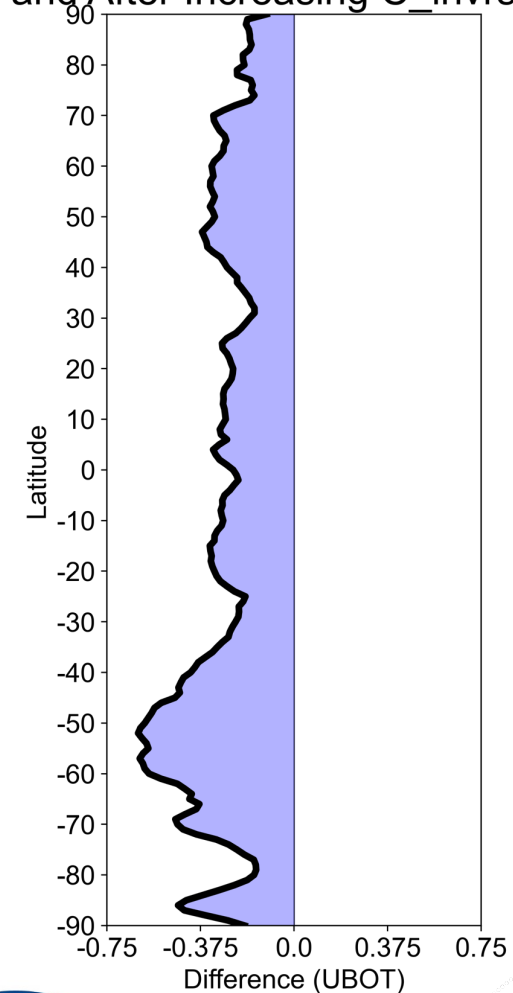
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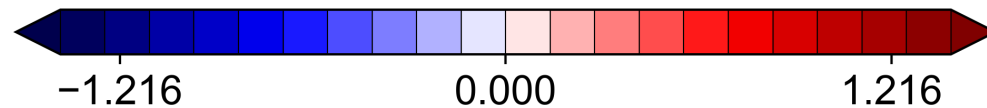
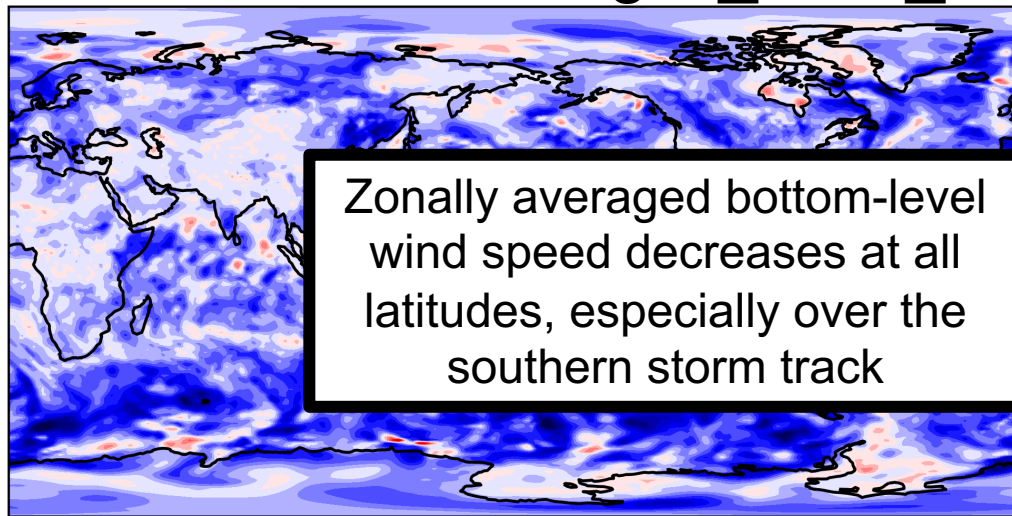
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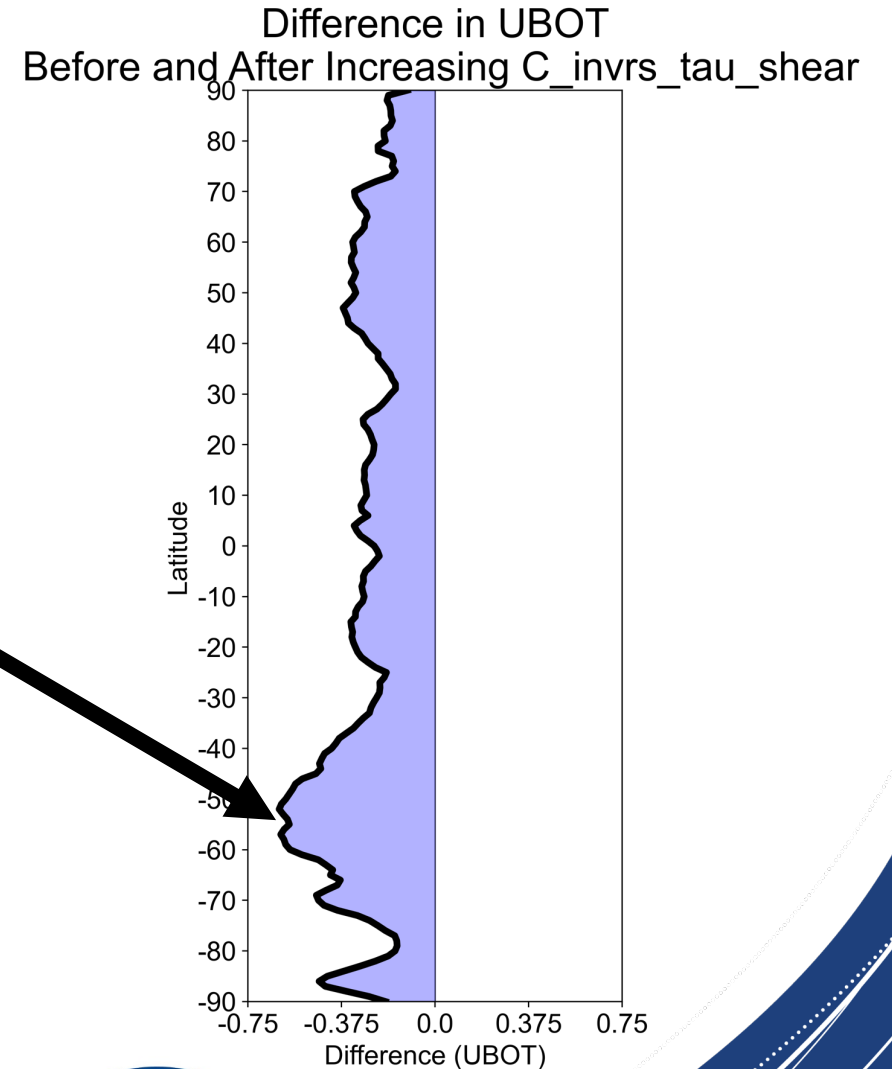


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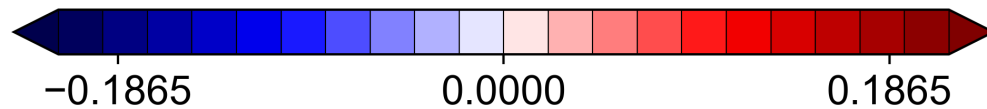
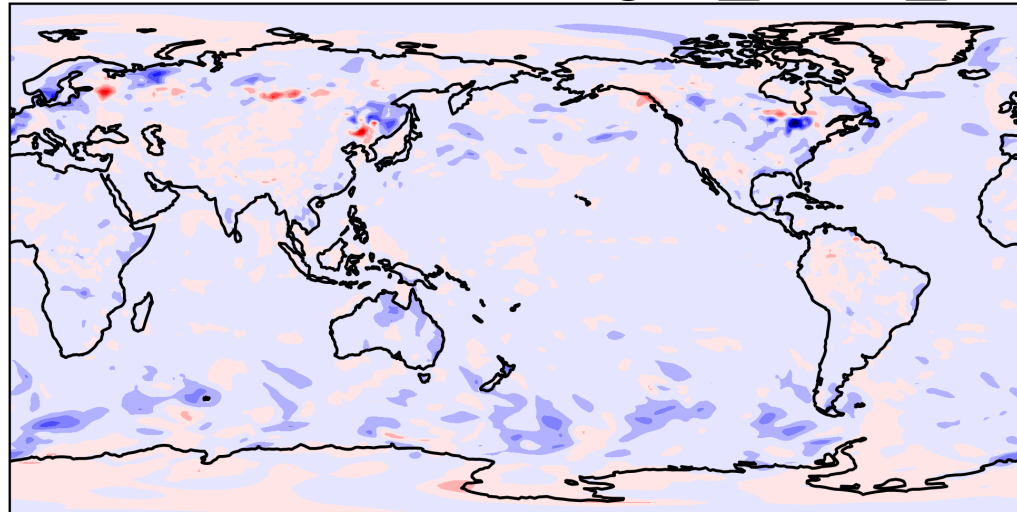
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Difference in TAU

Before and After Increasing $C_{invs_tau_shear}$

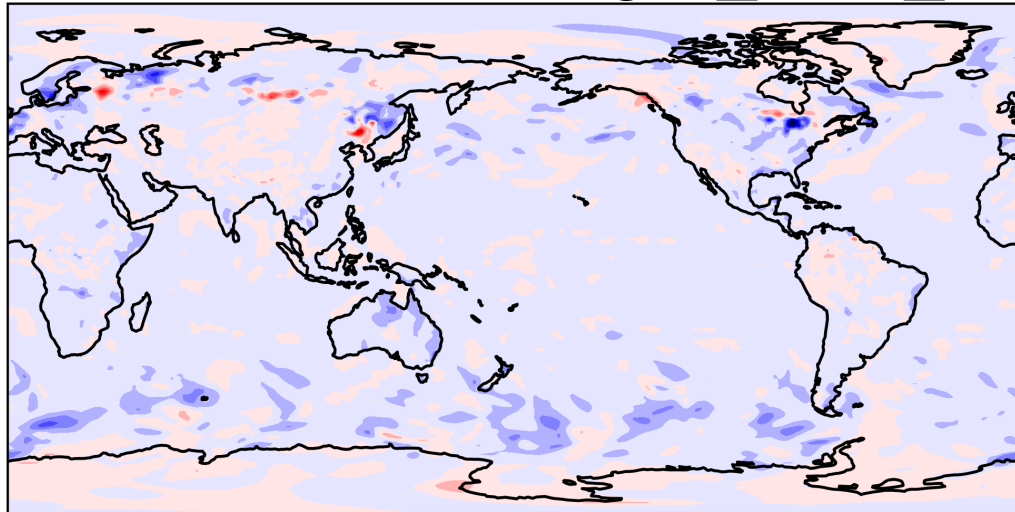


After Avg. = 0.10154 Before Avg. = 0.10774 Diff = -0.0062

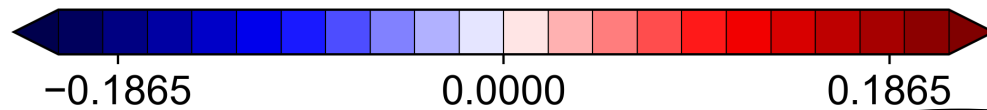


With reduced near-surface wind speeds, surface wind stress is reduced

Difference in TAU Before and After Increasing C_invrs_tau_shear



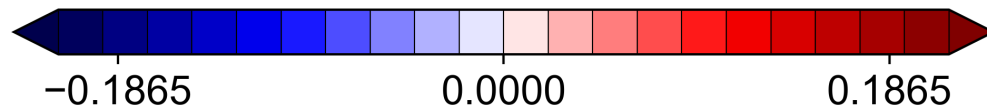
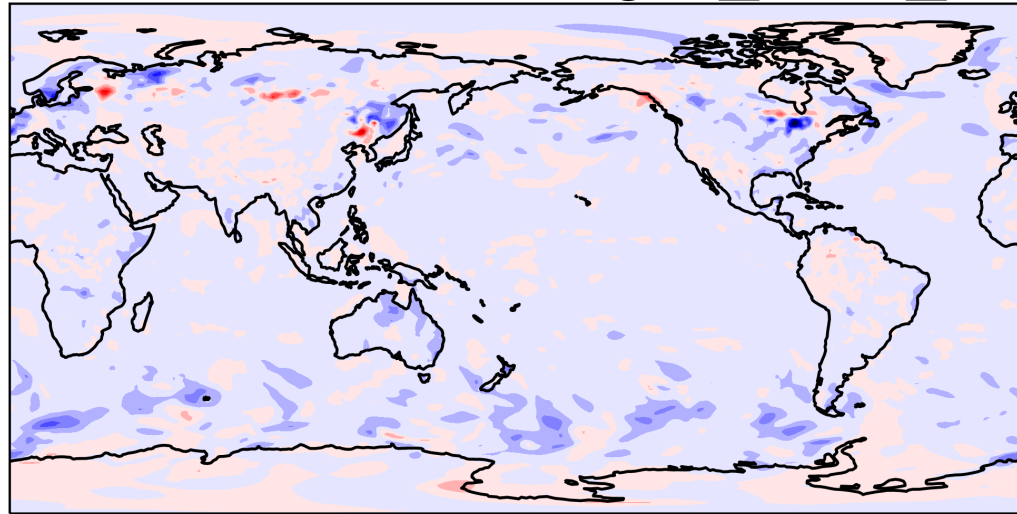
Difference in globally-averaged surface stress between "before" and "after" configs



