

# Wind shear impacts on convection — Implication for convective organization parameterization in climate models

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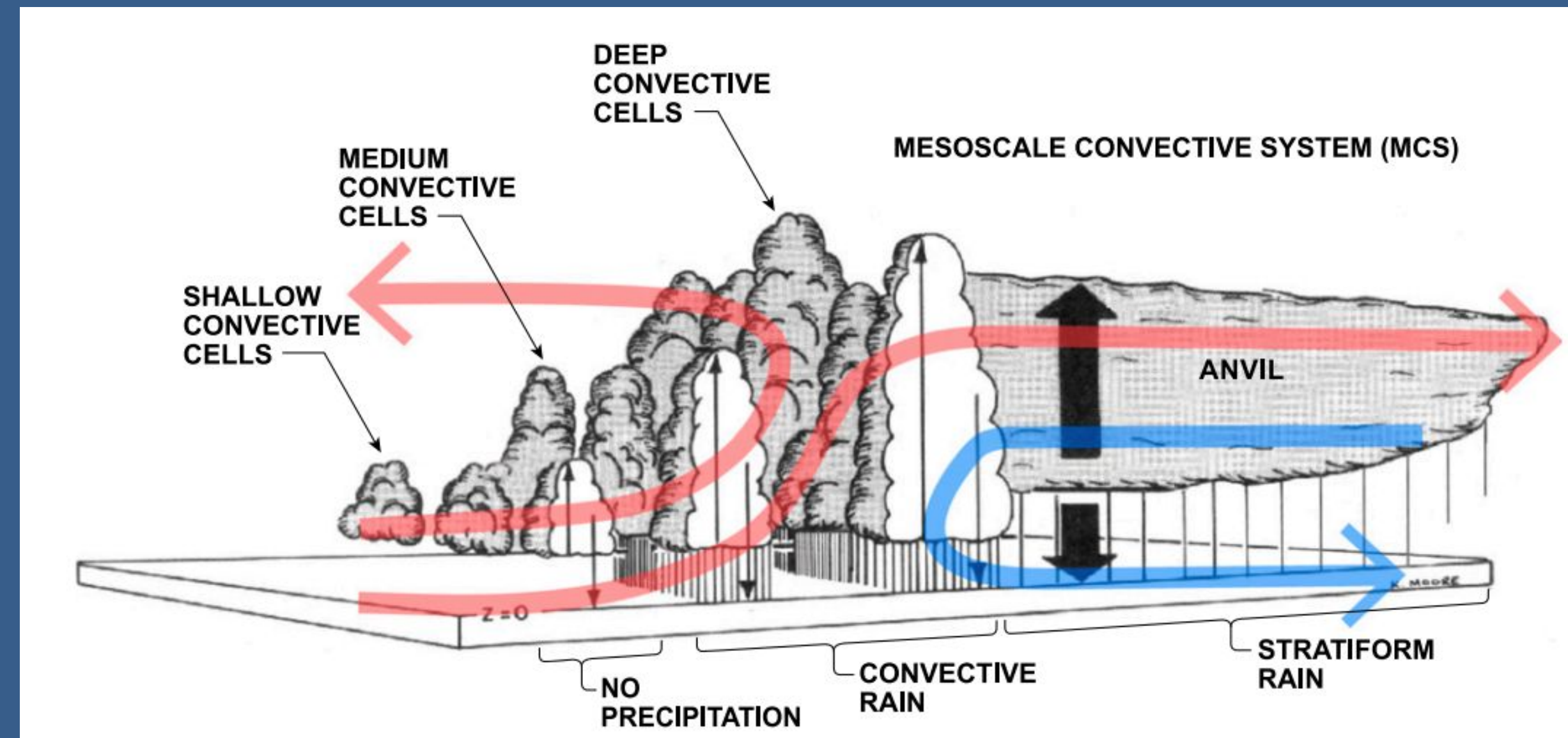
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# Past studies of mesoscale parameterization

## Multiscale Coherent Structure Parameterization (MCSP)



Slantwise layer overturning structure

Donner 1993;

Donner et al. 2001

Mapes and Neale 2011 (*org*)

Moncrieff et al. 2017 (MCSP)

$$Q = Q_c + Q_m$$

$$Q_m(p, t) = -\alpha_1 Q_c(t) \sin 2\pi \left( \frac{p_s - p}{p_s - p_t} \right)$$

Q: Convective heating

Q<sub>c</sub>: Heating from ZM deep convection scheme

Q<sub>m</sub>: Added mesoscale heating, depend on wind shear trigger



# Motivations: LES study of vertical wind shear on cumulus ensembles

- Include a more robust convective organization trigger/response related to wind shear
- Existing literatures are mostly focused on a single squall-line case or idealized warm bubble experiments
- We propose to get an ensemble convective response in a less storm-like and more realistic environment, and inform a physically-based convective organization-wind shear relationship to be used in climate models

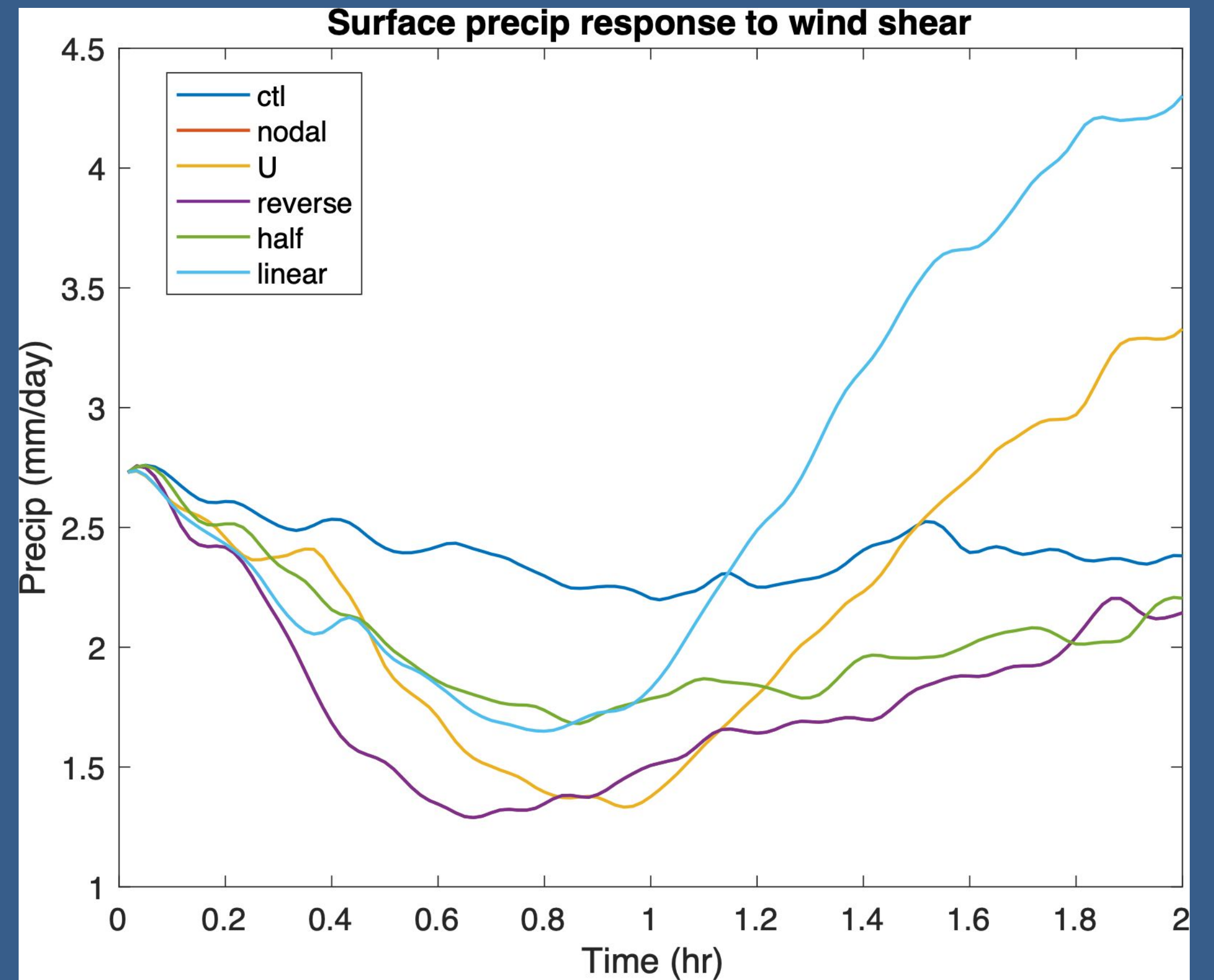
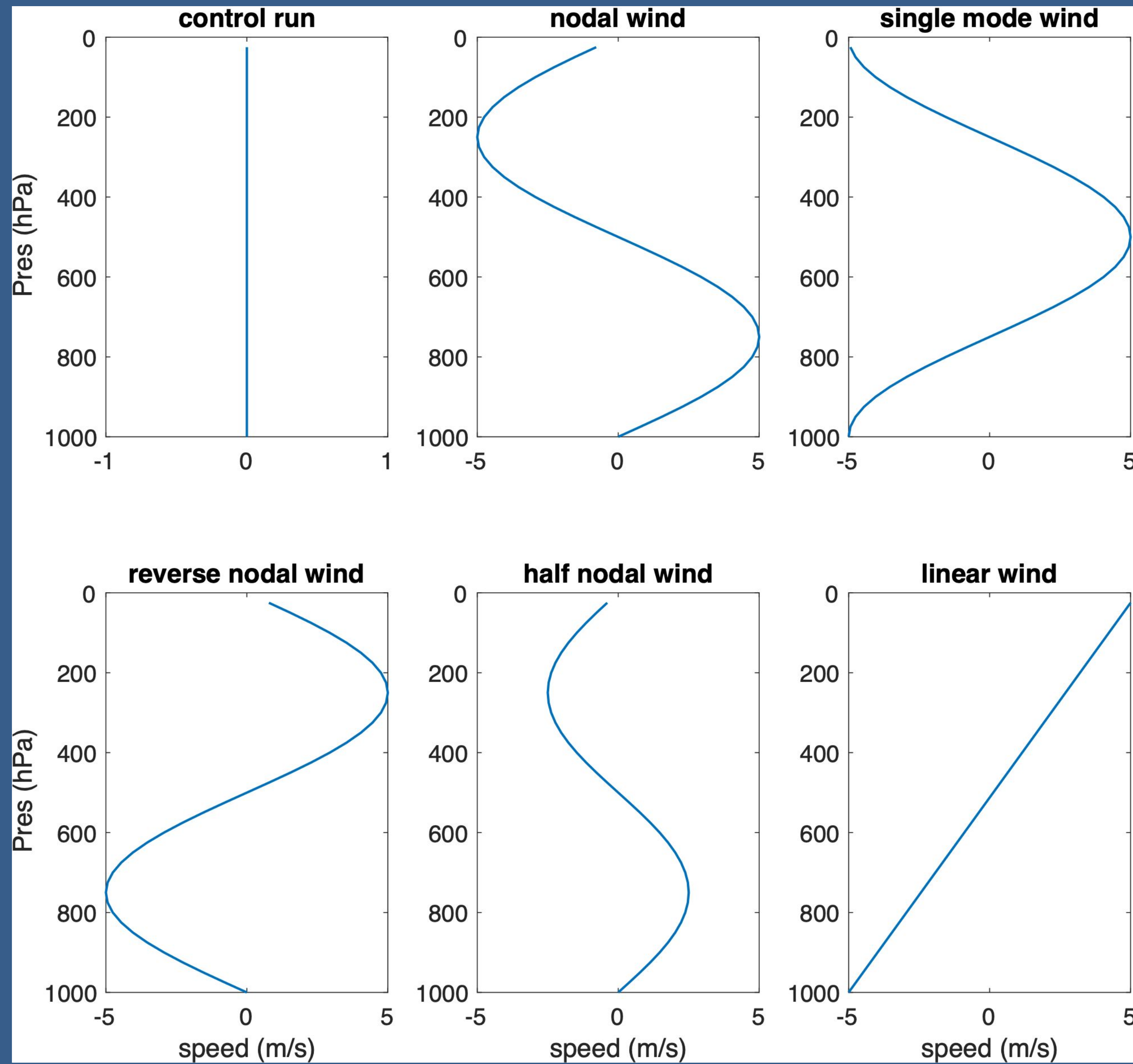


