North Atlantic Ocean Response to NAO Surface Heat Flux in Three Climate Models

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Mechanisms of decadal SPNA Variability







2023 CESM Workshop

Weak NAO-AMOC Relationship in Coupled Models

- Different representation of surface buoyancy (heat) fluxes associated with NAO
- Different efficacy of NAO buoyancy forcing for driving ocean response due to different mean states



CMIP6 piControl



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CMIP6 piControl





Allowing for assessing the robust responses and the differences arising from different background states

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Experimen

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- Obtained by regressing ERA5 surface heat flux onto an observed DJFM NAO index
- ±2σ applied over the SPNA for 10 years (winter only); run for additional 10-20 years without the forcing

Focus: link between surface water-mass transformation (WMT), AMOC and SPNA upper ocean temperature responses (ensemble mean differences)

WMT Response





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- **AWMT** directly induced by the forcing is relatively small
- ΔQ_c itself is small $\rightarrow \Delta WMT$ is largely induced by ΔA_c



Outcropping Area Response and Heat Flux



- ΔA_{o} mirrors the ΔWMT pattern
- ΔA_{ρ} is exposed to the background surface heat flux, which is stronger in CESM2 than in the other two models

Outcropping Area Response and Heat Flux Feedback



- Q' initially cools and makes the surface dense
- $A(\sigma'_2) > A(\sigma_2)$ $\rightarrow \sigma_2$ exposed to more Q +Q'
- σ_3 exposed to Q + Q' (WMT=0 before Q')
- More exposure to Q further expands $A(\sigma'_{2,3})$
- Because Q is larger in CESM2, ∆WMT is also larger





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AMOC (σ) Response

(c)

(f)

(h)

60



- Spin-up of the lower (denser) AMOC(σ) limb in the subpolar latitudes (first decade)
- Lower limb anomalies ~2x larger in CESM2, consistent with the WMT anomalies
- No significant anomalies in the upper (lighter) limb
- Southward propagation of the lower limb anomalies (second and third decades)
- Development of the upper limb anomalies

AMOC (σ)

ASSH/dx & AVVEL (45°N) TISE





- Anomalous dense water advects southward and is accumulate near the gyre boundary west of the Mid-Atlantic Ridge (MAR)
- Generating zonal SSH gradient
- Driving anomalous meridional flow (anomalous NAC)
- Bringing warm and salty subtropical waters into the SPNA (for +NAO)
- This mechanism working for all three models and consistent with that found by Yeager (2020) and Yeager et al. (2021) from both low- and hi-res CESM1

Upper Ocean Temperature Response



- Initial cooling due to direct forcing effect
- Delayed warming in the SPNA
- Dipole pattern with a anomaly of opposite sign off the Grand Banks (AMOC fingerprint)
- Warming penetrating into the Nordic Seas (third decade) → sea-ice response
- Patterns strikingly similar across the models
- Very similar patterns for the upper ocean salinity

Summar

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In response to the NAO surface heat flux forcing identically imposed in three CMIP6-class models, we found:

- Consistent mechanisms and patterns of the North Atlantic Ocean response (dense-water formation → AMOC → heat content in the SPNA)
- **Different amplitude** of the response
- Anomalous dense-water formation mainly occurs in the western SPNA
- Changes in isopycnal outcropping area and associated exposure to the background surface heat fluxes are the key for the ocean response
- Weak response directly driven by the imposed forcing
- The different background state can explain the inter-model amplitude difference
- Delayed SPNA warming due to a slow advection of anomalous dense waters and associated adjustment of the upper AMOC(σ)

Background WMT & AMOC





-4.8 -3.6 -2.4 -1.2 0 1.2 2.4 3.6 4.8 $\times 10^{-6} [kg m^{-2} s^{-1}]$





Δ WMT & Δ WMF for the First Decade







∆AMOC and SPNA UOT Time Series









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