After Avg. = 0.10154 Before Avg. = 0.10774 **Diff = -0.0062**

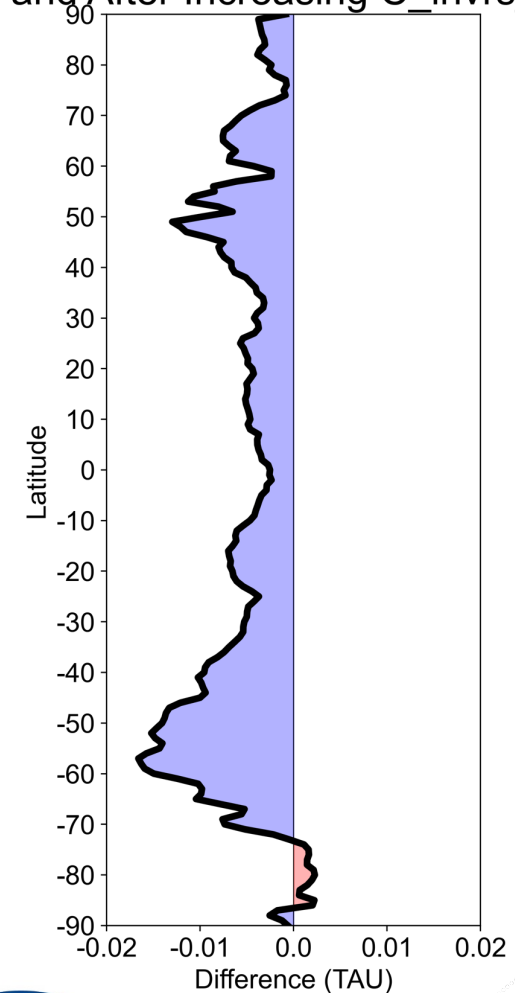
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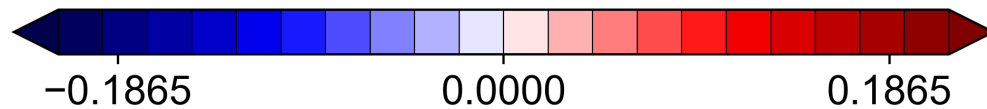
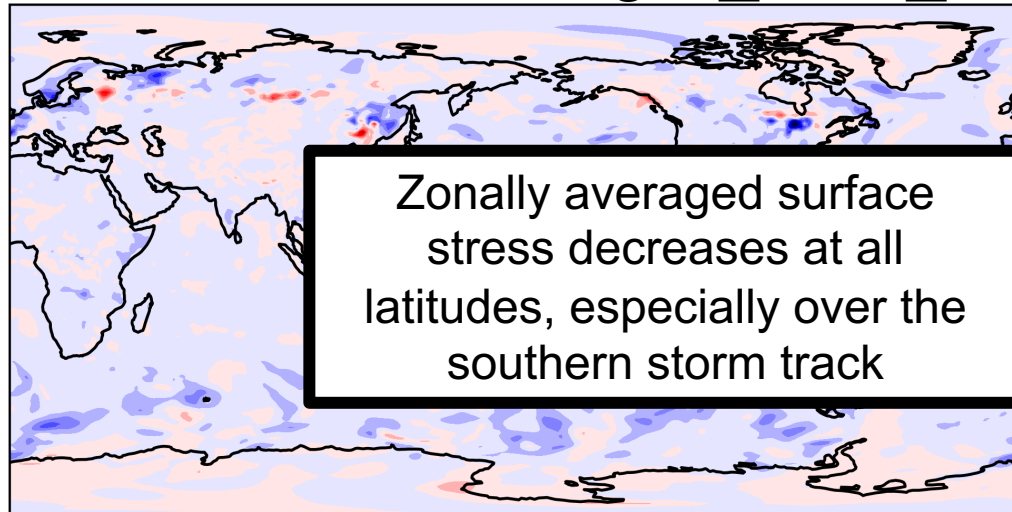
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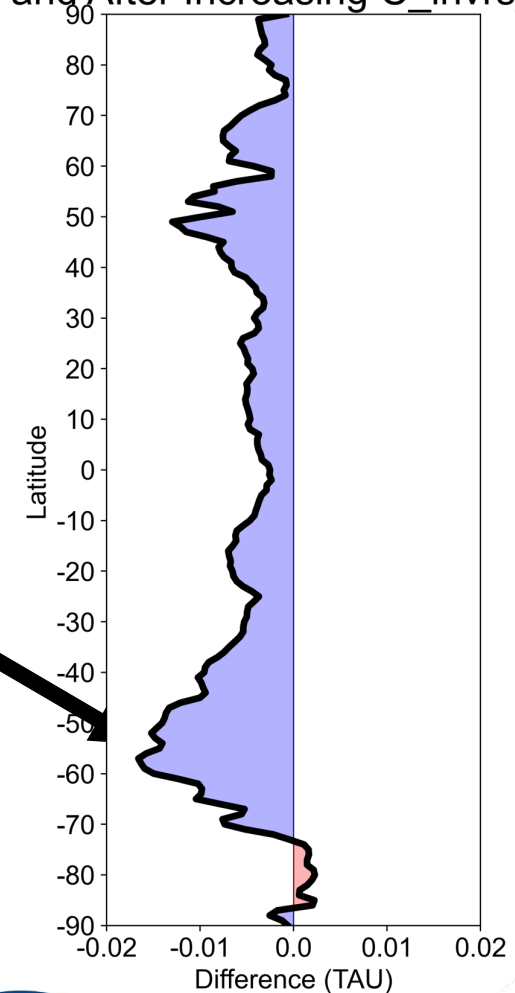
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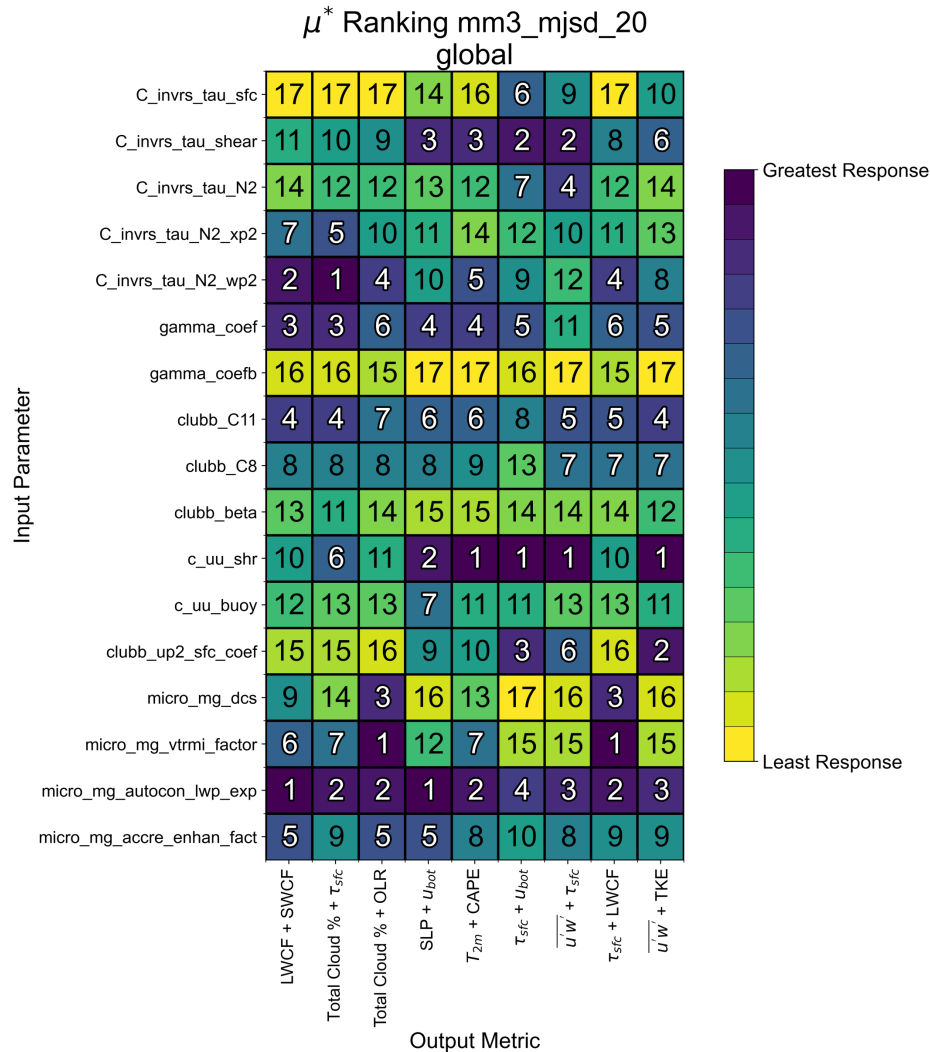