# Model setup

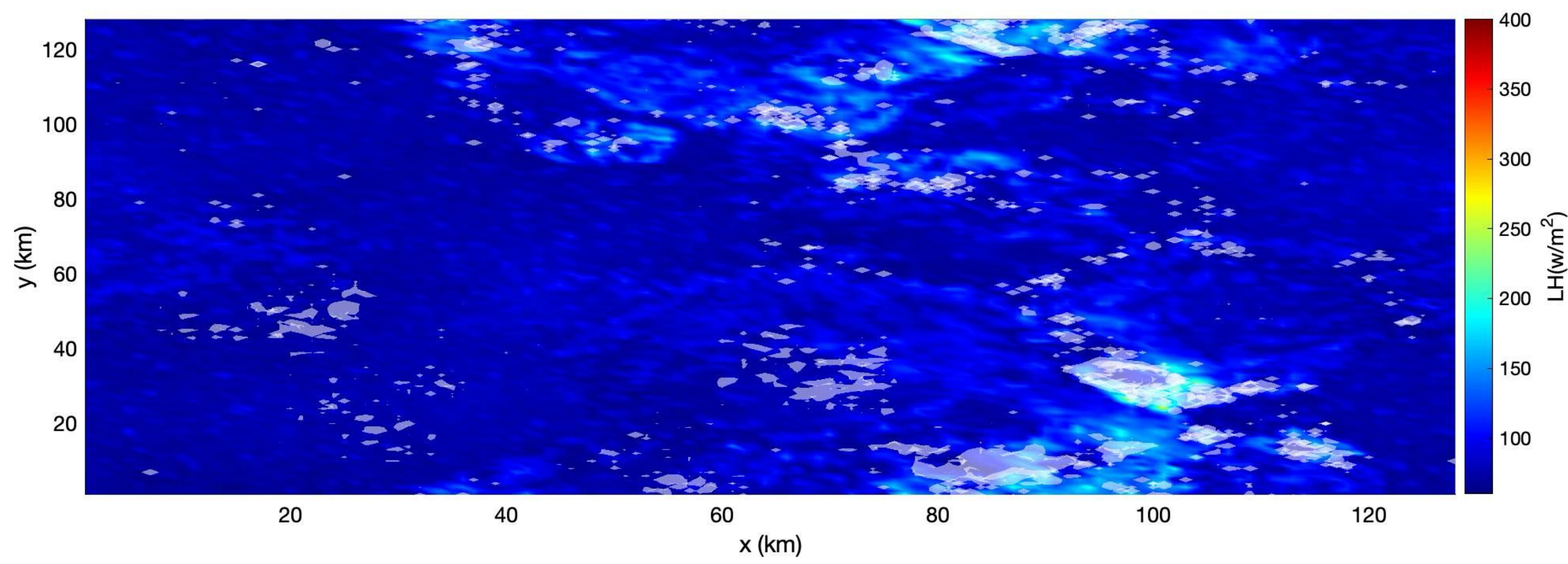
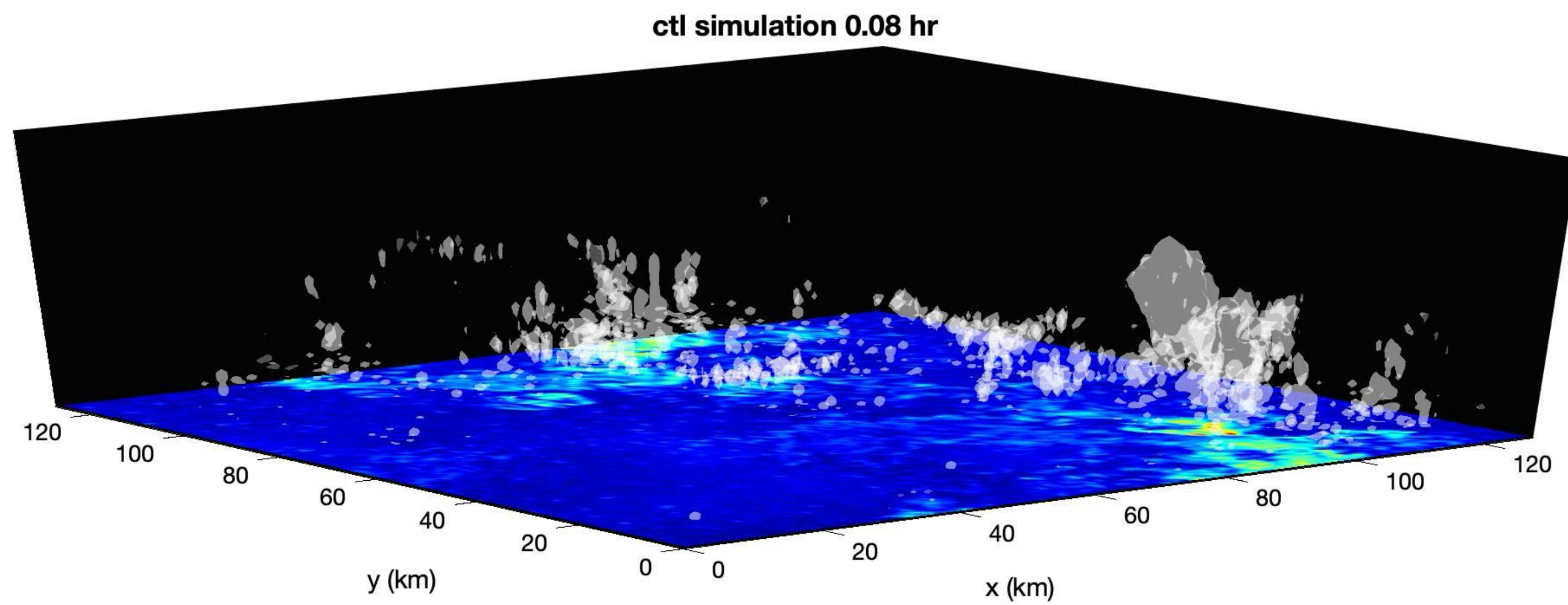
- Cloud-resolving model: SAM 6.11.6
- Radiative Convective Equilibrium (RCE) configurations, no large-scale forcing, except for the added zonal winds
- Model setup: 64 stretching vertical levels; 500m horizontal; 256 x 256 grid boxes; 10s temporal resolution
- After reaching RCE status on day 25, 20 ensemble members with each being run for 3 hours





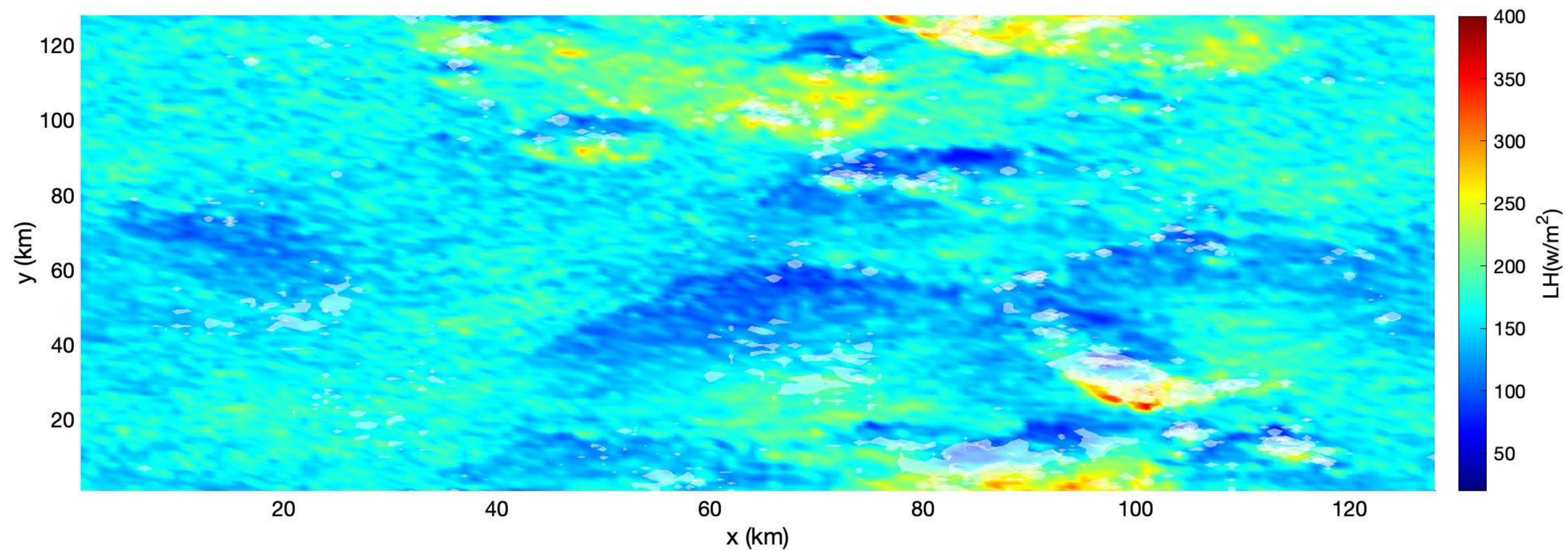
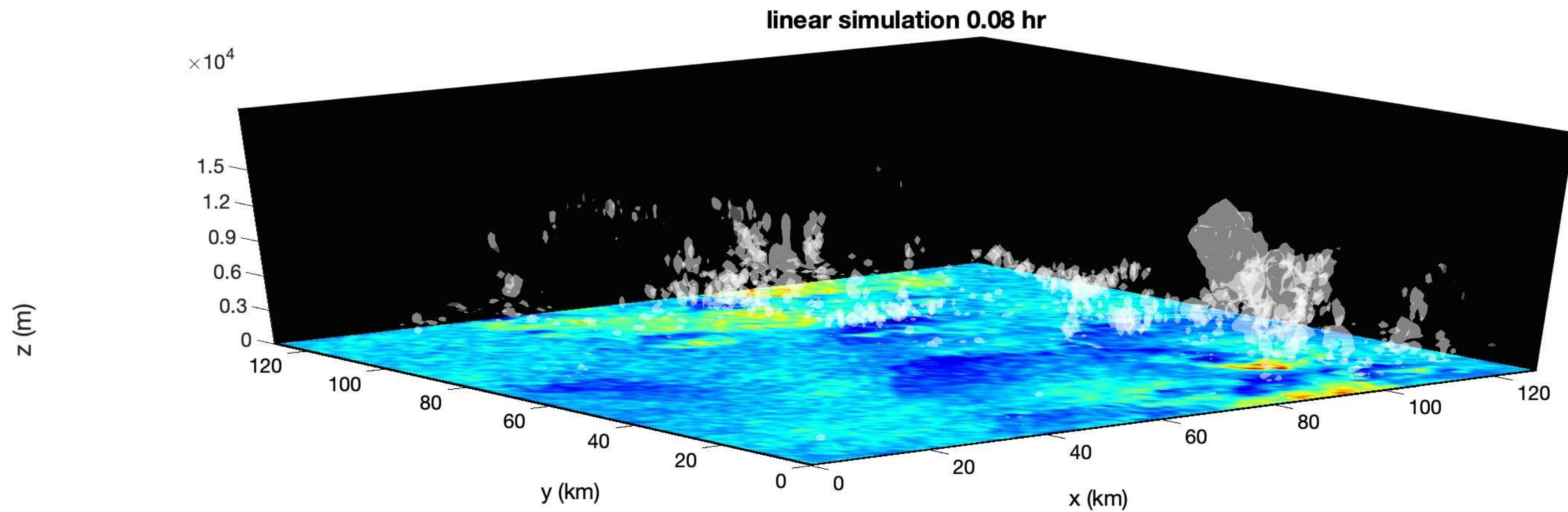






- For control scenario, 10 out of 20 ensemble members display organization (more circular)