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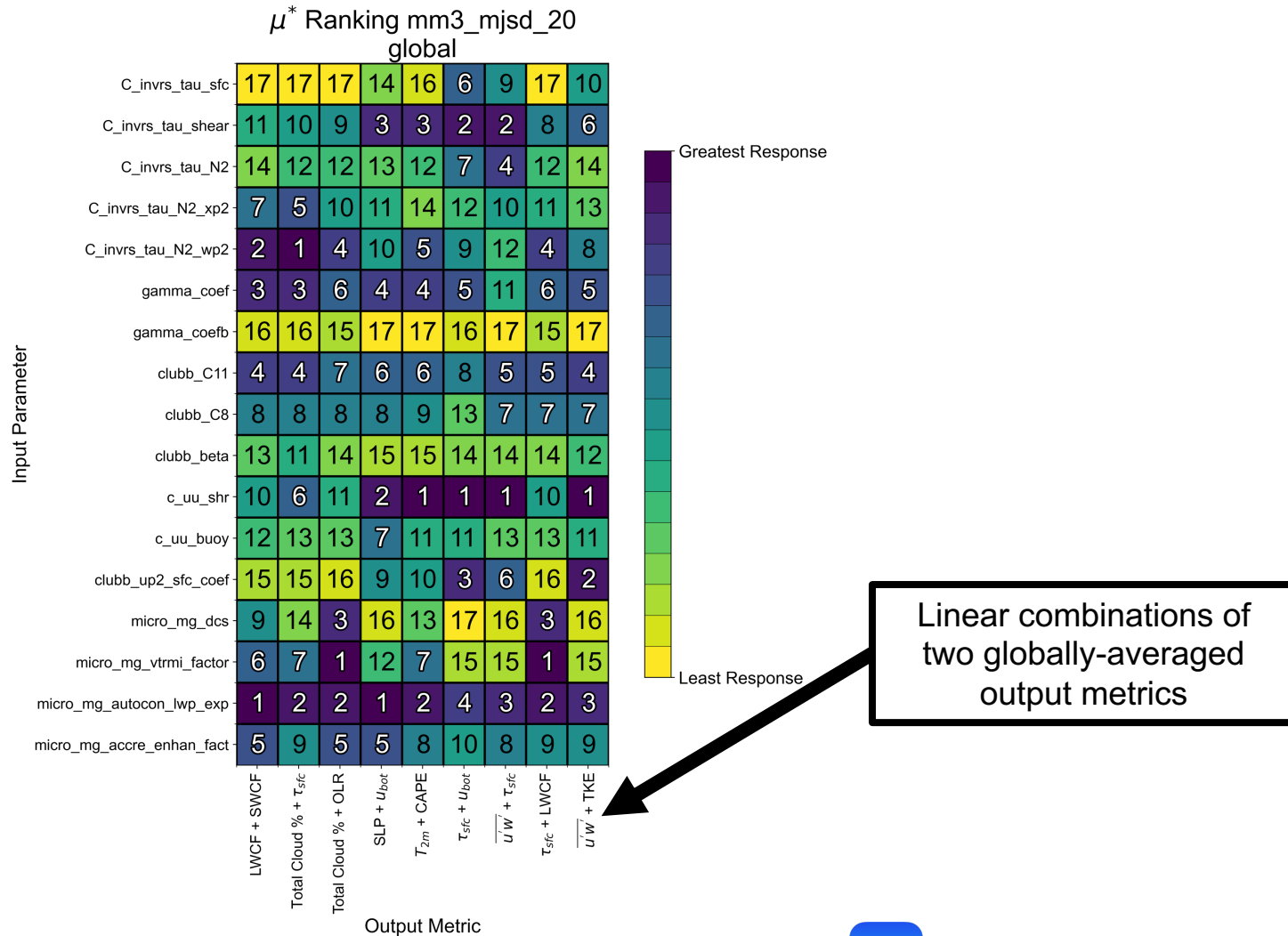
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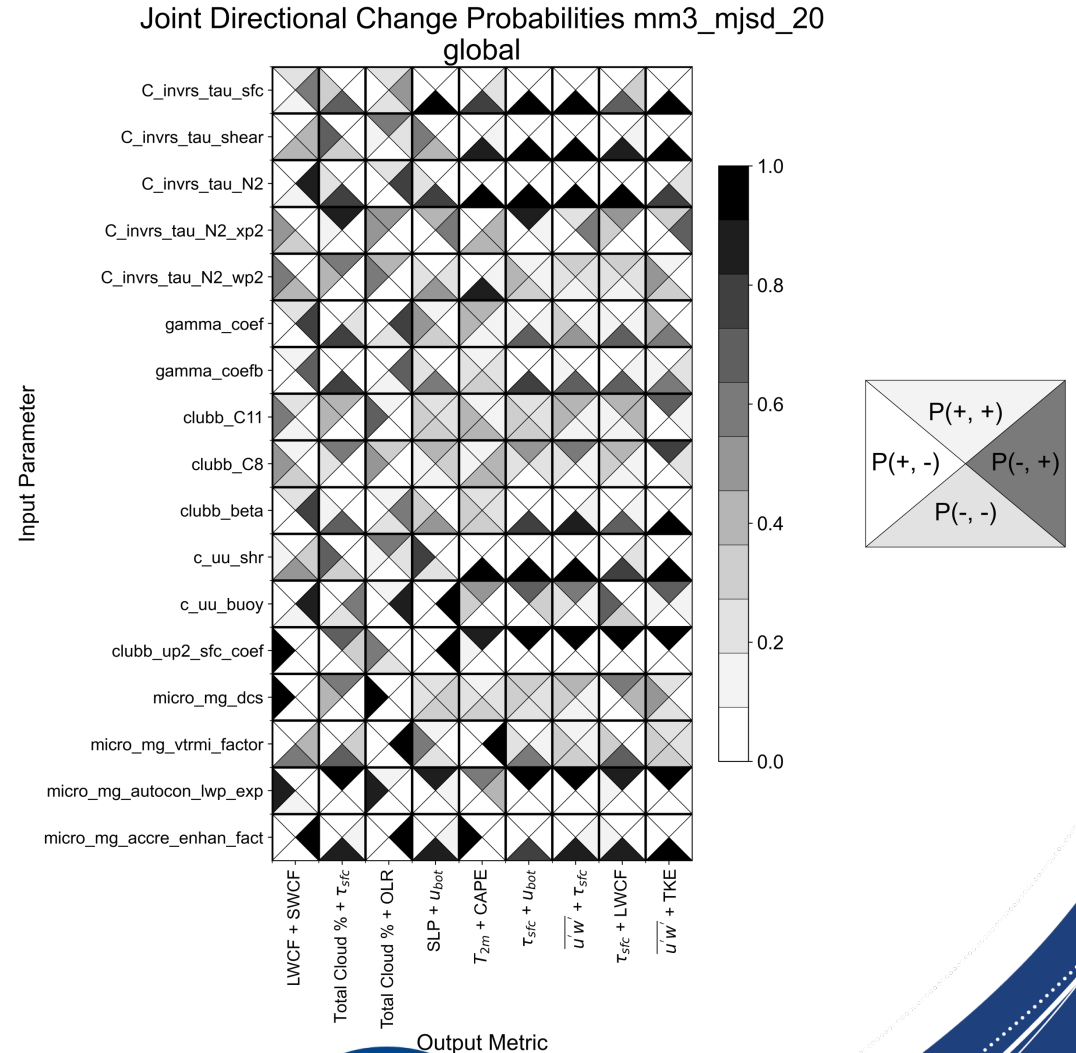
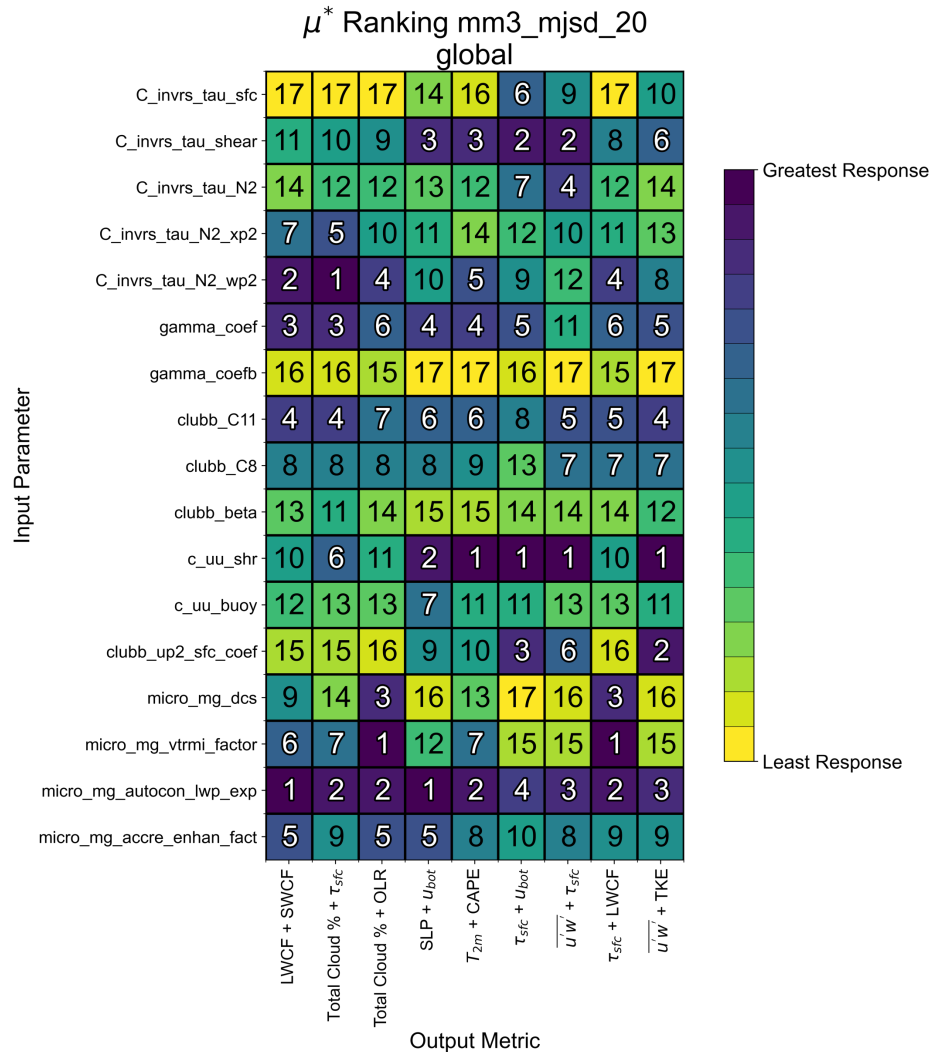
Ongoing Work: we can repeat the sensitivity analysis for combined globally-averaged metrics



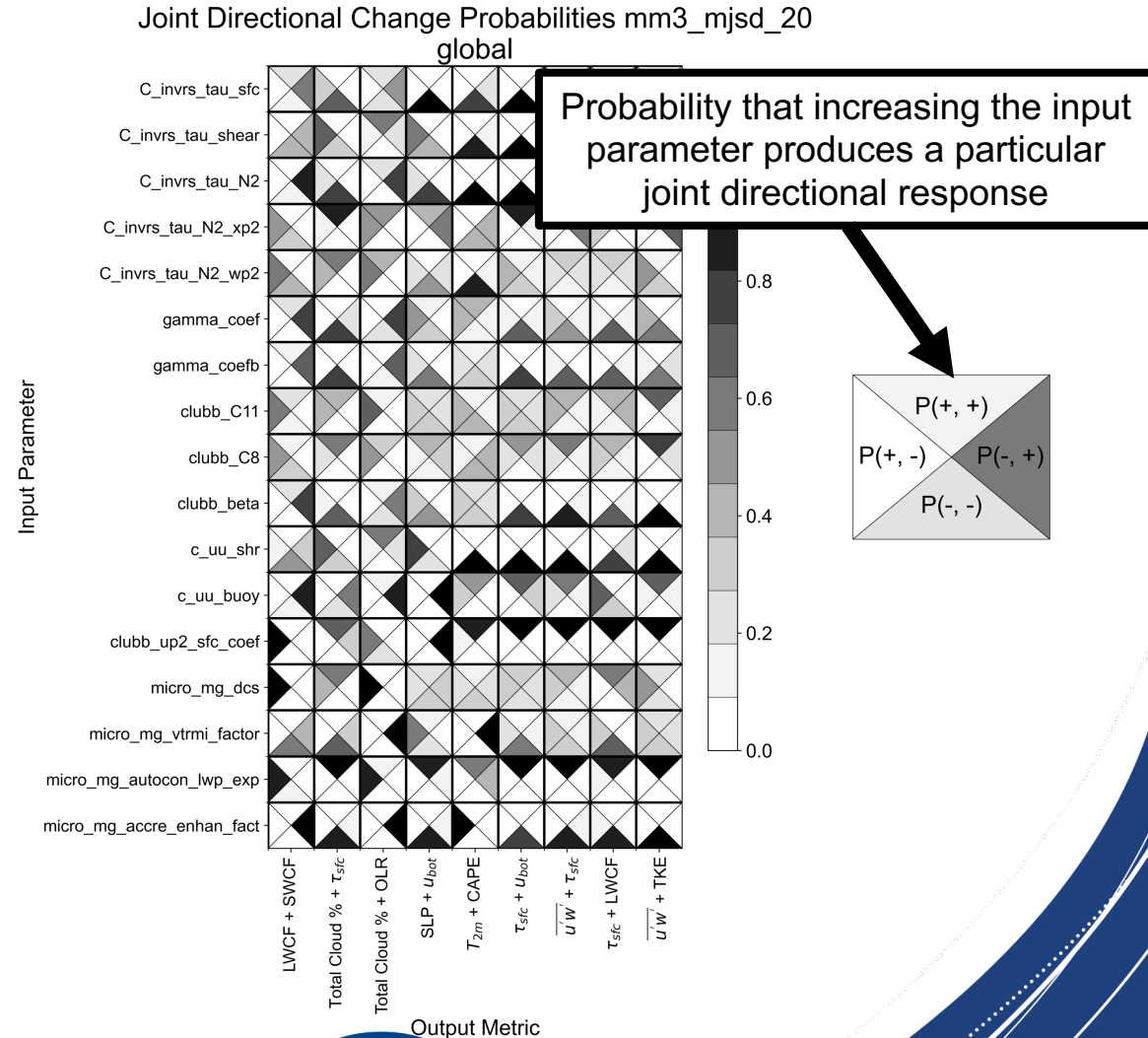
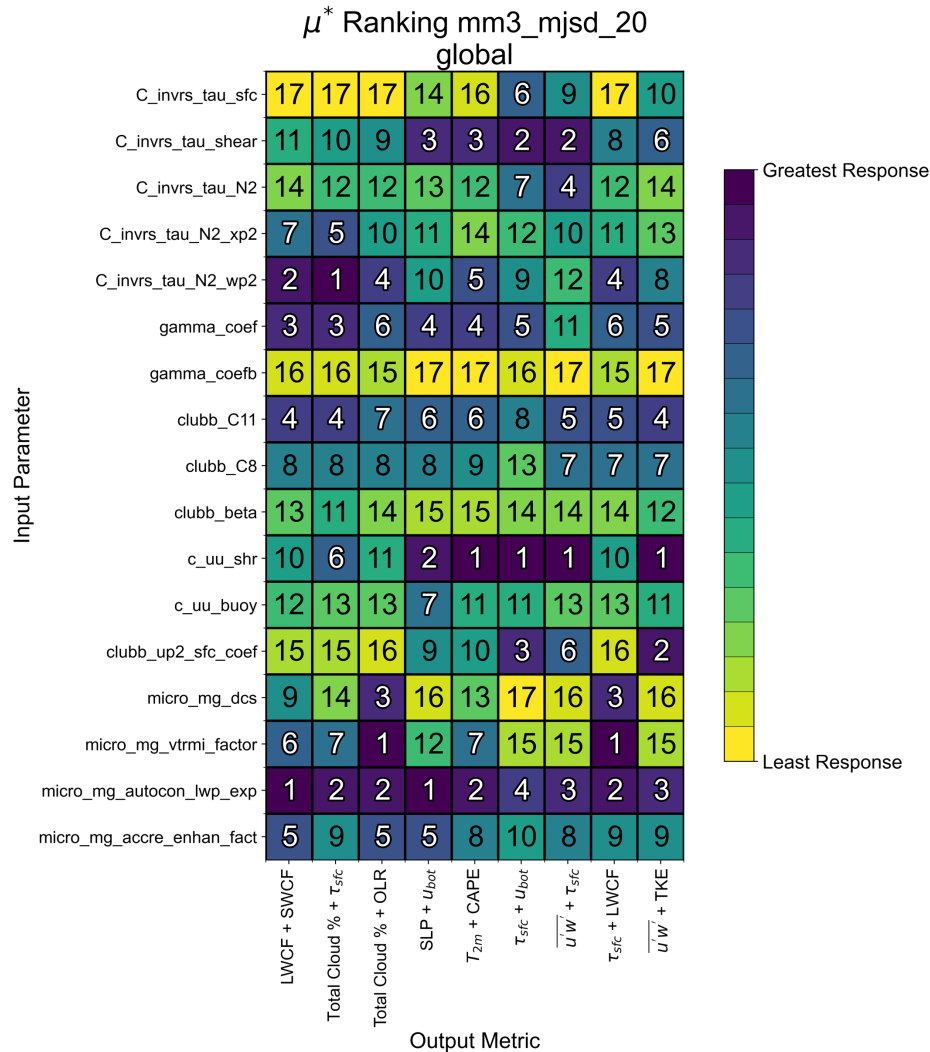
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Main Takeaways

Additional Questions?



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Main Takeaways

- We can use a sensitivity analysis to highlight input parameters in CLUBB that influence regional biases in fields like SWCF (Tropical Pacific) and surface wind stress (Southern Ocean)

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- The impacts of perturbing these input parameters are related changes in budgets of turbulent fluxes like heat/moisture flux and momentum flux

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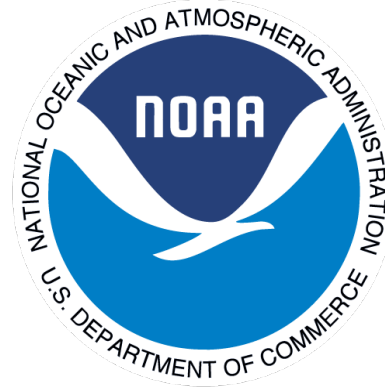
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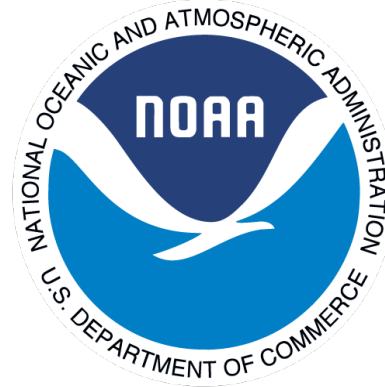
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- **Can we perturb a handful of input parameters in a 10-year, free-running simulation and reduce certain regional biases?**

We thank our partners in this work



PennState

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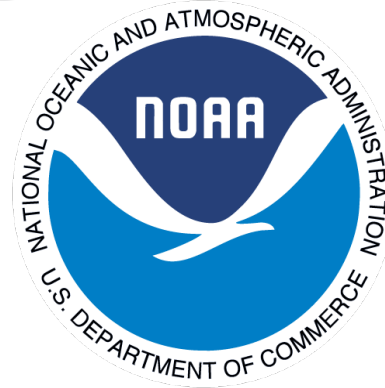
For more details about this work...

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Main Takeaways

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Extra Slides

CLUBBX's formulation of inverse eddy turnover time scale depends on environmental conditions



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$$L = \tau \bar{e}^{\frac{1}{2}} \quad \leftarrow$$

Vertical turbulent length scale is the product of the eddy turnover time scale and the square root of TKE



CLUBBX's formulation of inverse eddy turnover time scale depends on environmental conditions

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Where the eddy time scale is the sum of dissipating processes...



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Where the eddy time scale is the sum of dissipating processes...

$$\frac{1}{\tau} = C_{bkgnd} \frac{1}{\alpha} + C_{sfc} \frac{u^*}{\kappa} \frac{1}{(z - z_{sfc} + d)} + C_{shear} \sqrt{\left(\frac{\partial \bar{u}}{\partial z}\right)^2 + \left(\frac{\partial \bar{v}}{\partial z}\right)^2} + C_{N2} \sqrt{N^2}$$



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Background eddy dissipation
at all levels



CLUBBX's formulation of inverse eddy turnover time scale depends on environmental conditions

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Eddy dissipation due to surface effects

CLUBBX's formulation of inverse eddy turnover time scale depends on environmental conditions

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Vertical turbulent length scale is the product of the eddy turnover time scale and the square root of TKE

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Eddy dissipation in sheared flow

CLUBBX's formulation of inverse eddy turnover time scale depends on environmental conditions

$$L = \tau \bar{\epsilon}^{\frac{1}{2}} \longleftarrow$$

Vertical turbulent length scale is the product of the eddy turnover time scale and the square root of TKE

Where the eddy time scale is the sum of dissipating processes...

$$\frac{1}{\tau} = C_{bkgnd} \frac{1}{\alpha} + C_{sfc} \frac{u^*}{\kappa} \frac{1}{(z - z_{sfc} + d)} + C_{shear} \sqrt{\left(\frac{\partial \bar{u}}{\partial z}\right)^2 + \left(\frac{\partial \bar{v}}{\partial z}\right)^2} + C_{N2} \sqrt{N^2}$$

Eddy dissipation in
stable environment

CLUBBX's formulation of inverse eddy turnover time scale depends on environmental conditions

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Vertical turbulent length scale is the product of the eddy turnover time scale and the square root of TKE

Where the eddy time scale is the sum of dissipating processes...

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- The coefficients attached to each term on the RHS are tunable within CLUBBX



CLUBBX's formulation of inverse eddy turnover time scale depends on environmental conditions

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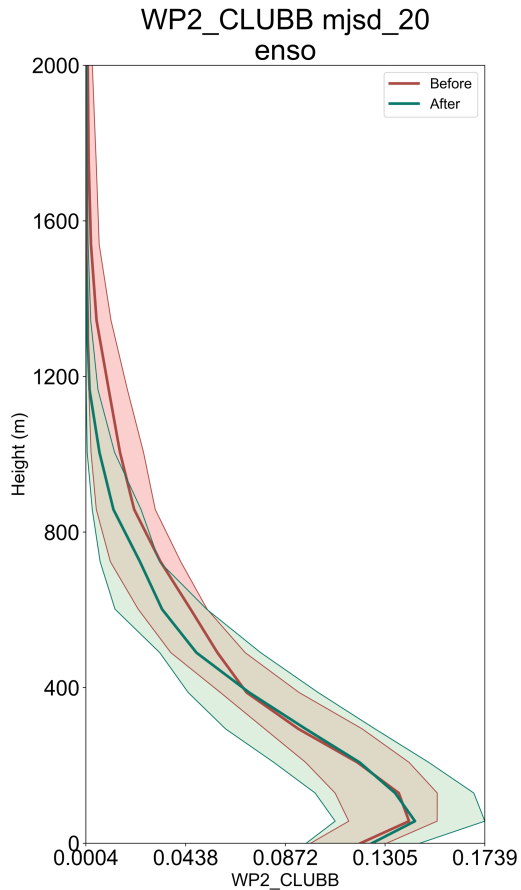
- The coefficients attached to each term on the RHS are tunable within CLUBBX
- This allows the dissipation of turbulent eddies to be tailored to a specific atmospheric regime (e.g., stable boundary layer)



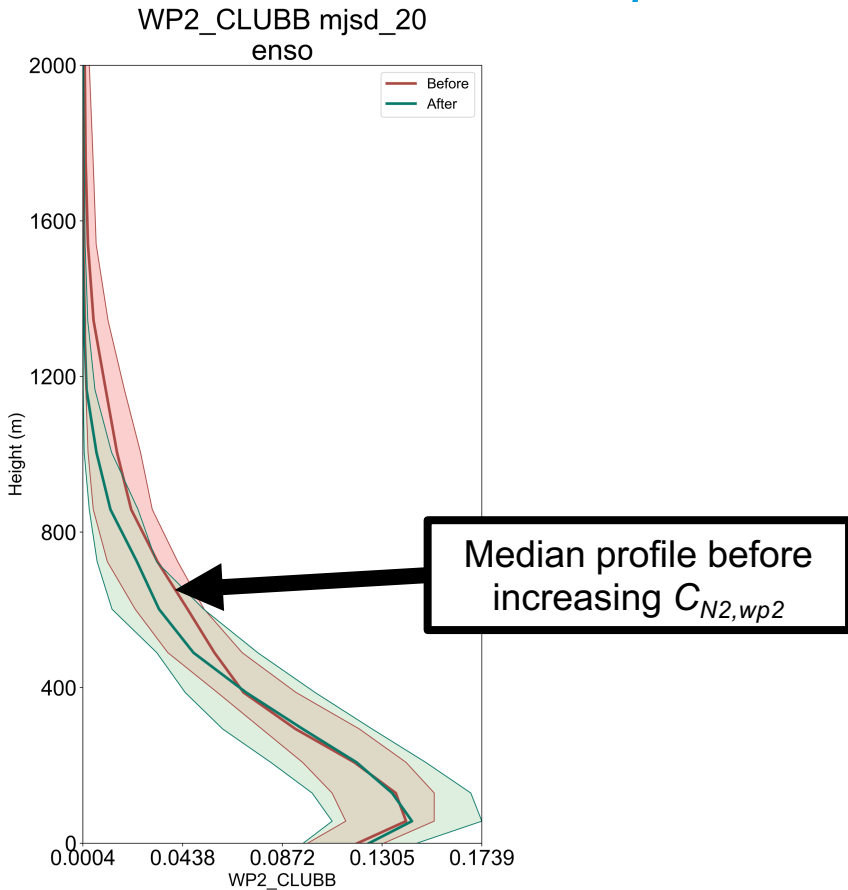
We can track the impact of increasing one of these terms ($C_{N2,wp2}$) through the CLUBBX equations