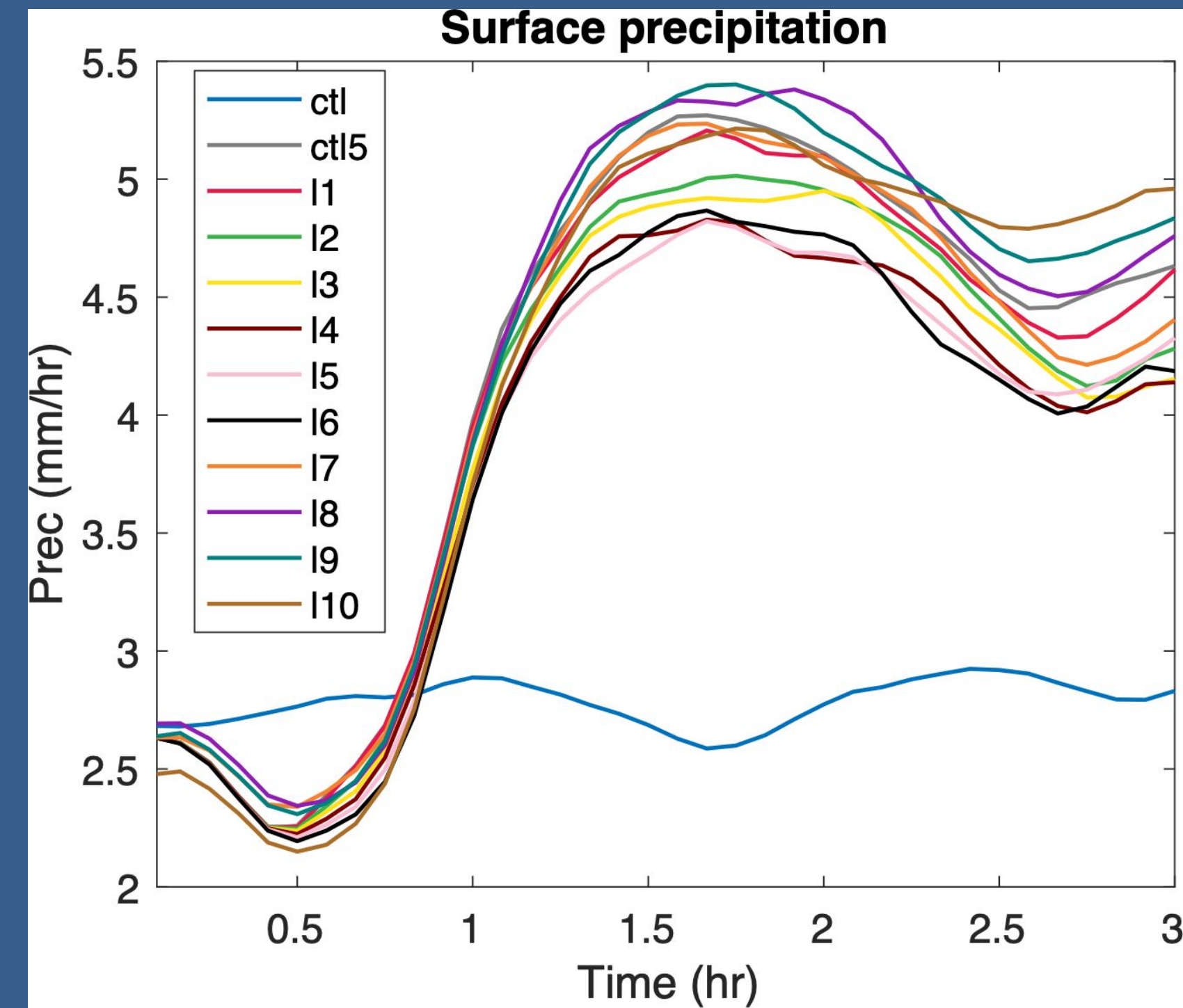
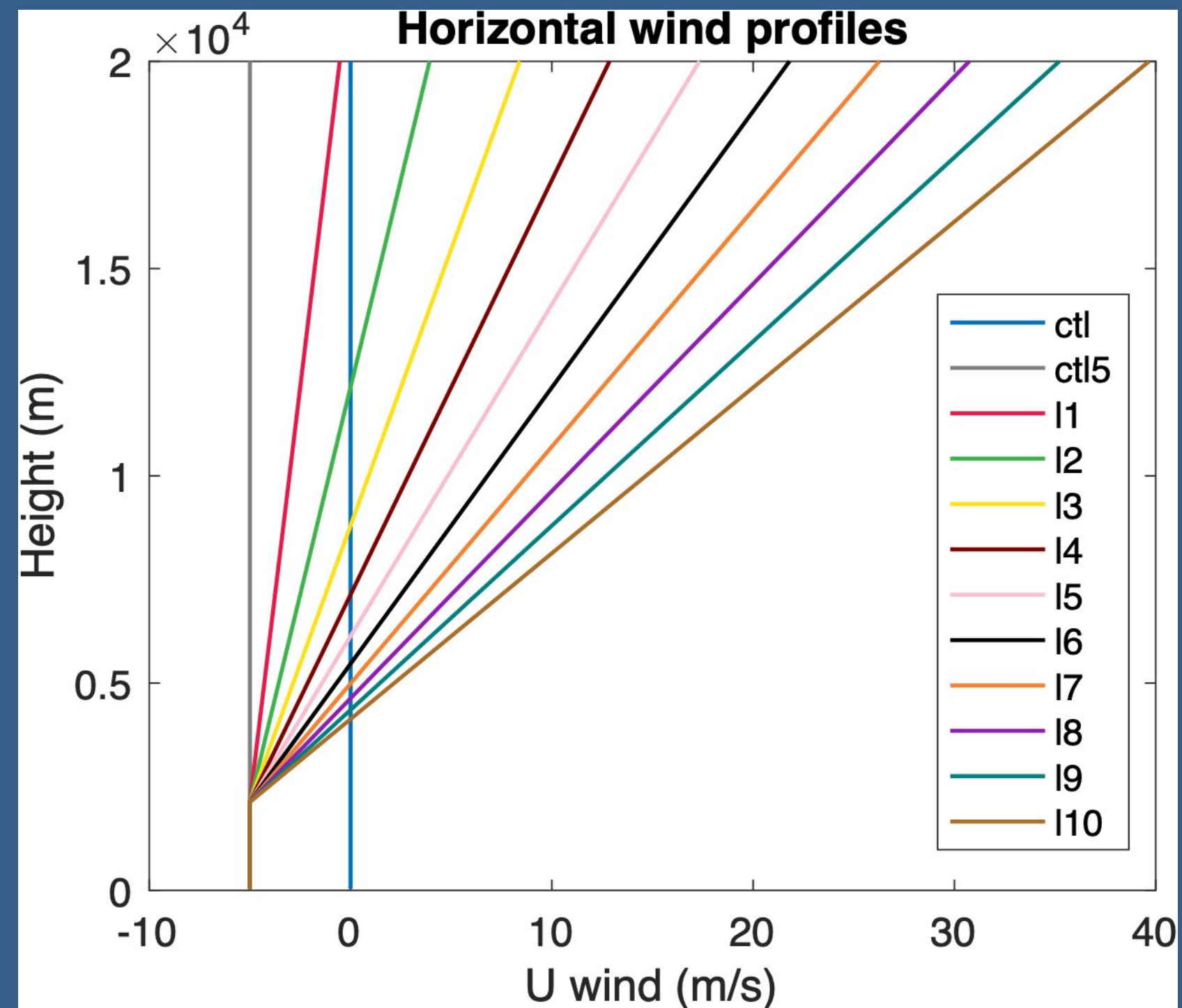




- For linear shear scenario, 15 out of 20 ensemble members display convective organization

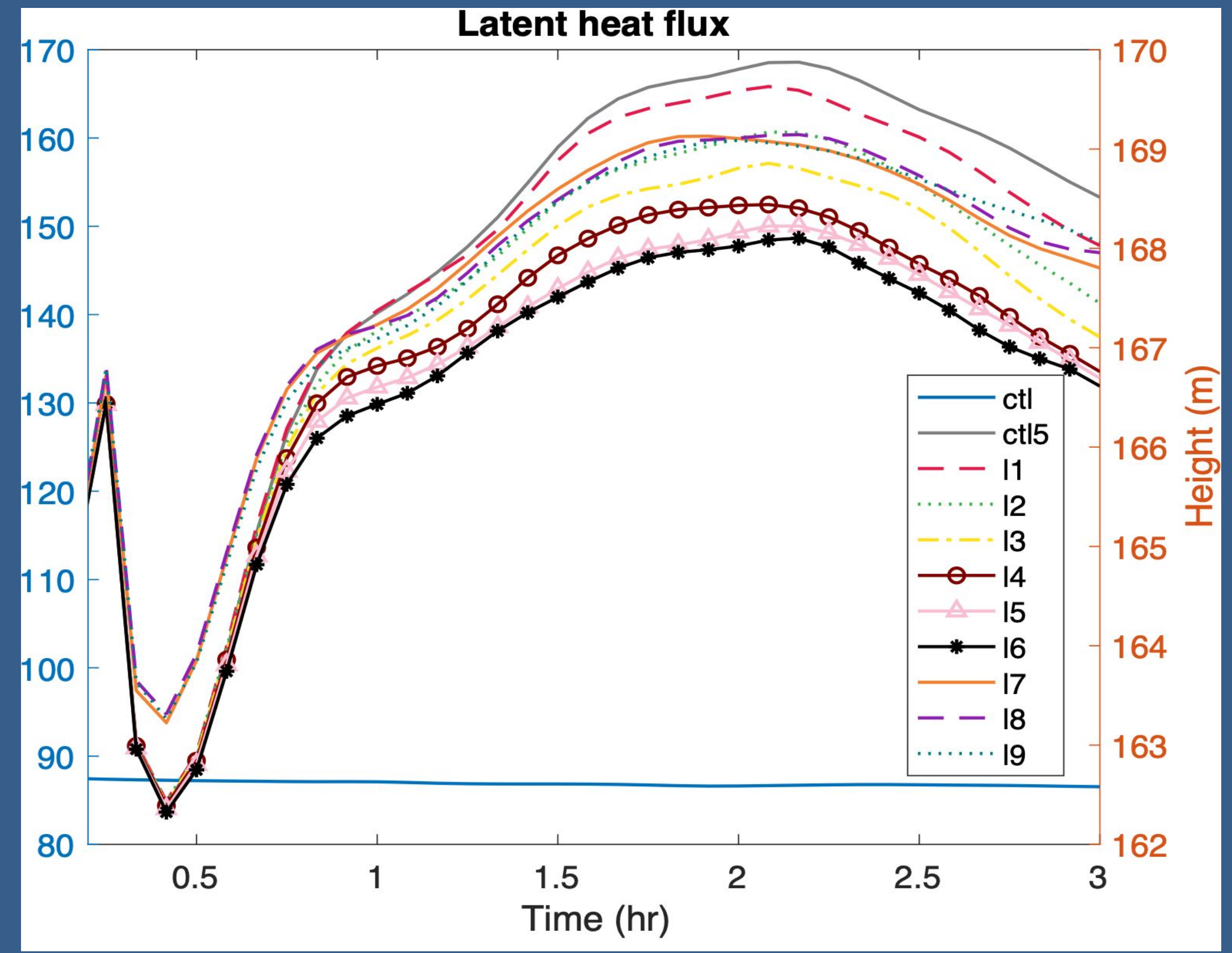
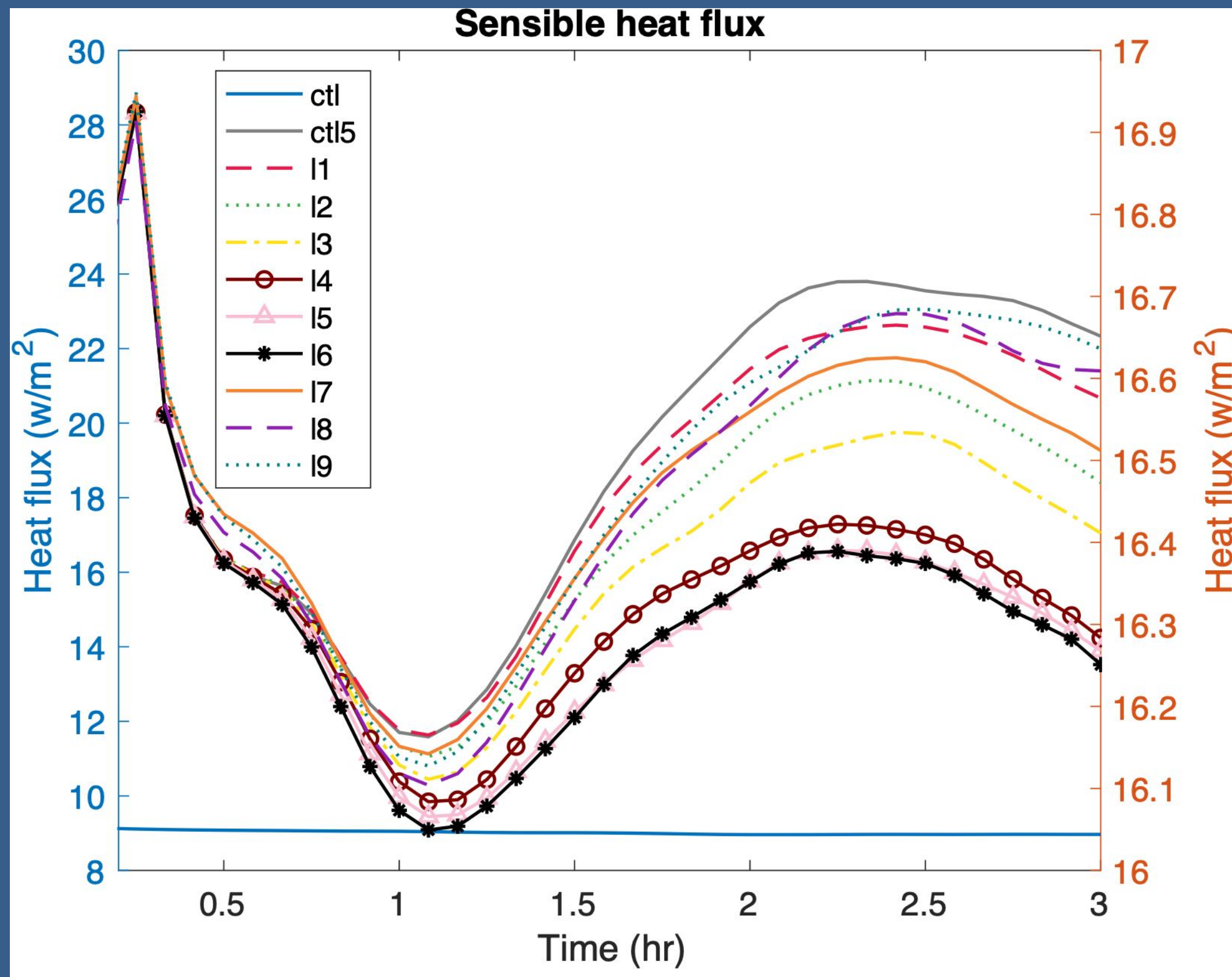


# Precipitation responses to various magnitudes of shear



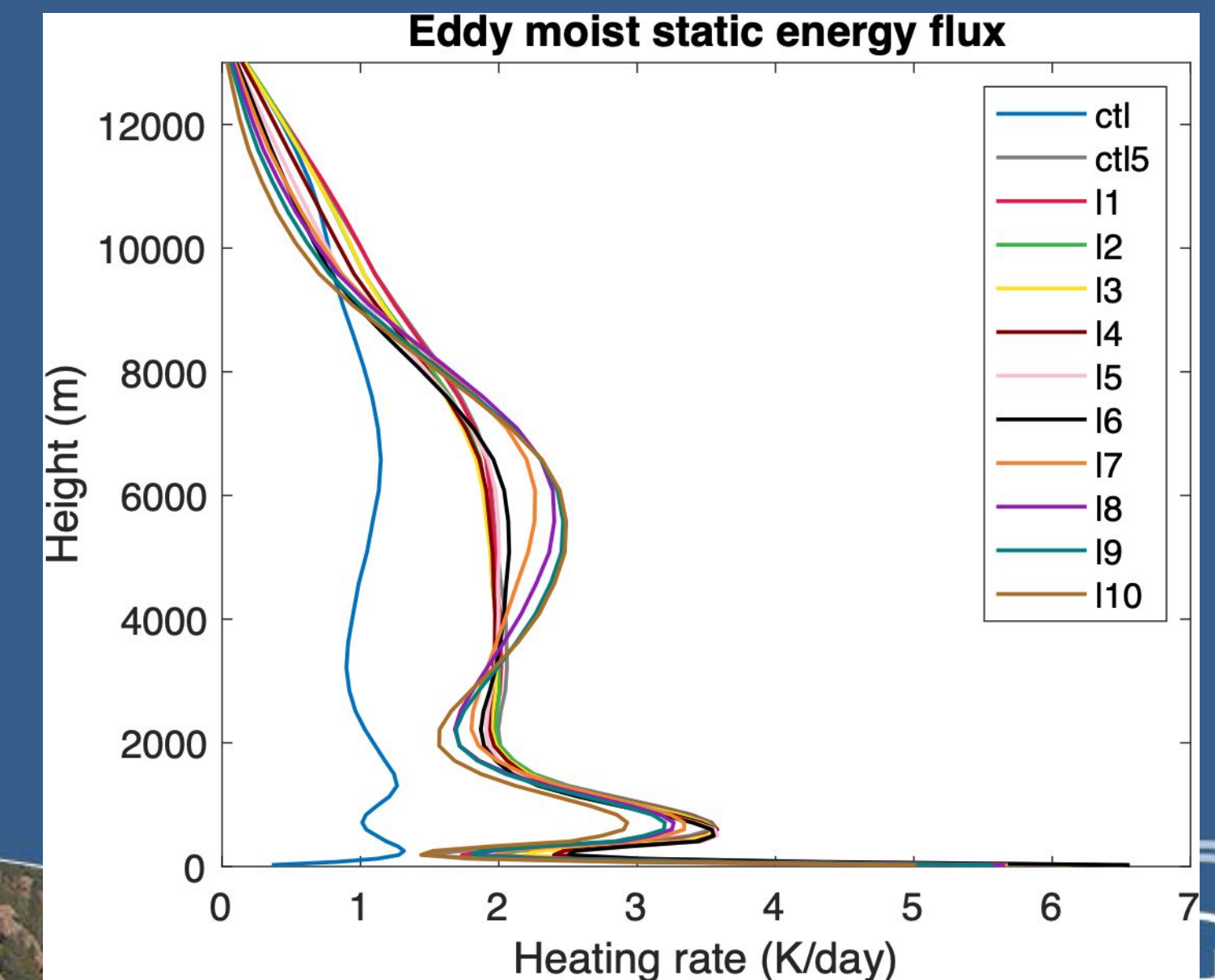
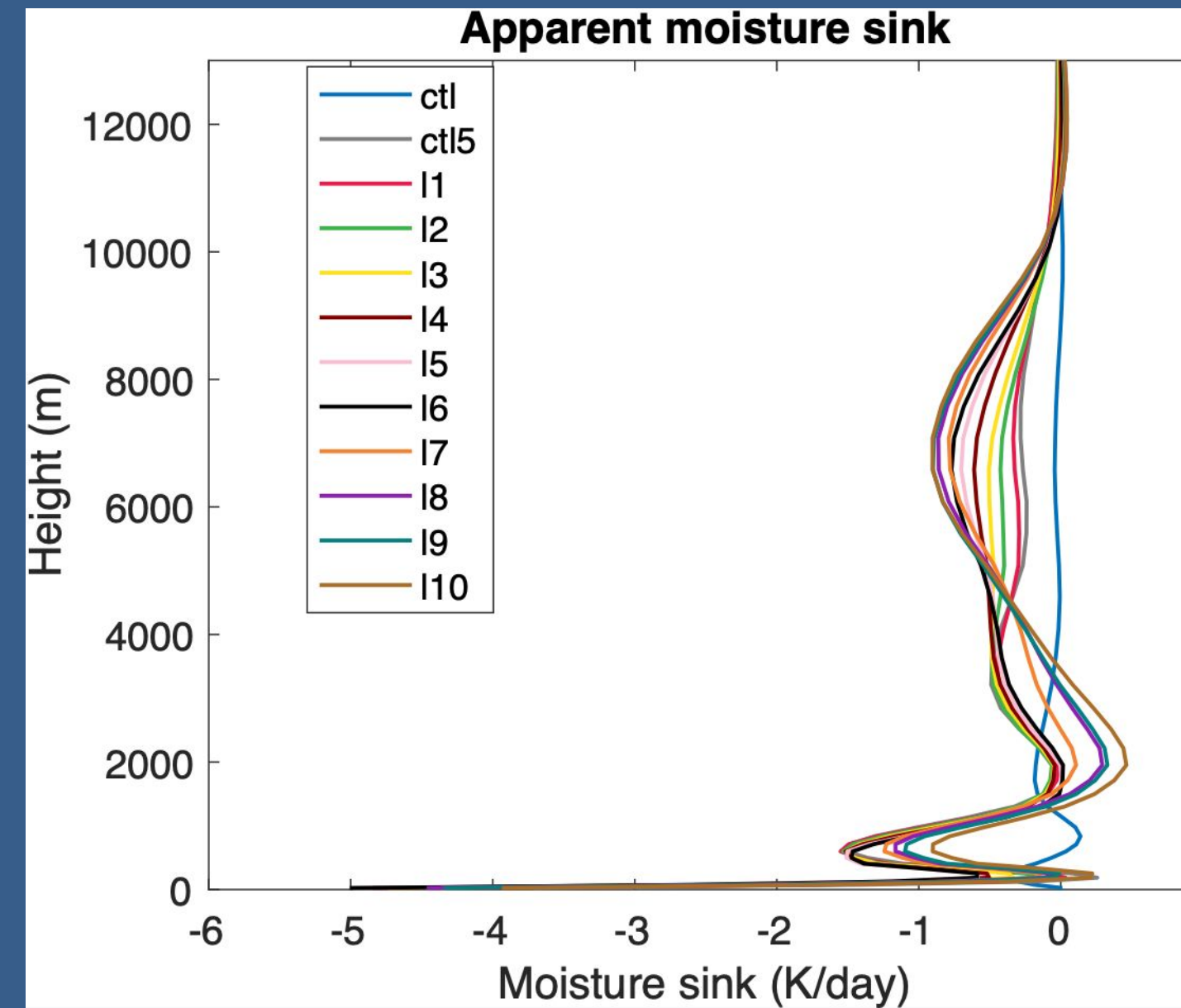
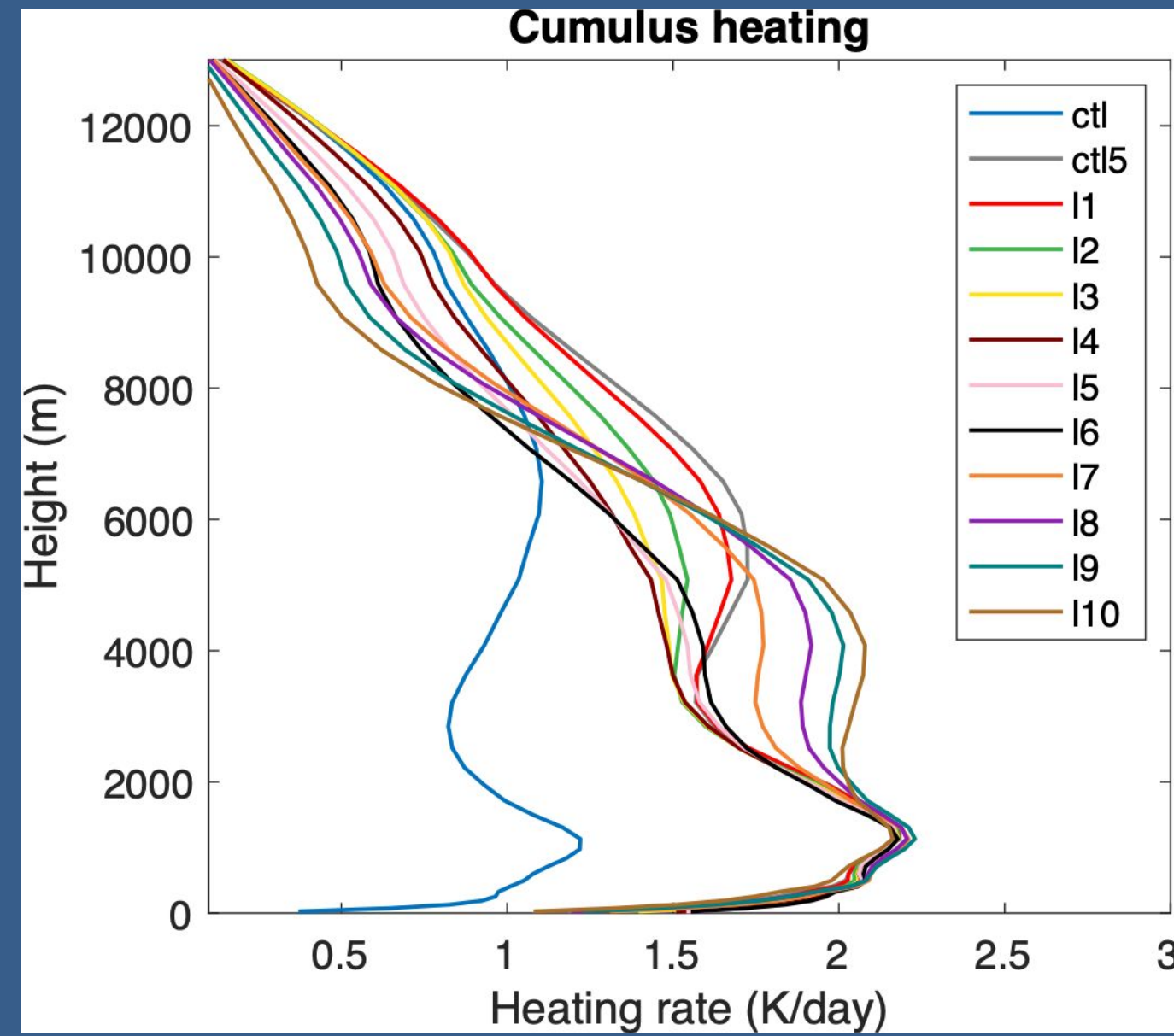