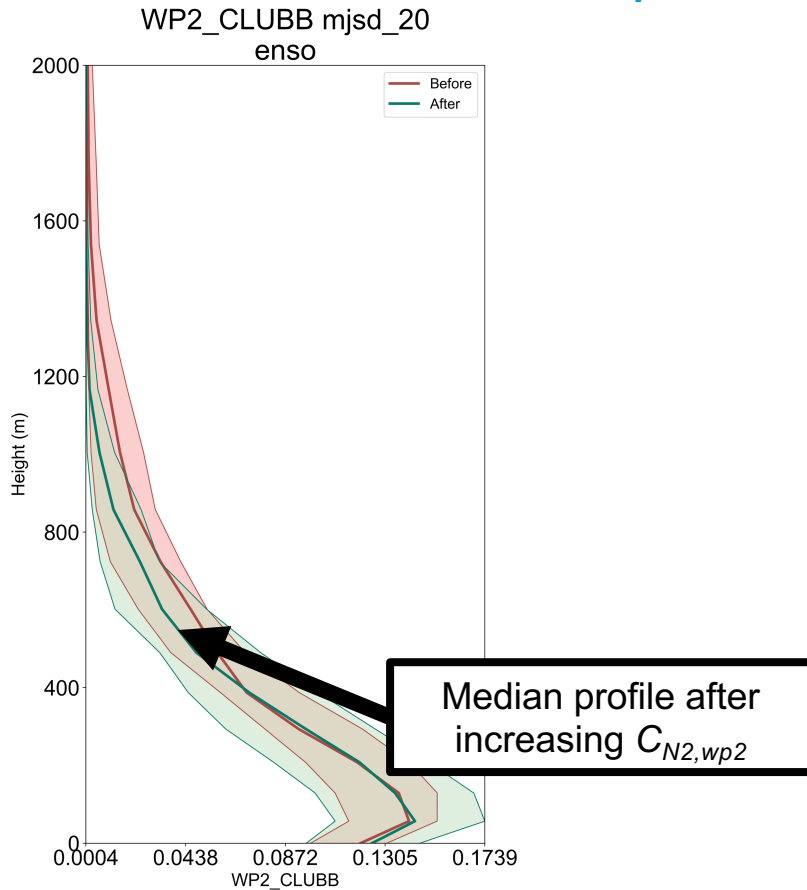
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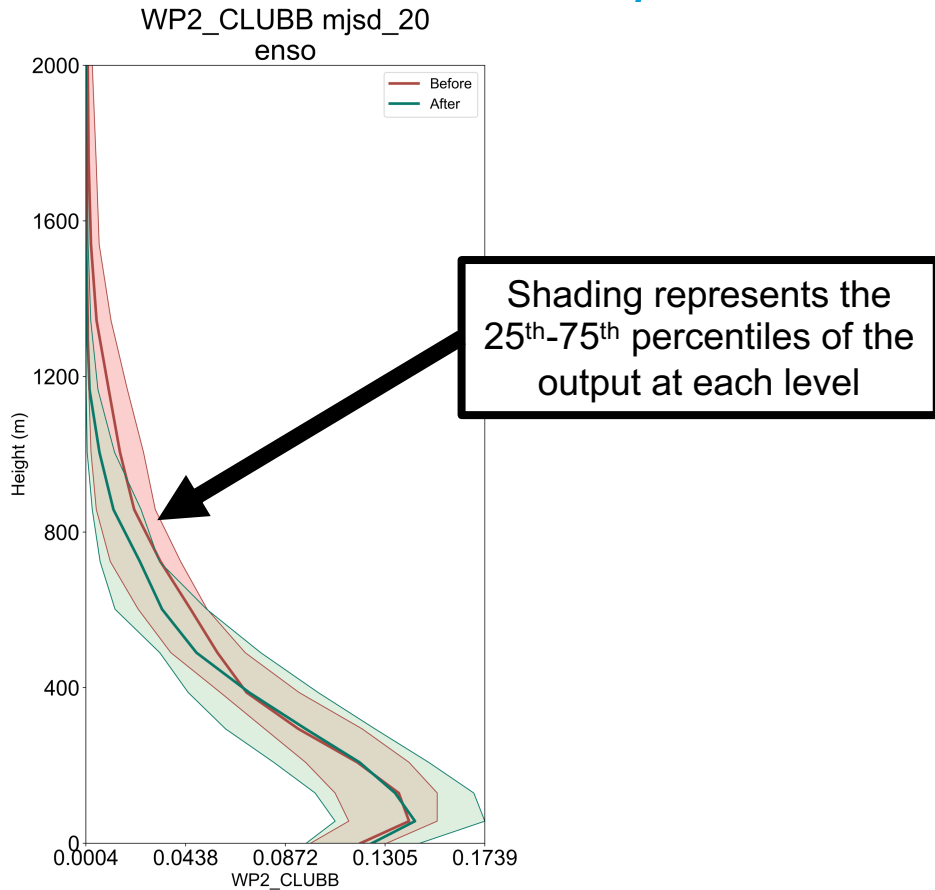
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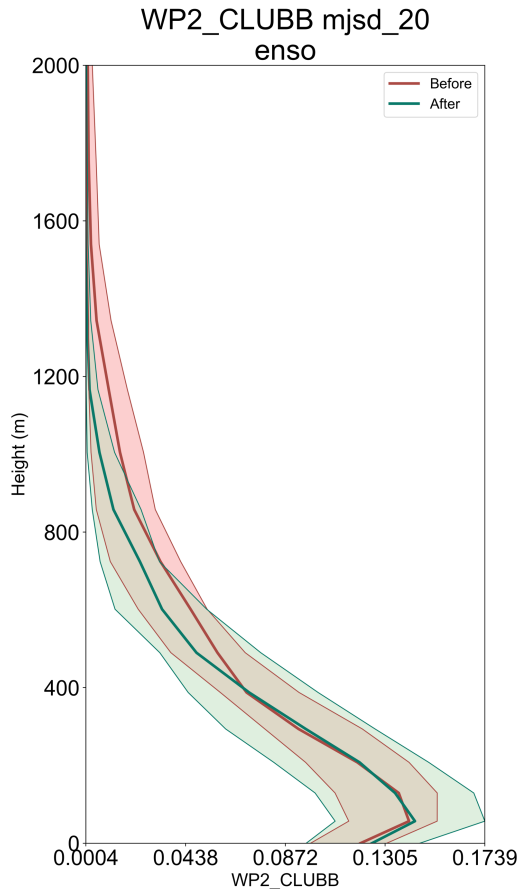
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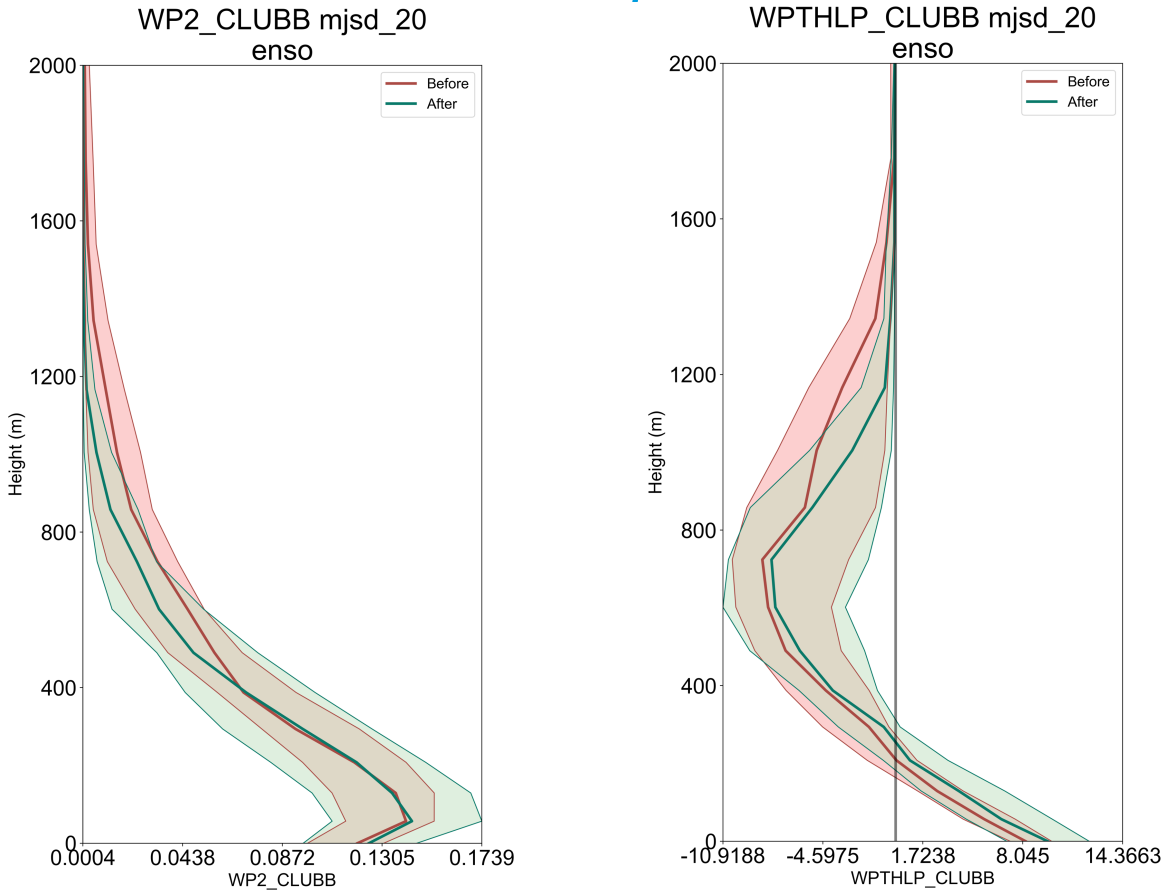
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Increasing $C_{N2,wp2}$ dampens the vertical velocity variance where stable stratification exists, thus reducing wp2



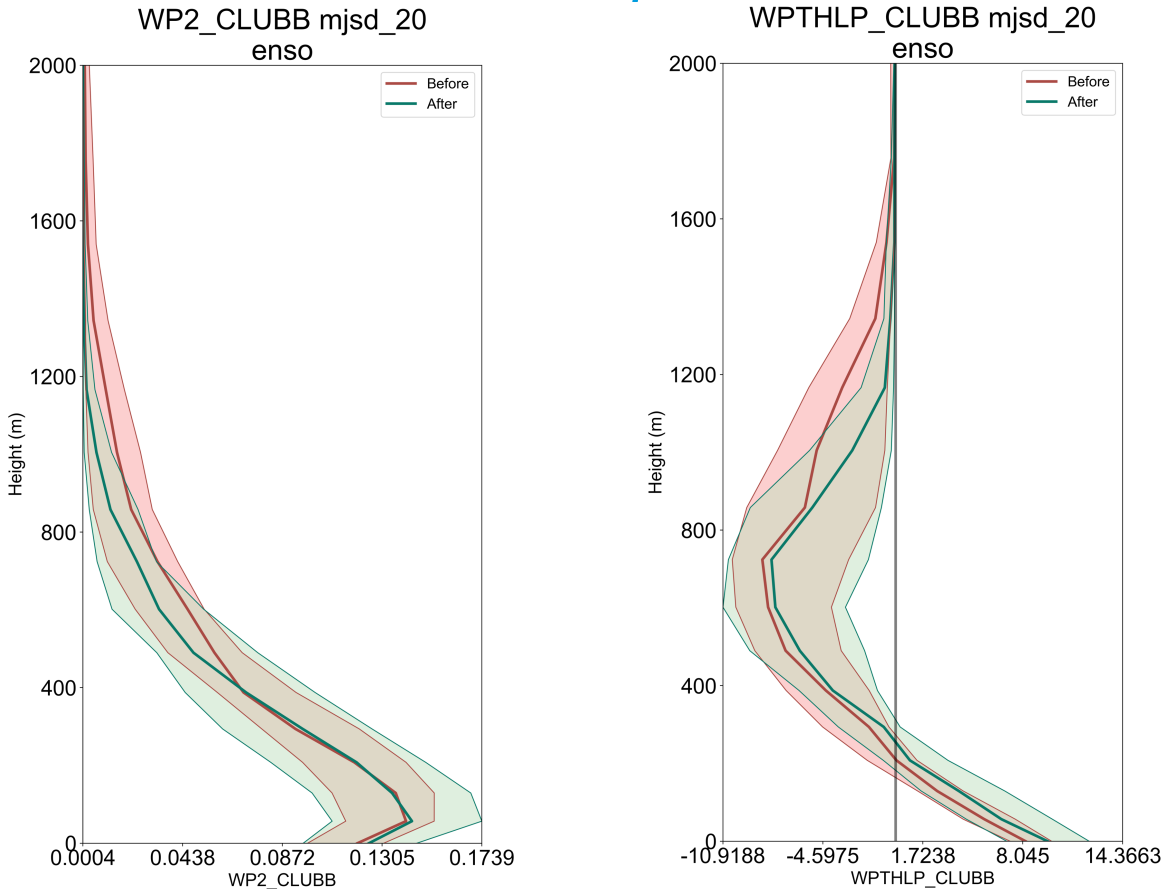
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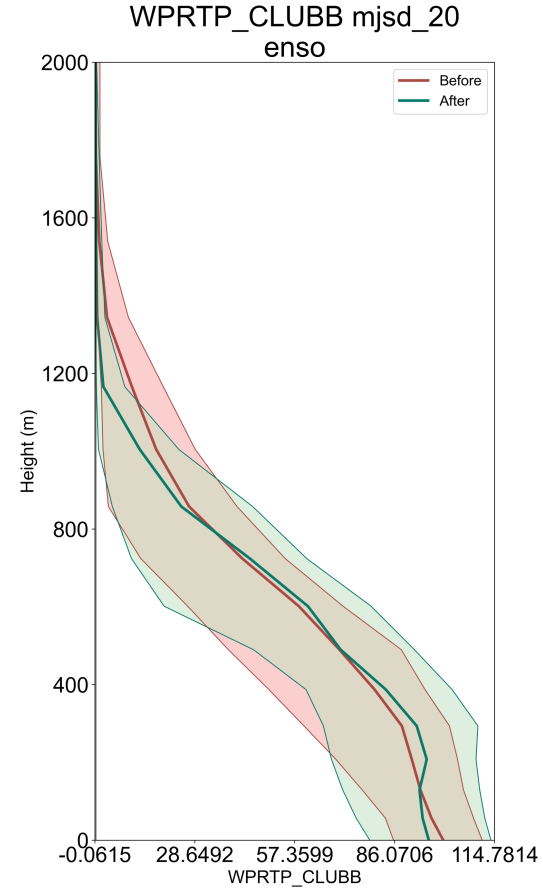
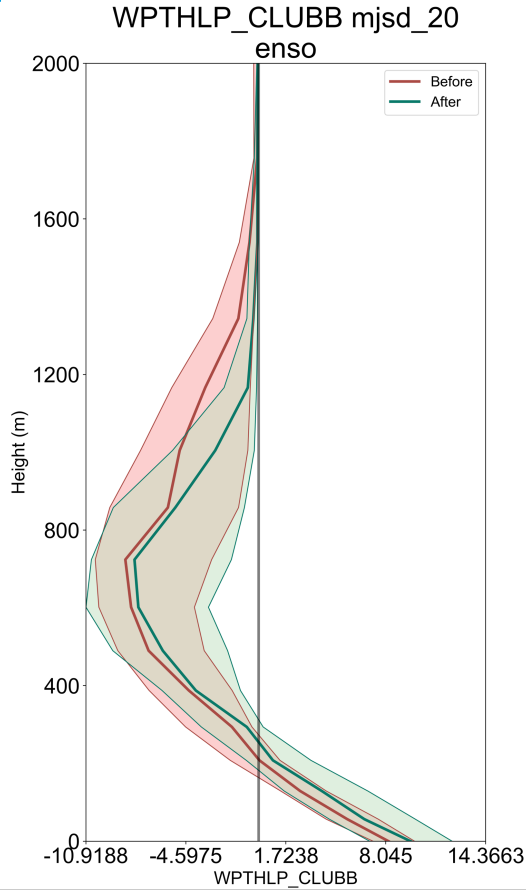
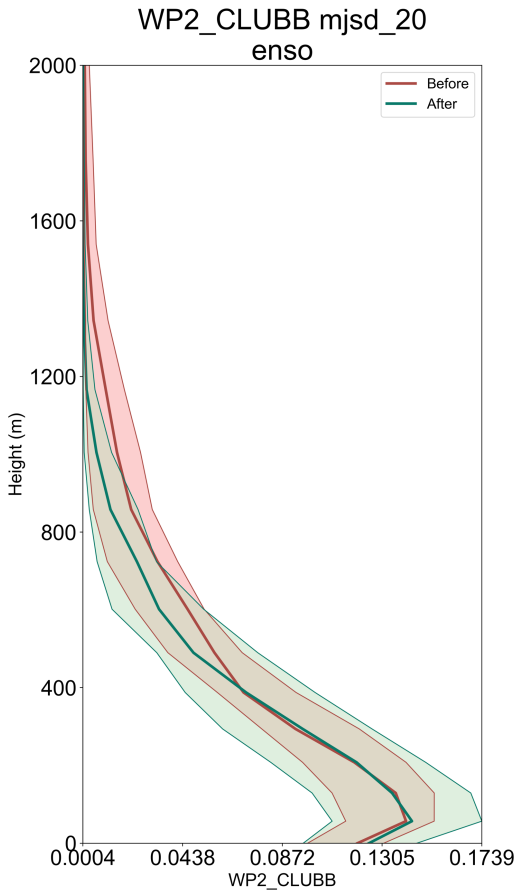