# Other responses to various magnitudes of shear



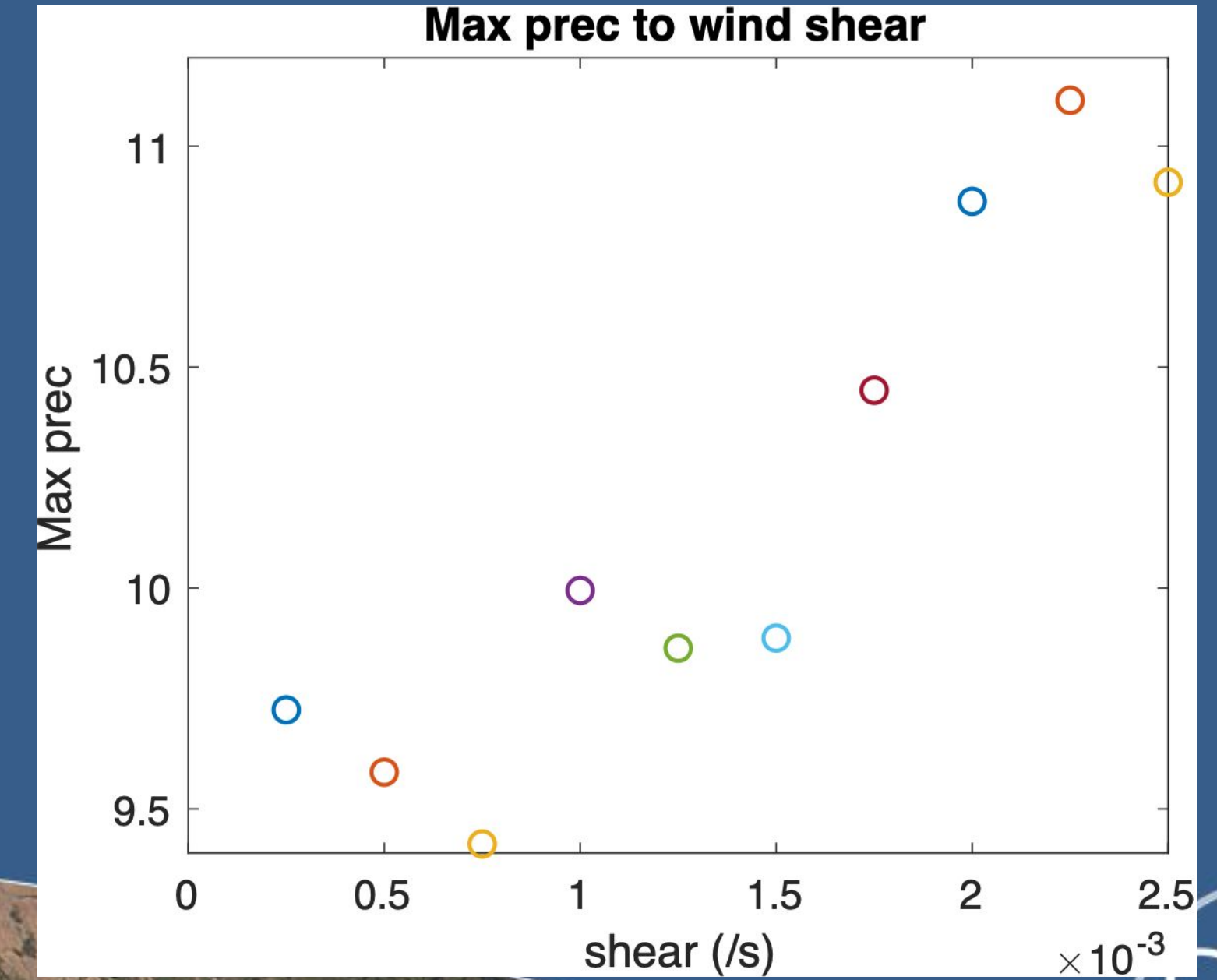
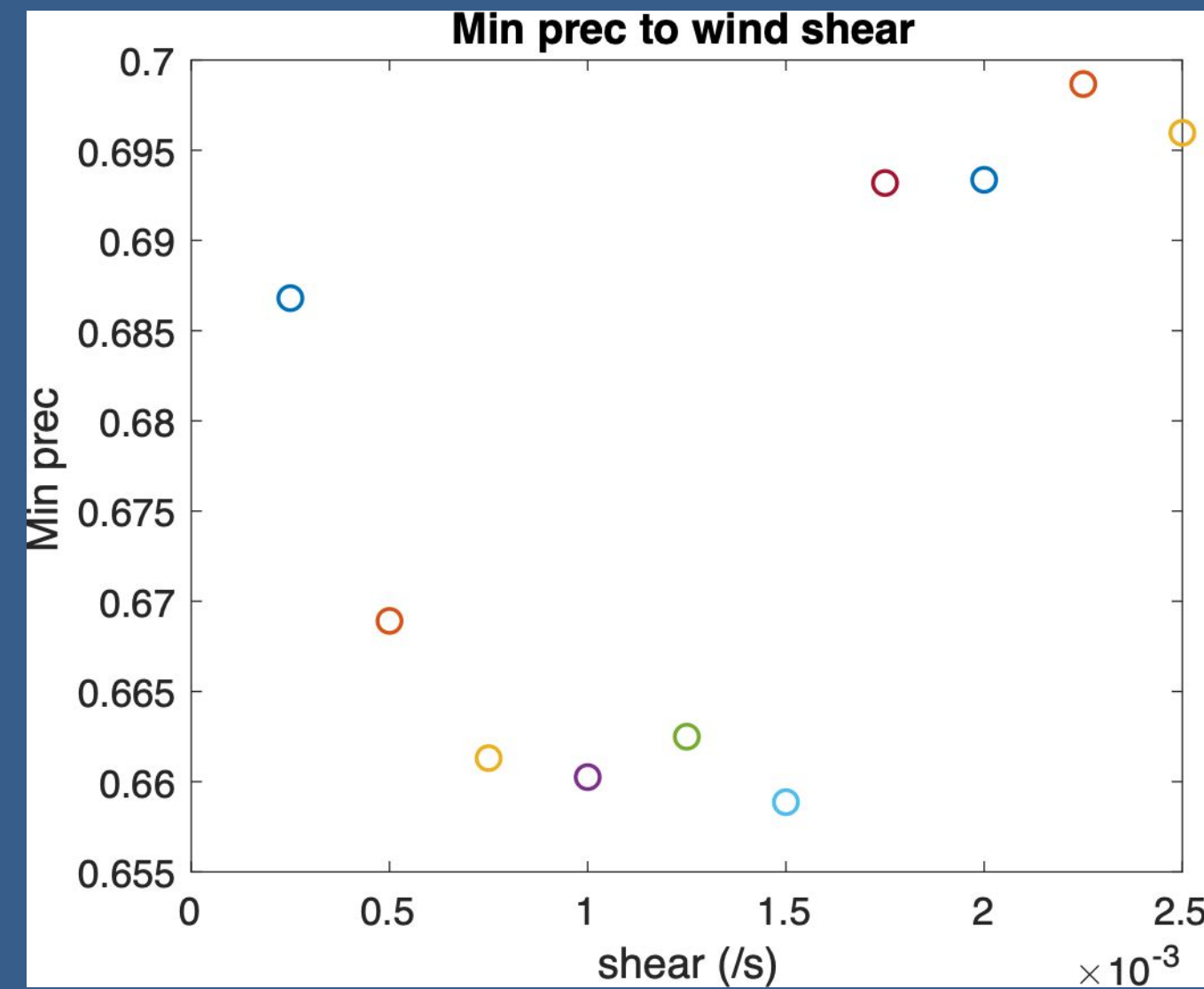
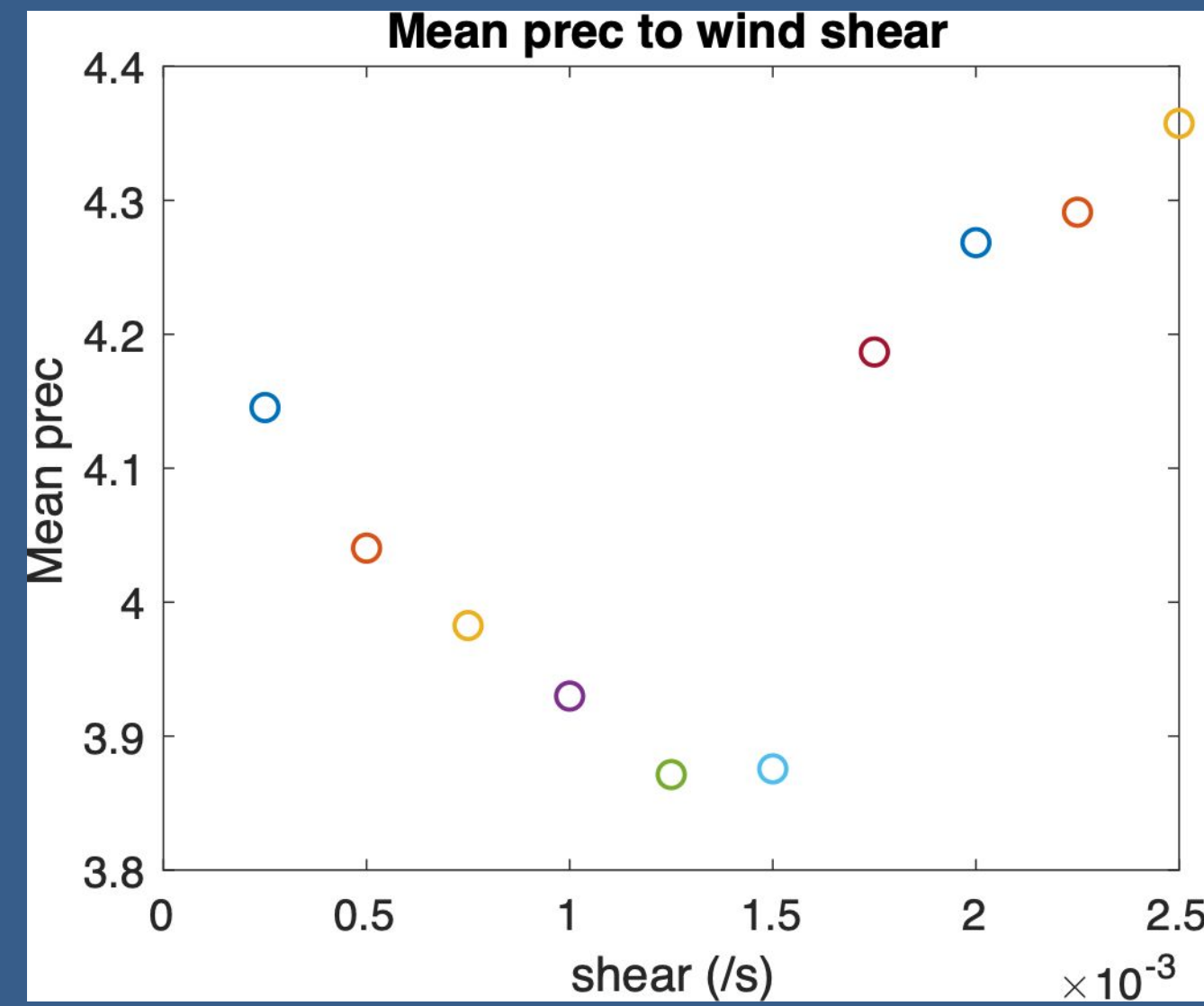