Increasing $C_{N2,wp2}$ dampens the vertical velocity variance where stable stratification exists, thus reducing wp2

Increasing $C_{N2,wp2}$ decreases the downward transport of higher θ_v air from above (increased stability)



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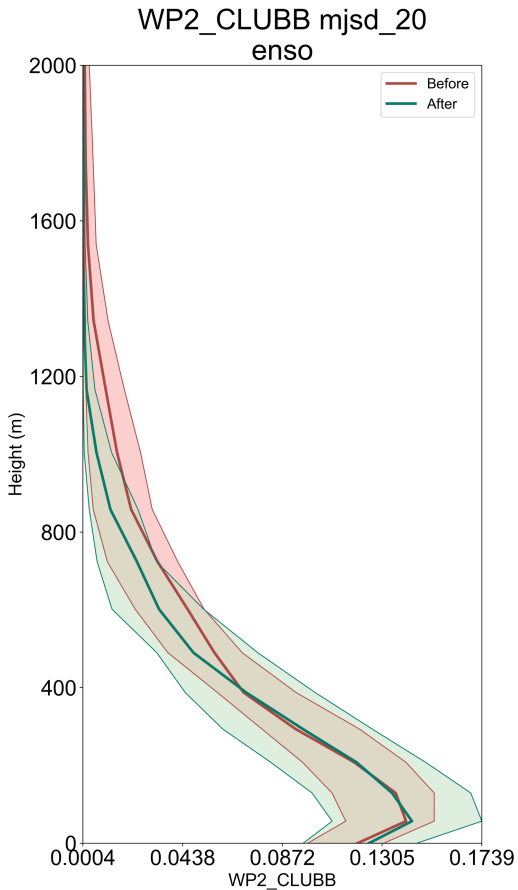


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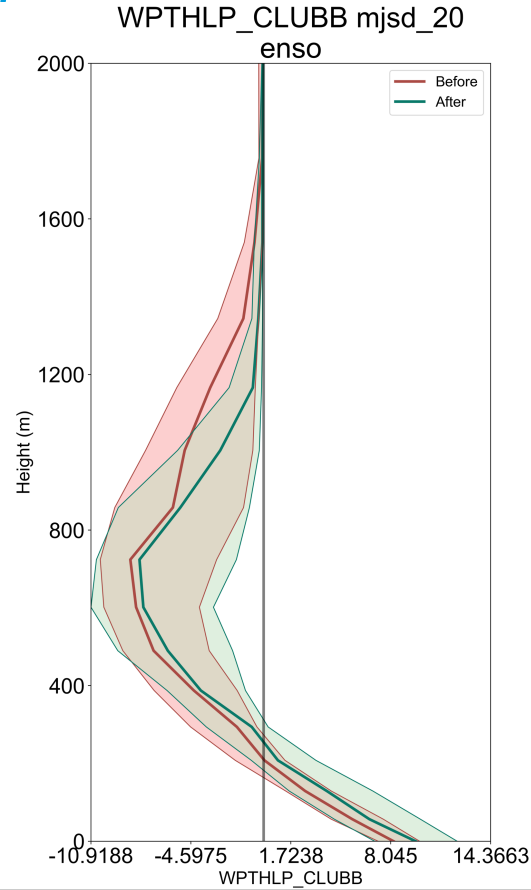
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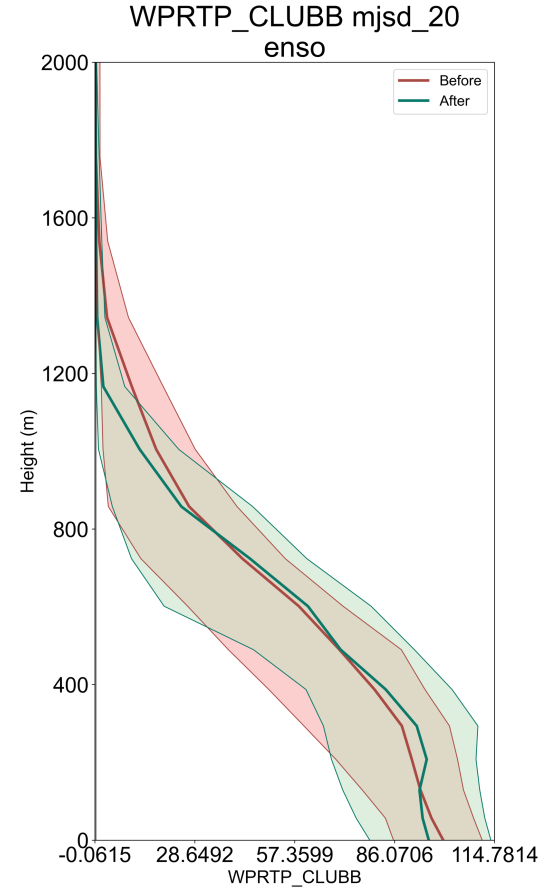
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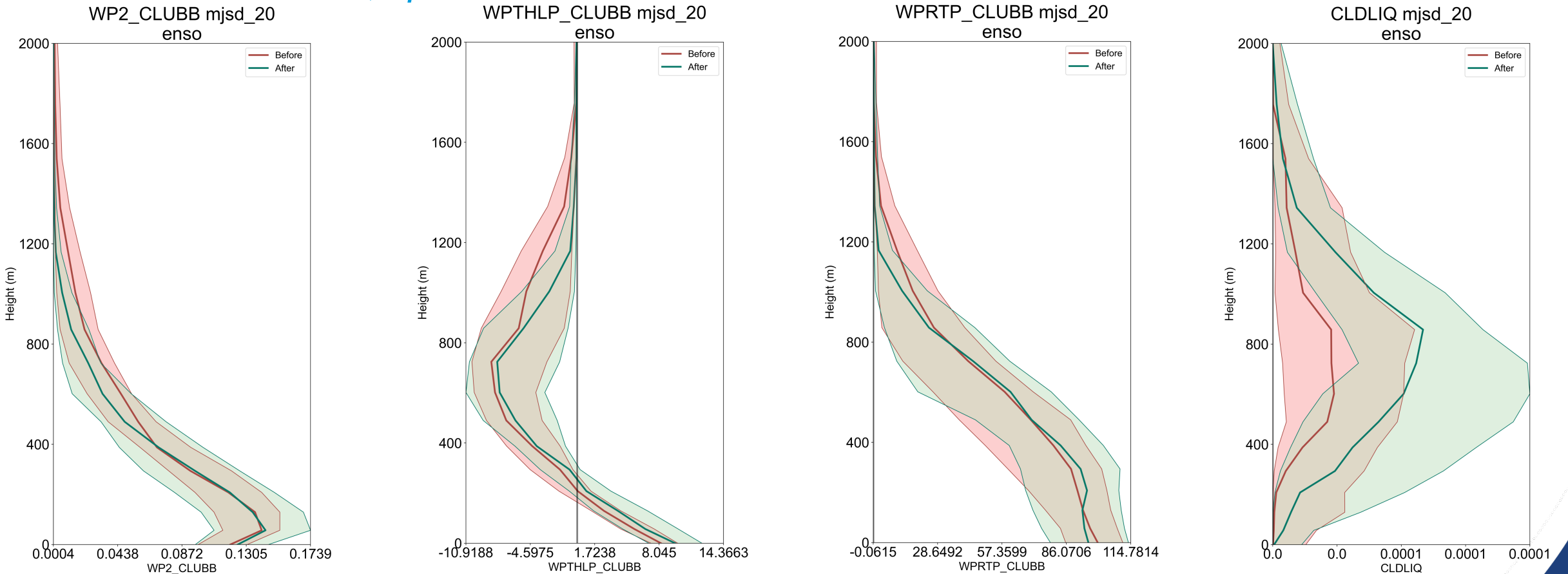
Increasing $C_{N2,wp2}$ decreases the downward transport of higher θ_v air from above (increased stability)



Above about 800 m, there's a decrease in the magnitude of downward transport of drier air from above



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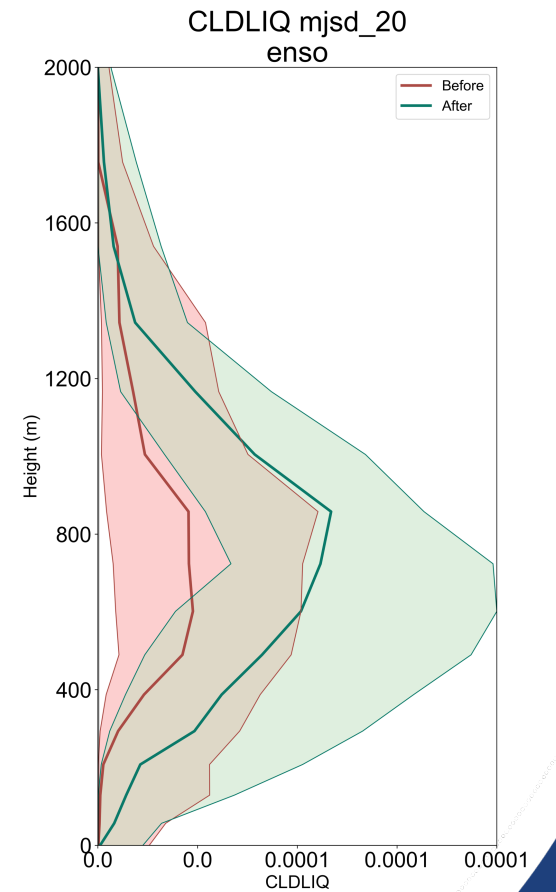
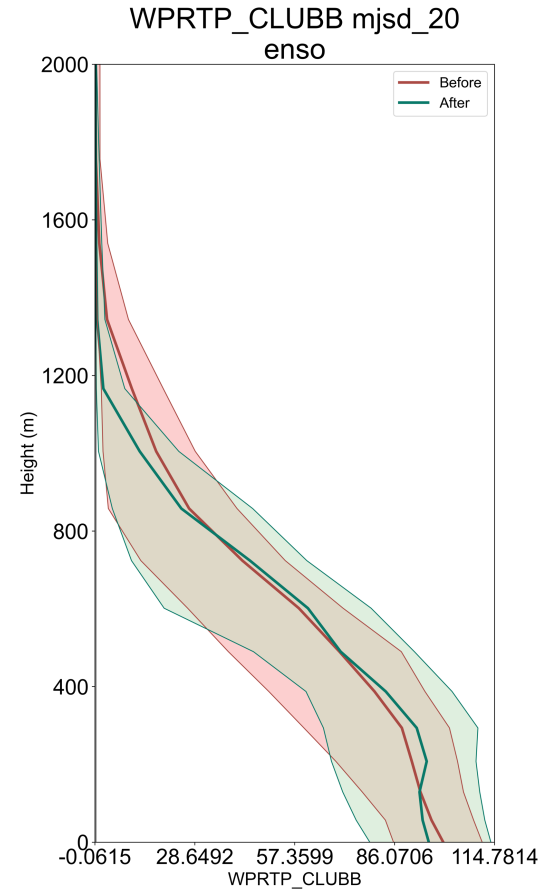
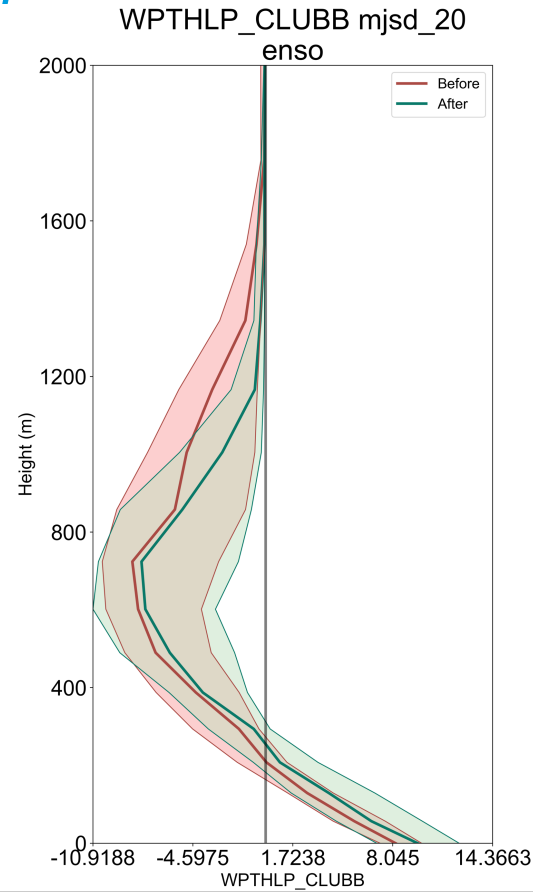
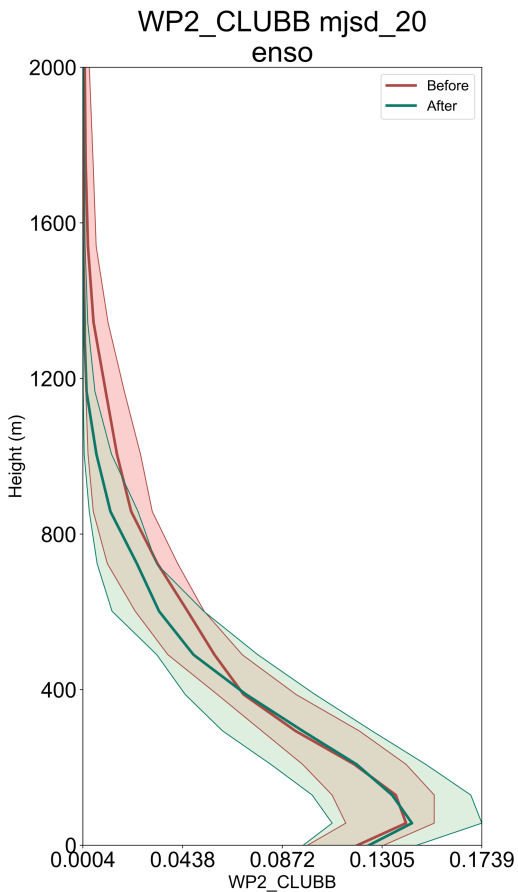
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With reduced downward transport of drier air and increased stability, cloud liquid water content increases