# Other convective responses to various magnitudes of shear





# Precipitation response to wind shear



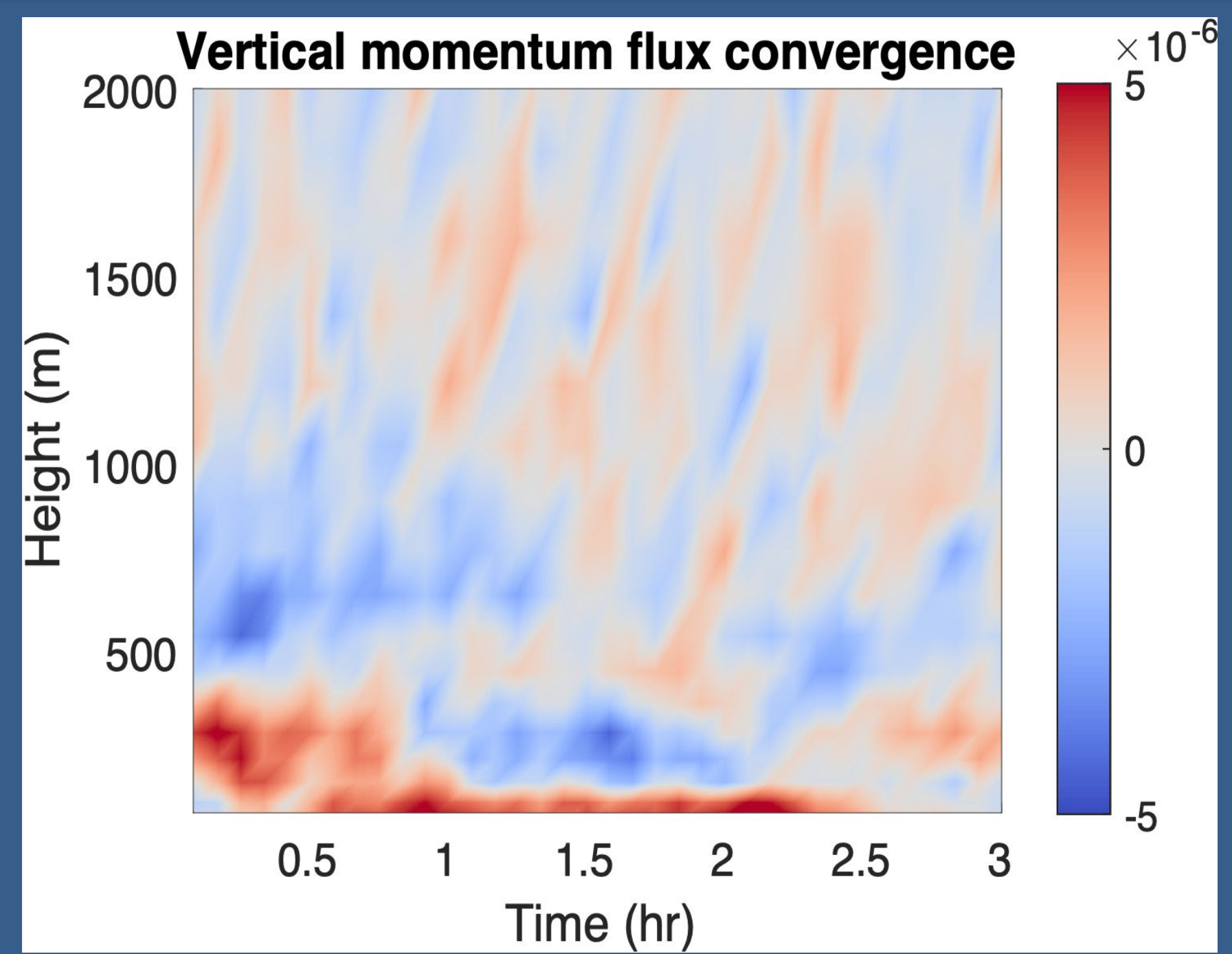
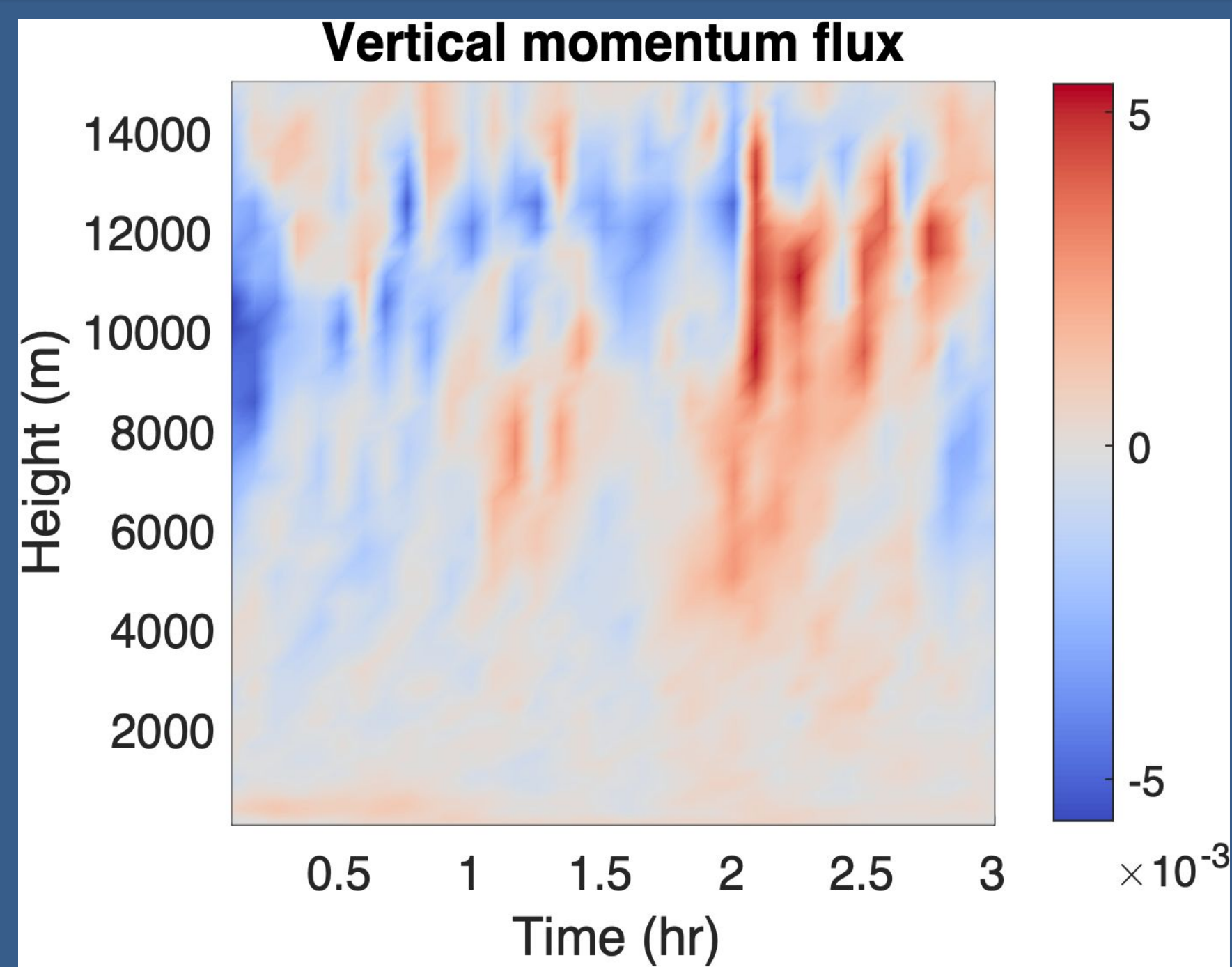


# Scientific questions

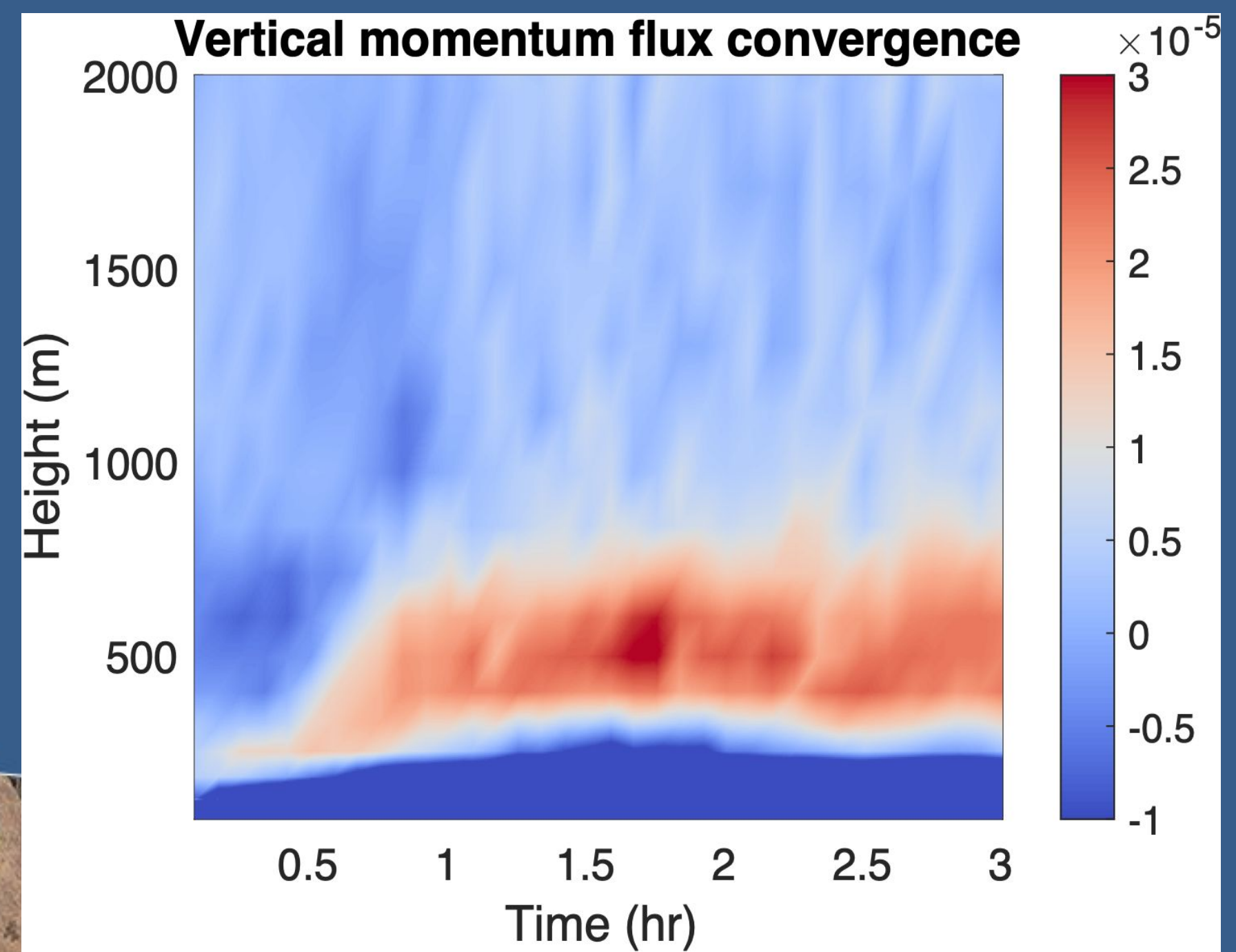
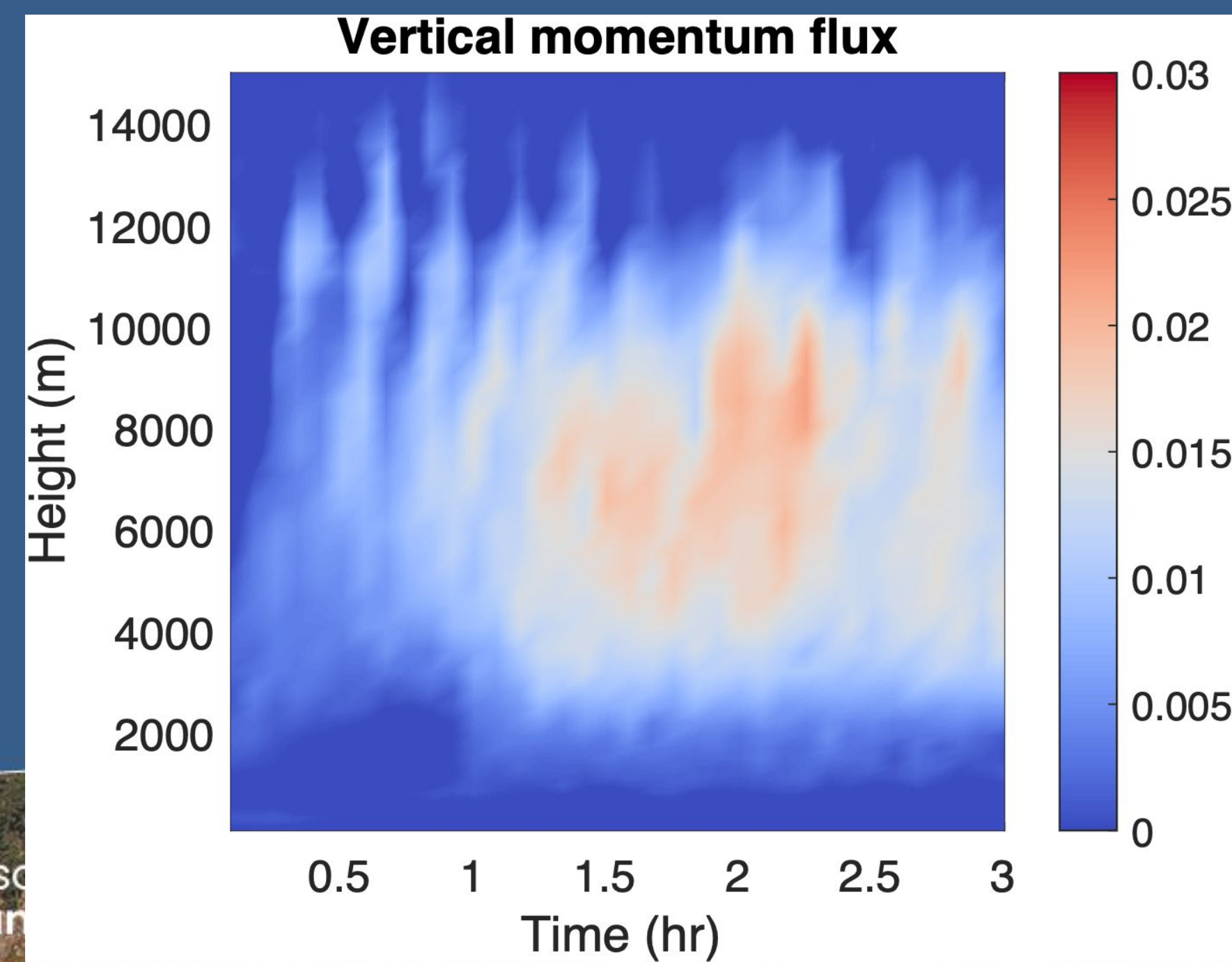
- Why is there surface precipitation suppression in the first hour of simulation?
- Why is the convective response to wind shear non-monotonic?
- What determines the critical wind shear value?



Ctl



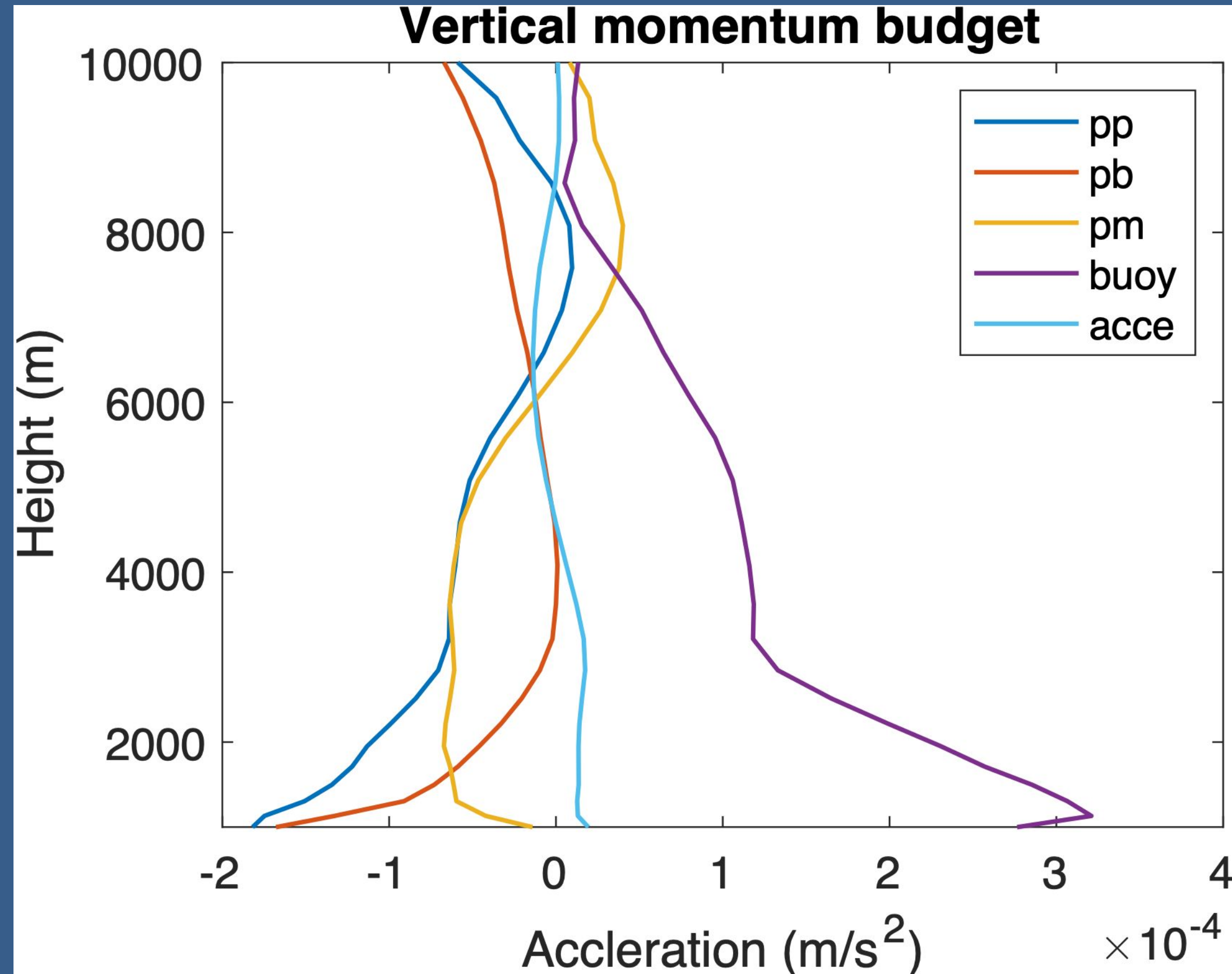
Ctl5





# Vertical momentum budget analysis

Cloudy updrafts buoyancy, PP, acceleration



$$\frac{1}{2} \frac{\partial \bar{w}^2}{\partial z} = \bar{B} - \frac{1}{\rho} \frac{d\bar{p}'_B}{dz} - \frac{1}{\rho} \frac{d\bar{p}'_M}{dz} - \epsilon \bar{w}^2$$