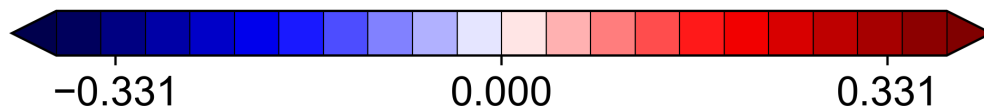
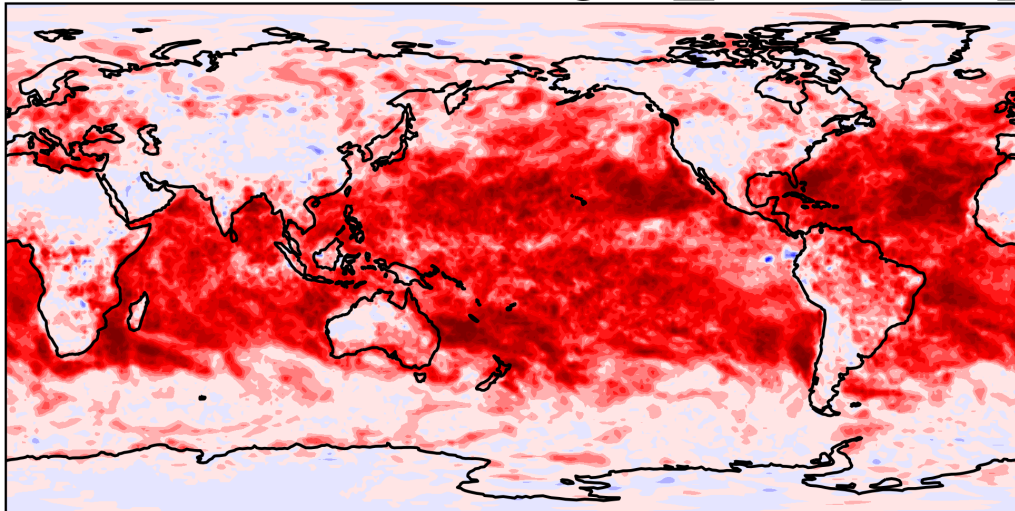
With an increase in cloud liquid water content in the PBL, there's an appreciable increase in low cloud fraction



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Difference in CLDLLOW

Before and After Increasing C_invrstau_N2_wp2



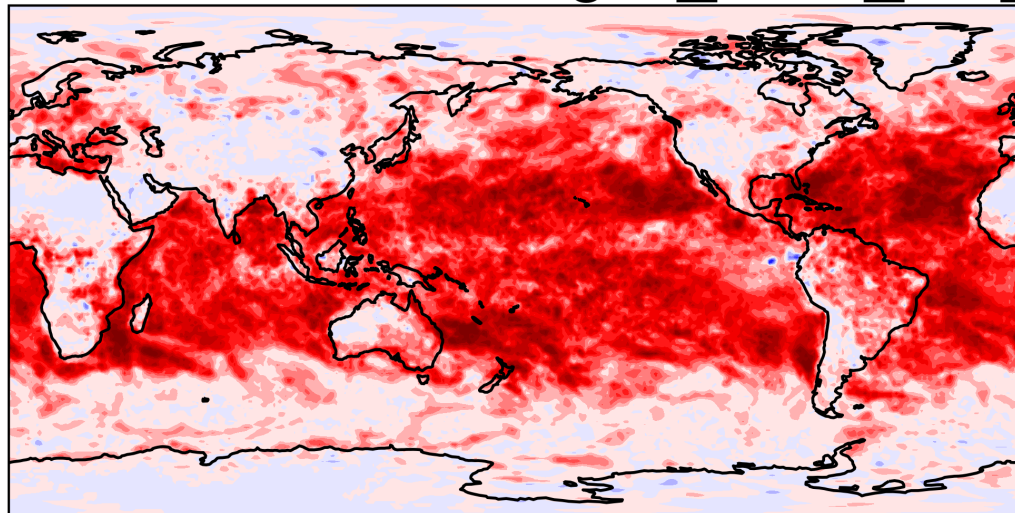
After Avg. = 0.6922 Before Avg. = 0.57557 Diff = 0.11663



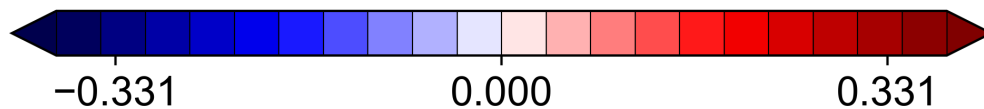
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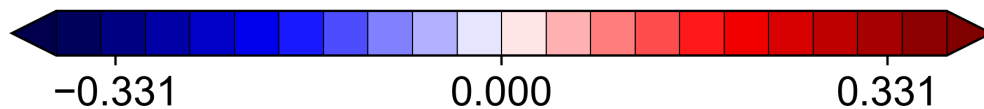
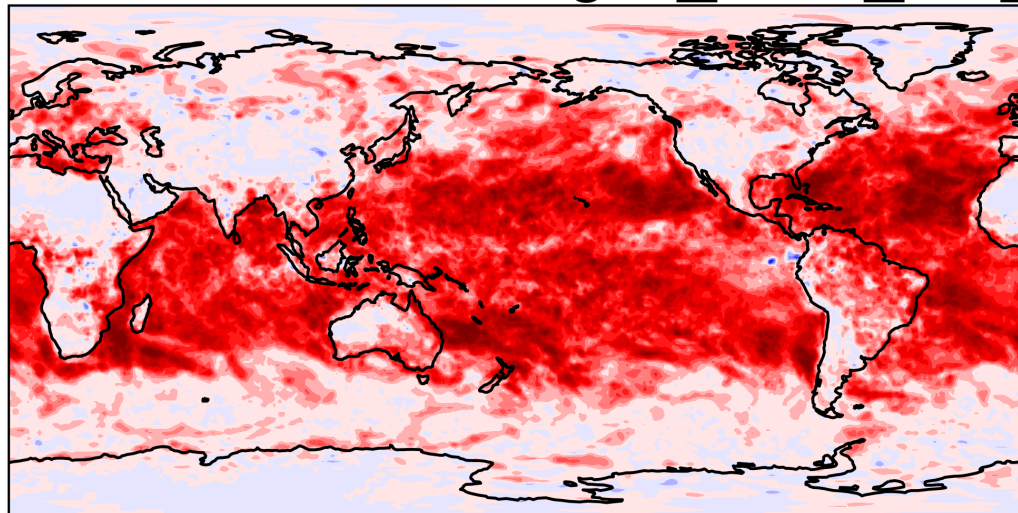
Difference in globally-averaged low cloud between "before" and "after" configs



After Avg. = 0.6922 Before Avg. = 0.57557 **Diff = 0.11663**

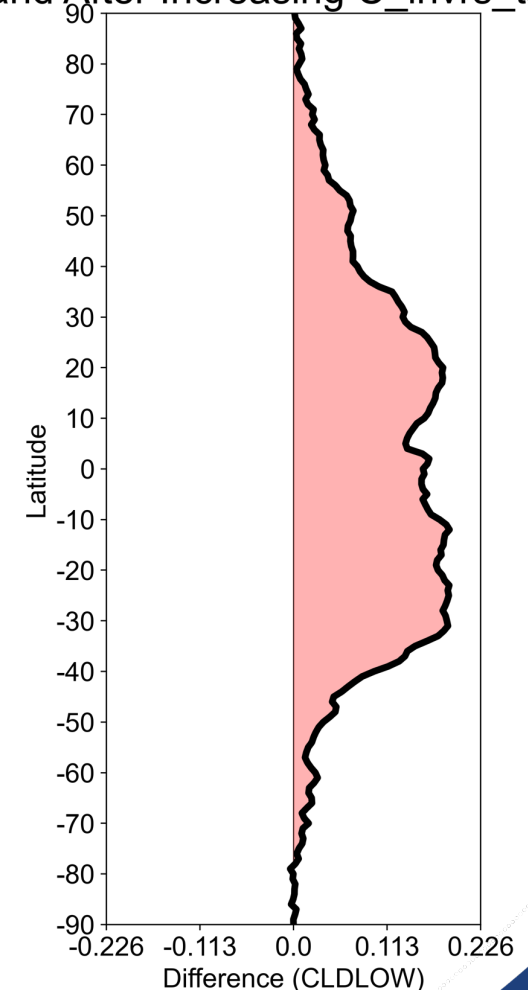
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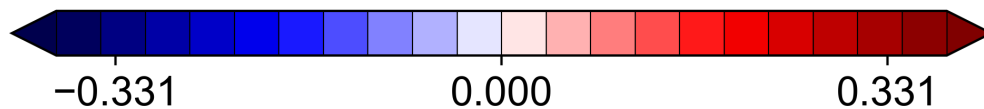
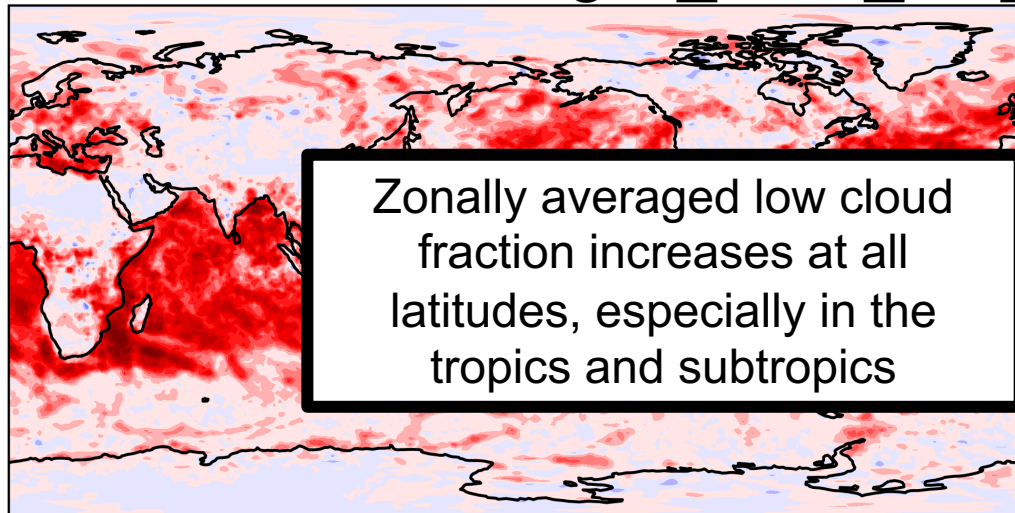
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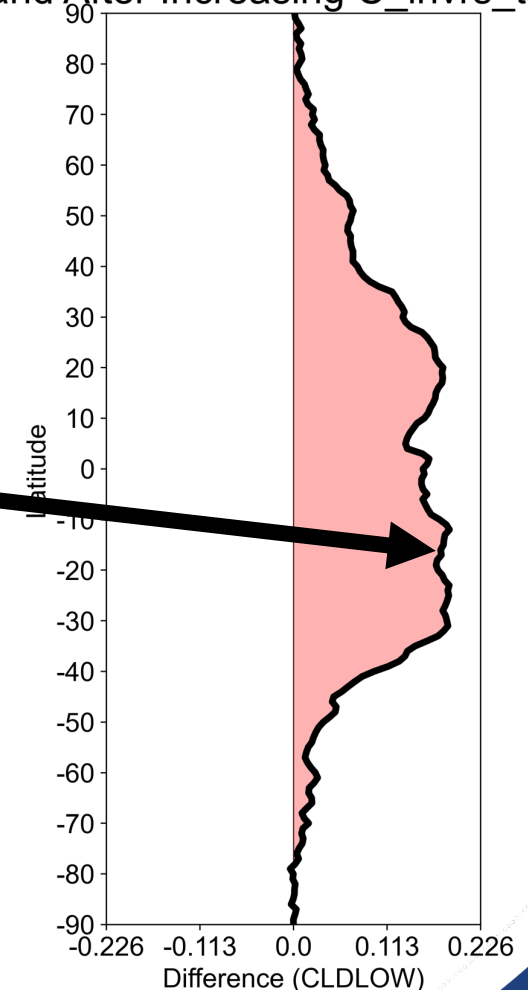
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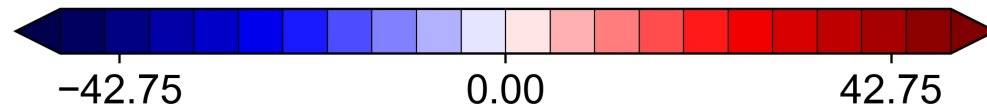
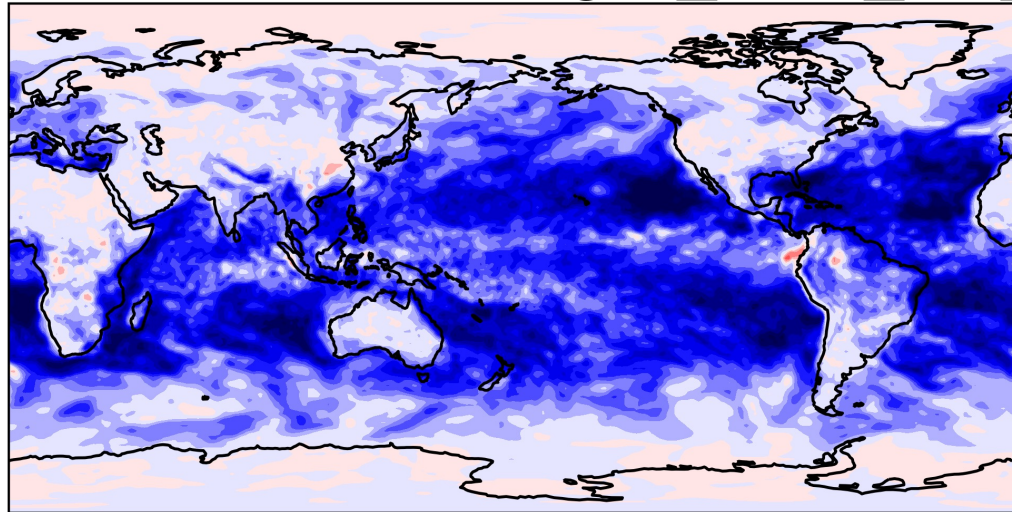
With an increase in low cloud fraction, there's a considerable decrease in SWCF (increased magnitude)



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Difference in SWCF

Before and After Increasing $C_{\text{invers_tau_N2_wp2}}$

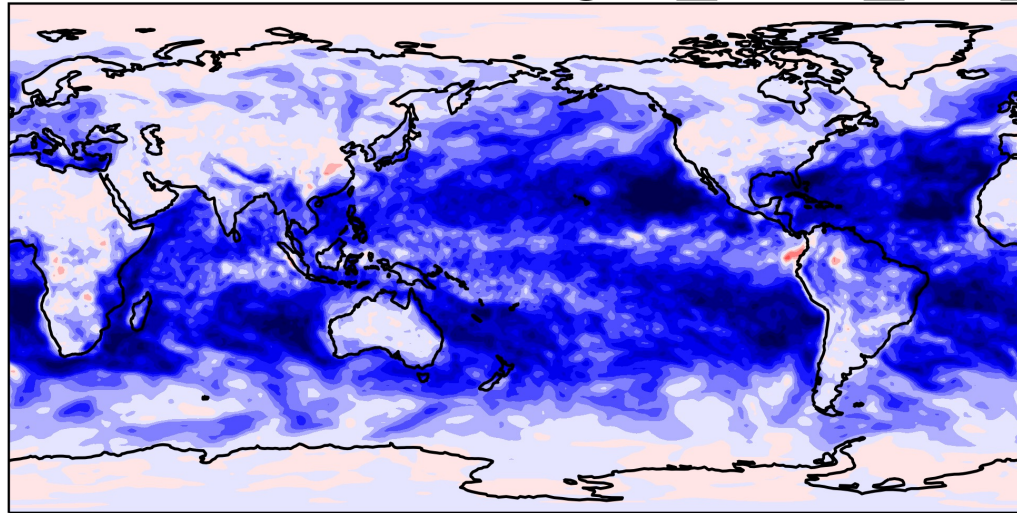


After Avg. = -71.81175 Before Avg. = -55.45565 Diff = -16.3561

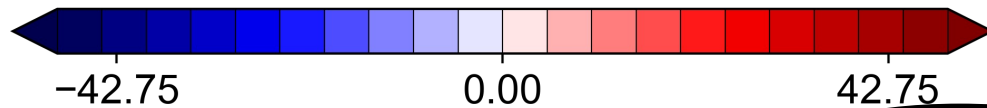
With an increase in low cloud fraction, there's a considerable decrease in SWCF (increased magnitude)

Difference in SWCF

Before and After Increasing C_invrs_tau_N2_wp2



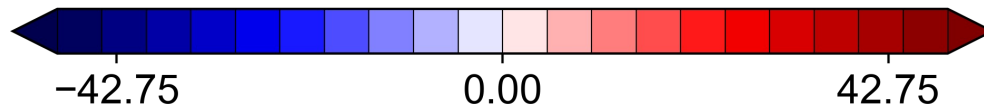
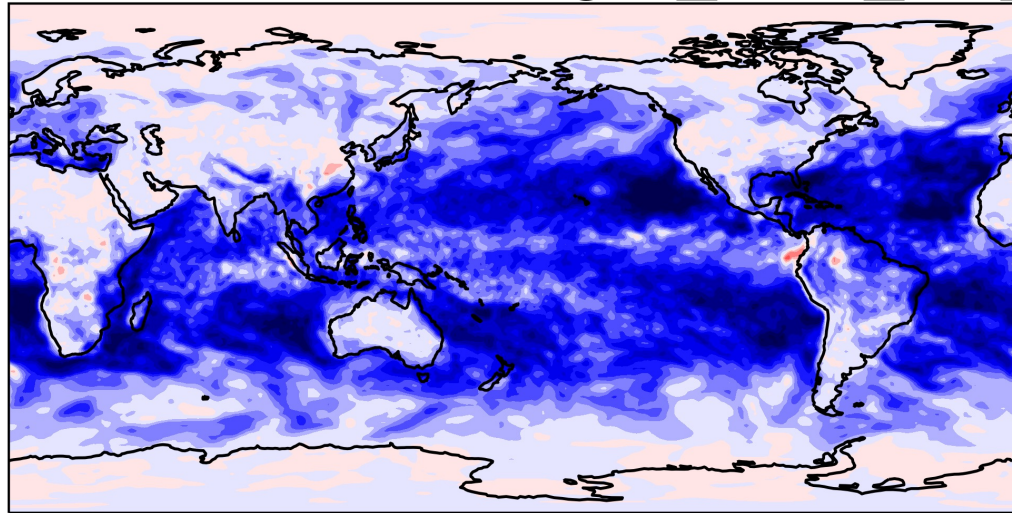
Difference in globally-averaged SWCF between "before" and "after" configs



After Avg. = -71.81175 Before Avg. = -55.45565 Diff = -16.3561

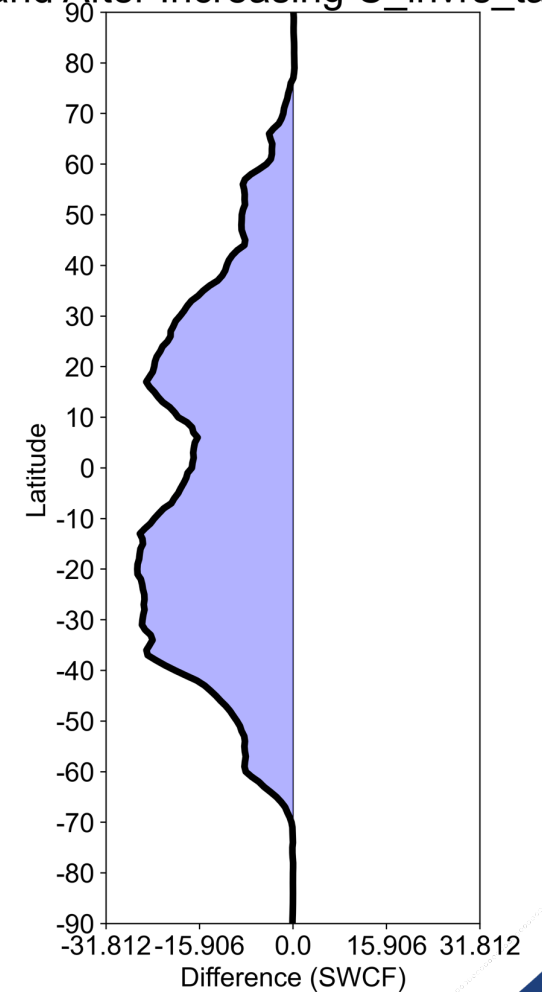
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Difference in SWCF
Before and After Increasing C_invrs_tau_N2_wp2



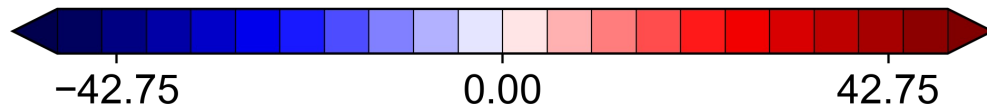
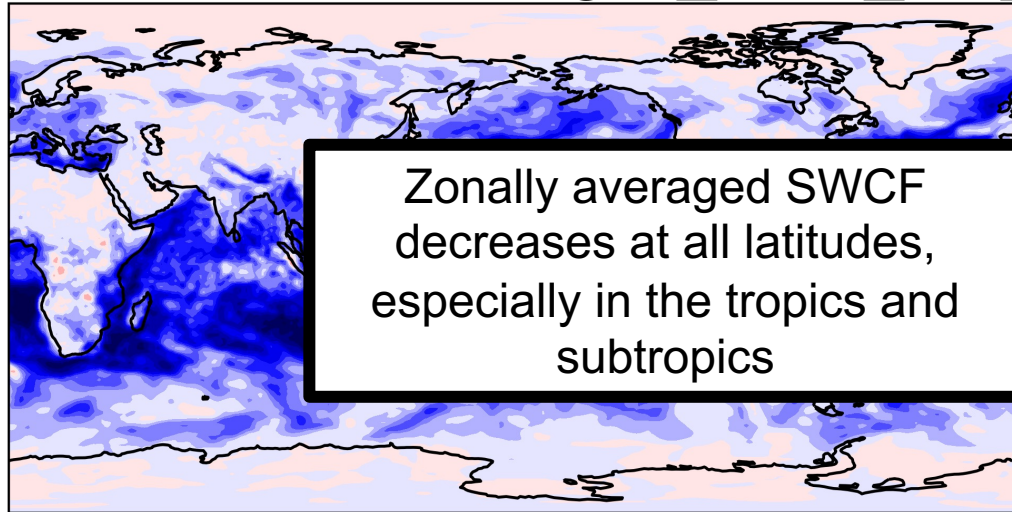
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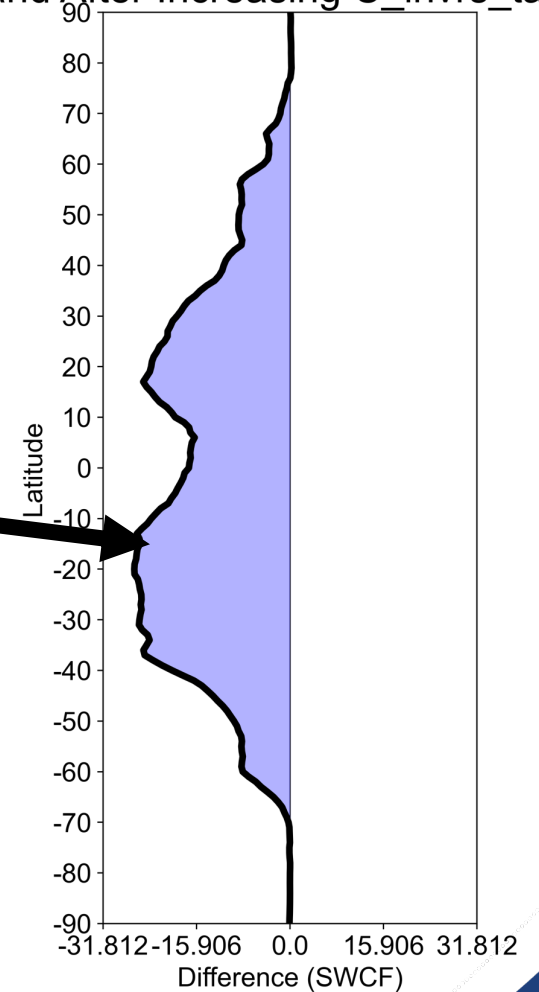
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Difference in SWCF
Before and After Increasing $C_{invs_tau_N2_wp2}$



After Avg. = -71.81175 Before Avg. = -55.45565 Diff = -16.3561

Difference in SWCF
Before and After Increasing $C_{invs_tau_N2_wp2}$



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C_{shear} : weighting term for eddy dissipation from wind shear
Increasing this term **also reduces the vertical turbulent length scale L**

$$C_{shear} \sqrt{\left(\frac{\partial \overline{u}}{\partial z}\right)^2 + \left(\frac{\partial \overline{v}}{\partial z}\right)^2}$$