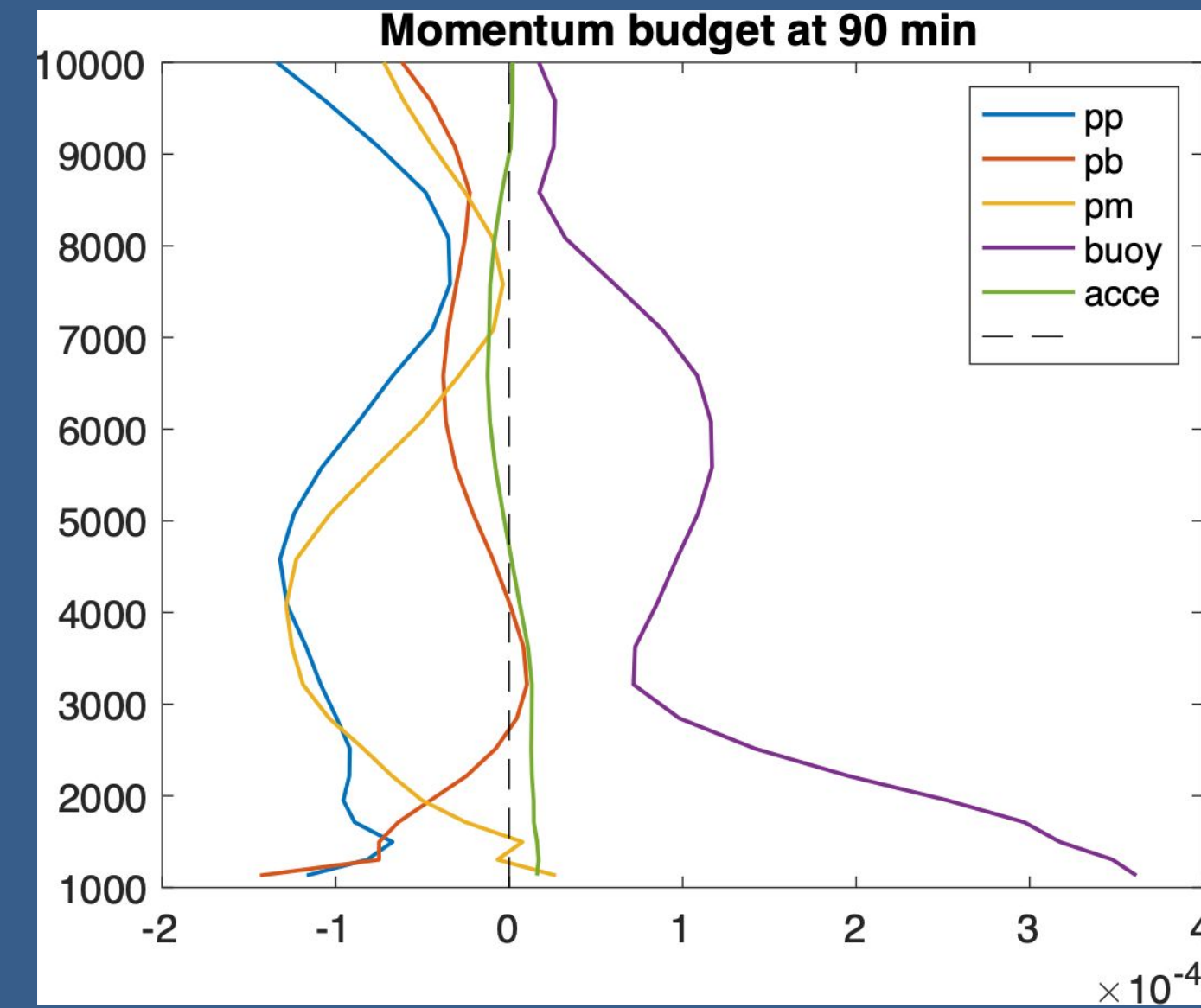
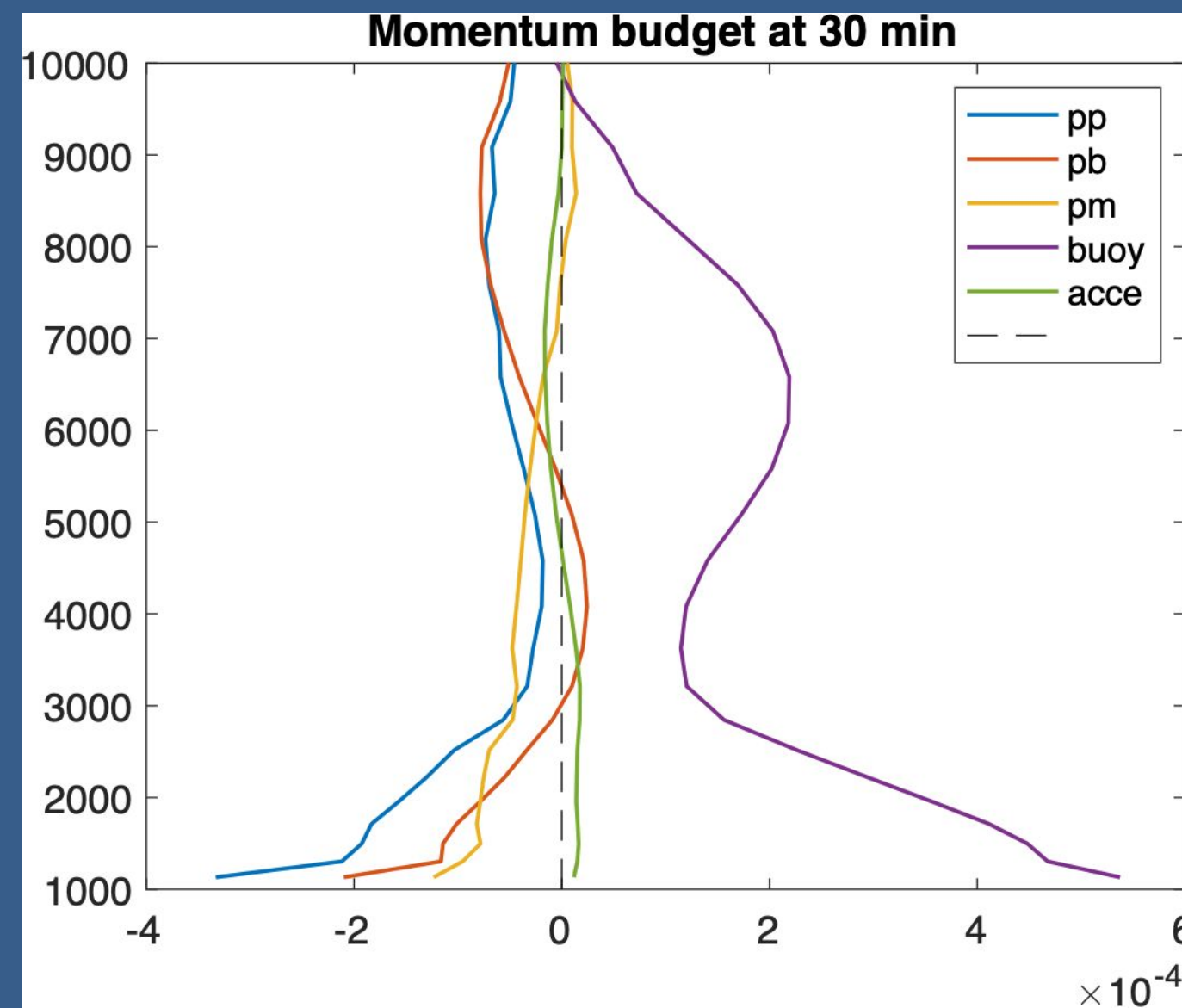
$$\nabla^2 p'_B = \partial_z(\rho_0 B),$$

$$\nabla^2 p'_M = -\nabla \cdot (\rho_0 \vec{v} \cdot \nabla \vec{v}).$$

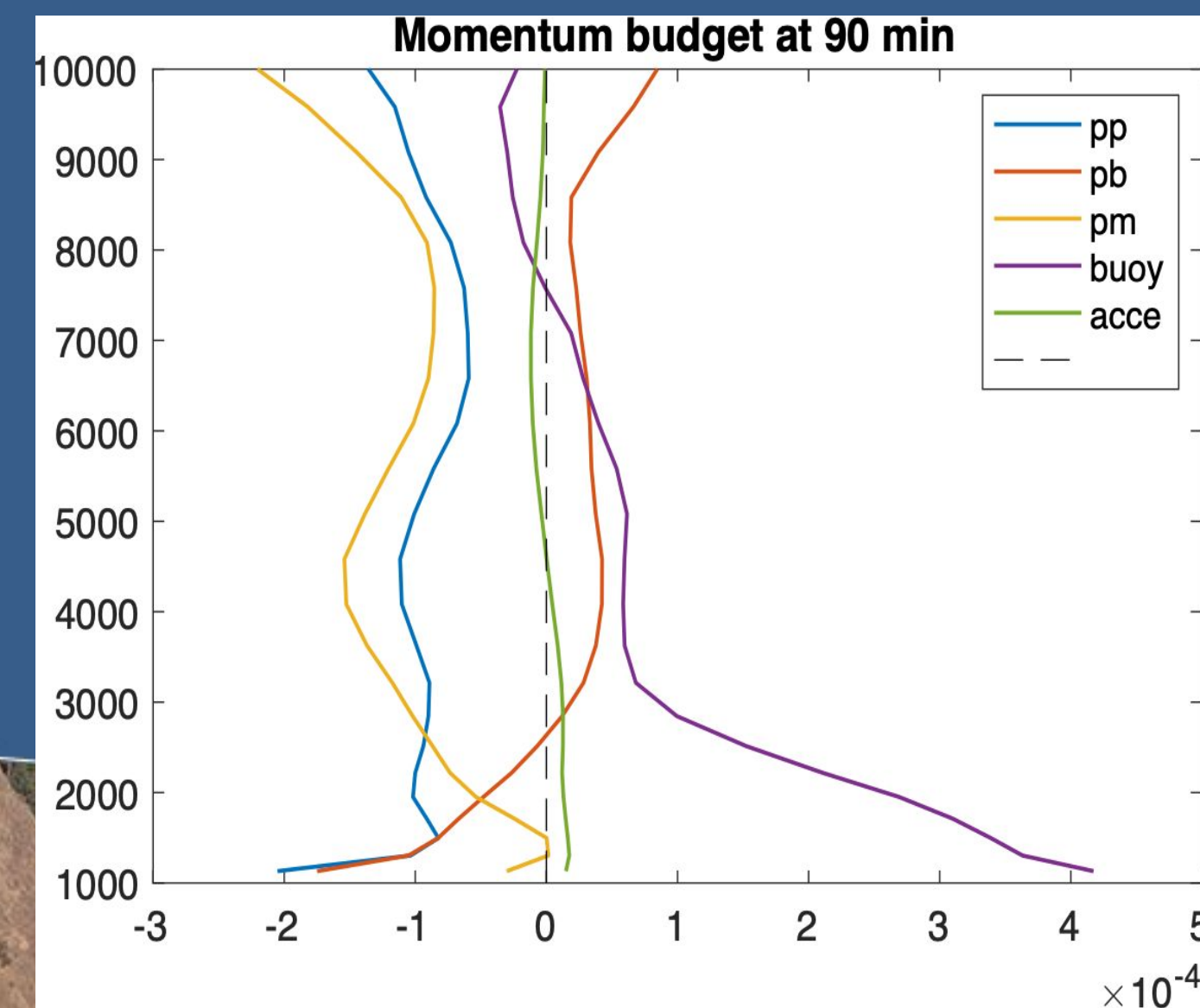
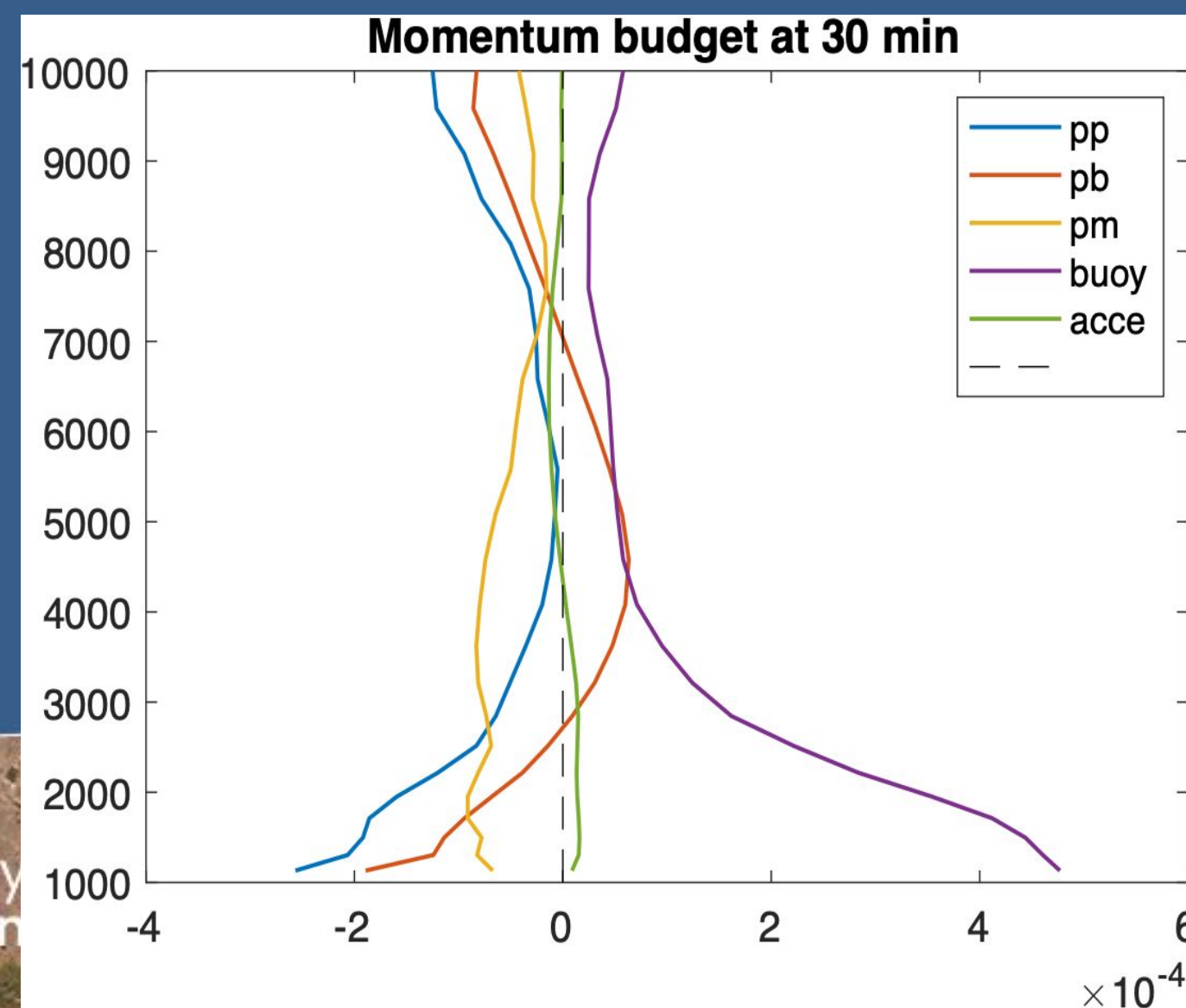


# Momentum budget temporal evolution of cloudy updrafts

CtI5



Linear6





# Take-home messages



Cumulus ensembles response to imposed wind shear in a non-monotonic behavior



A physically- and process- based convective trigger could be implemented to better characterize MCSs in coarse-resolution climate models.



A competing mechanism between pressure drag and surface fluxes. The immediate reduction can be related to pressure drag, then it takes about an hour for a convective cell to reach to the upper troposphere, that's where the surface fluxes kick in.





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